

Ocean Data Assimilation for Seasonal Forecasting

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Troccoli, Jerome Vialard, ...

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Outline

- Why Ocean Data Assimilation?
- The Operational Ocean Analysis System
 - Main ingredients
 - Systematic error
 - Impact of data assimilation
- What next?
 - Future plans for seasonal forecasts
 - The challenge of decadal variability (ENACT)
 - Ocean analysis for shorter range forecasts (?)
- Summary

The Basis for Seasonal Forecasts

- *Atmospheric point of view: Boundary condition problem*

- Forcing by lower boundary conditions changes the PDF of the atmospheric attractor

“Loaded dice”

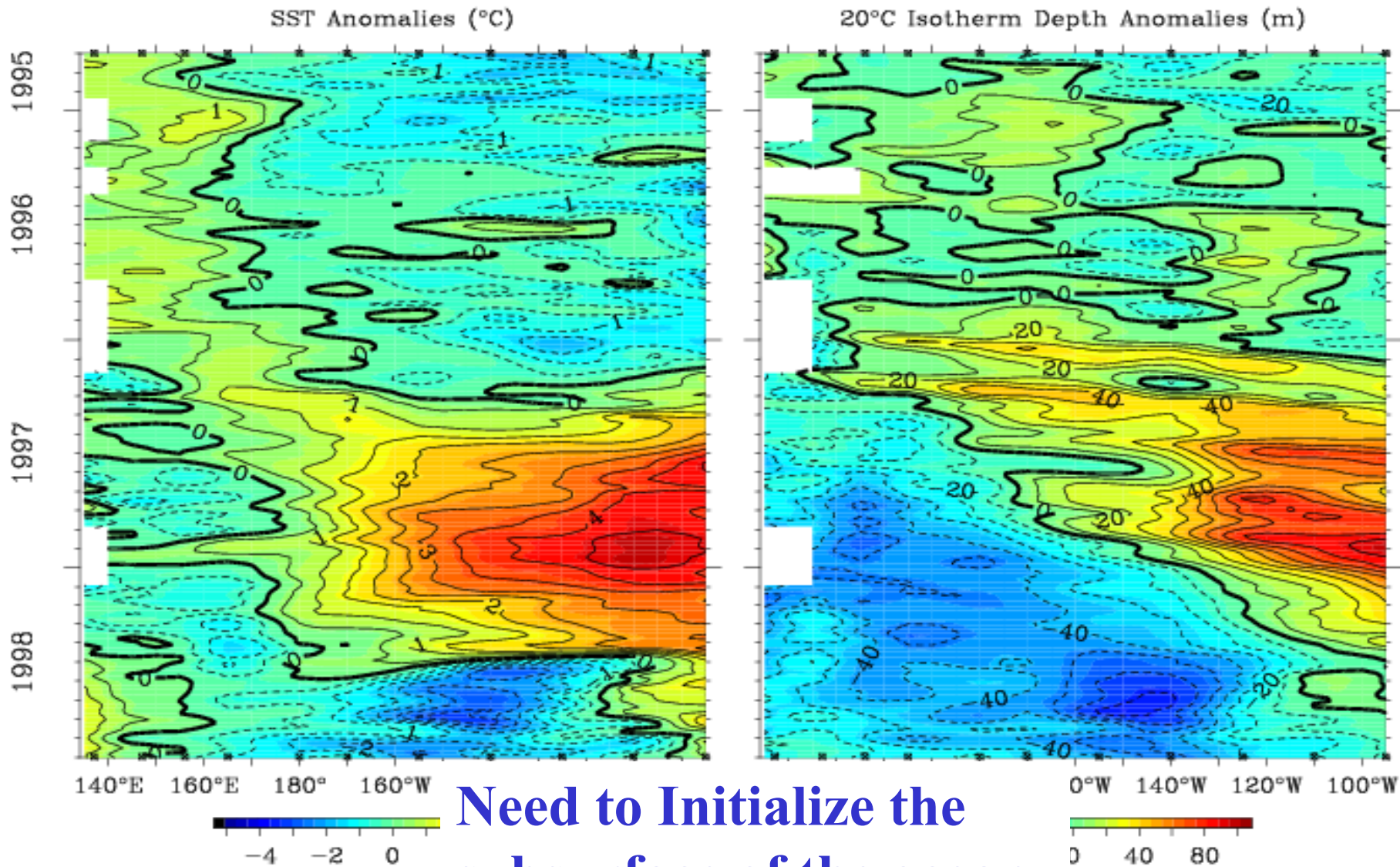
- The lower boundary conditions (SST, land) have longer memory

- Higher heat capacity (Thermodynamic argument)
- Predictable dynamics

- *Oceanic point of view: Initial value problem*

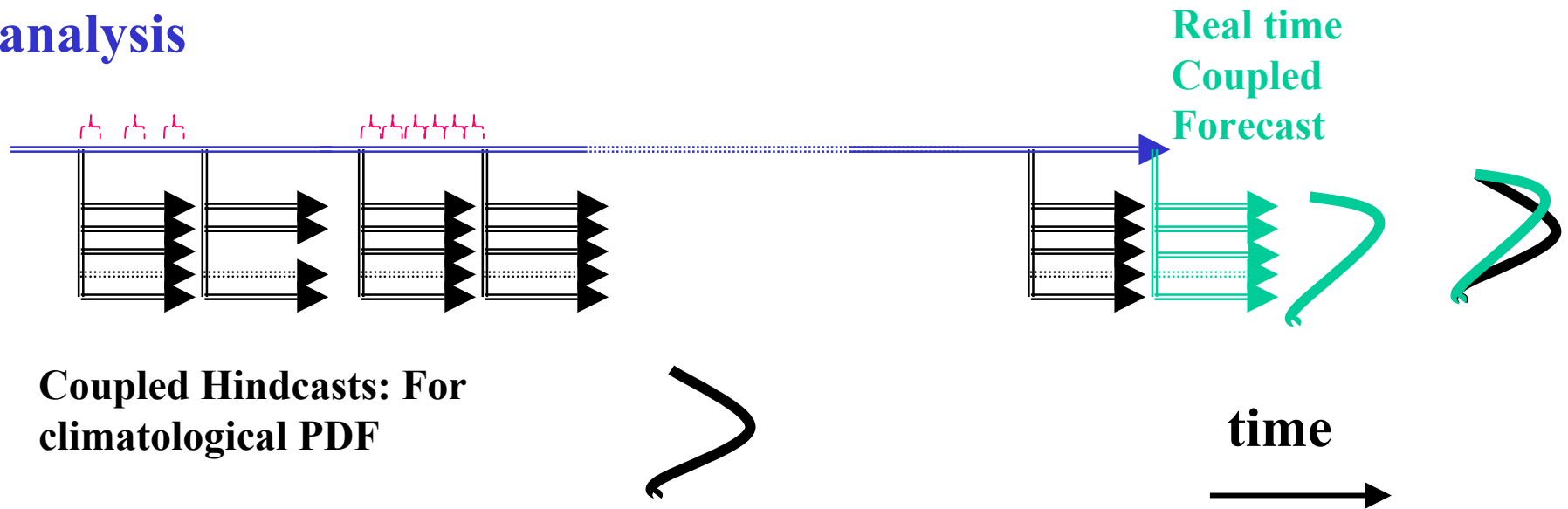
- *Prediction of tropical SST*

Five-Day SST and 20°C Isotherm Depth 2°S to 2°N Average



Need to Initialize the subsurface of the ocean

Ocean analysis



The seasonal forecasts need the model climatological PDF

=>Need for several years of ocean (re-)analysis

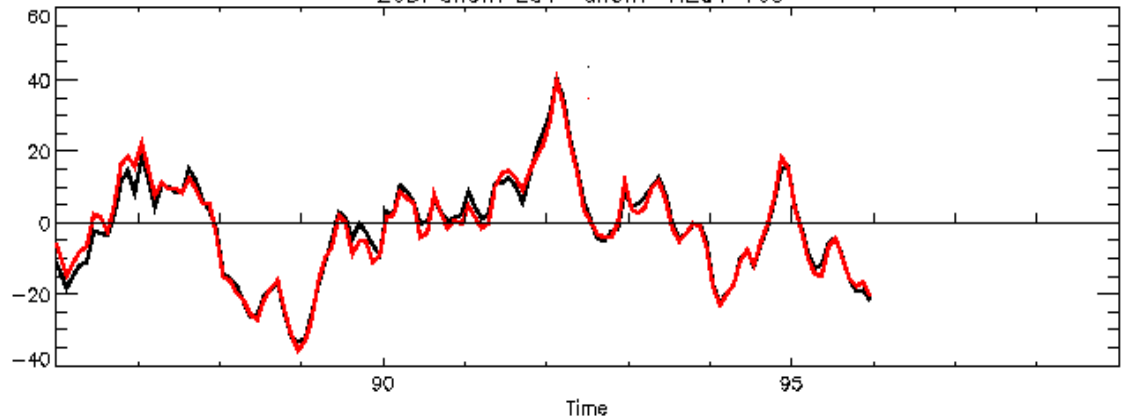
Creation of Ocean Initial conditions:

- A) Ocean model driven by surface fluxes :
 - Momentum,
 - Heat (short and long wave)
 - PME
- B) Daily Atmospheric Fluxes
- C) Real time: from the NWP
- D) Calibration Period: Reanalysis products.

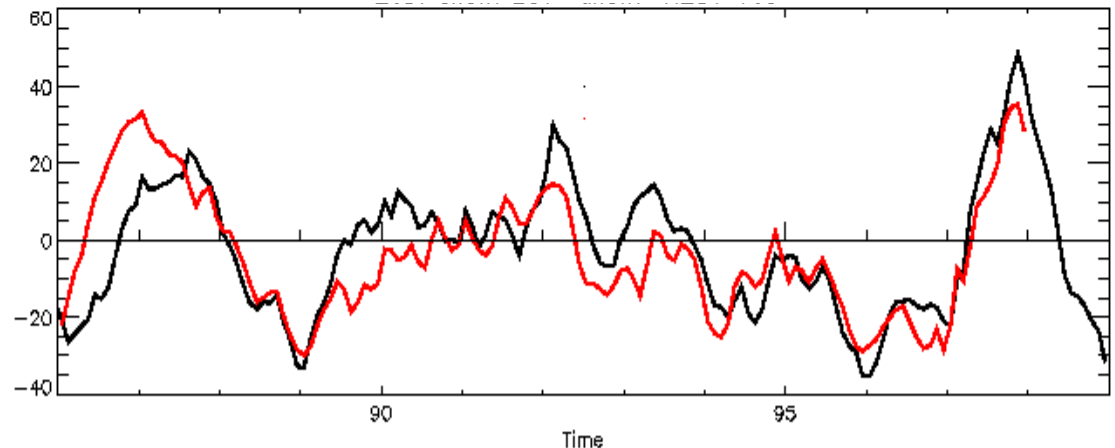
Oceanic Initial conditions

- A) The state of the upper ocean is largely determined by the wind forcing
- B) Large uncertainty in wind products.
- D) Need Data assimilation to constrain the ocean state.

D20 anomaly (with data assimilation)



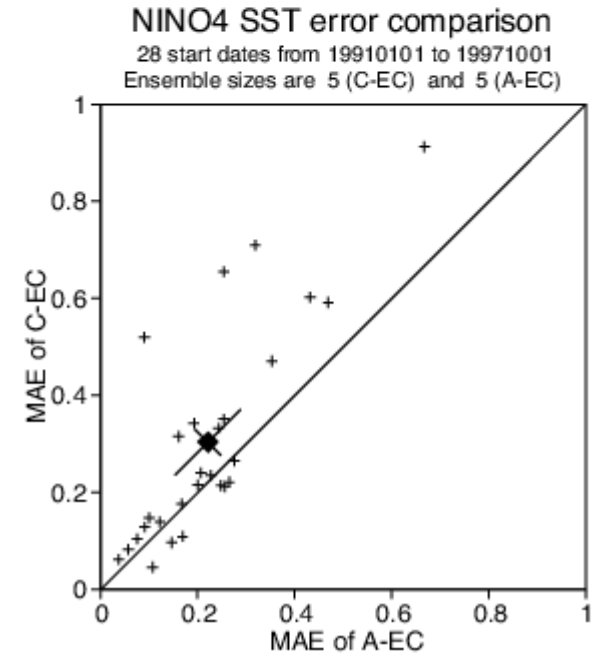
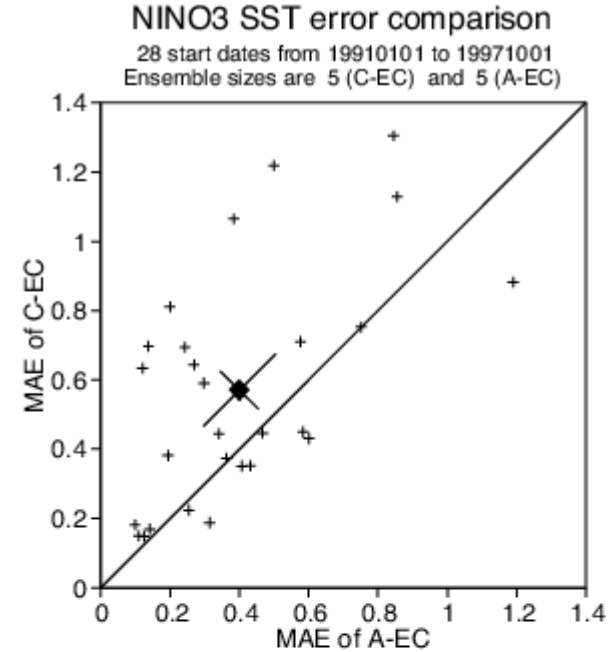
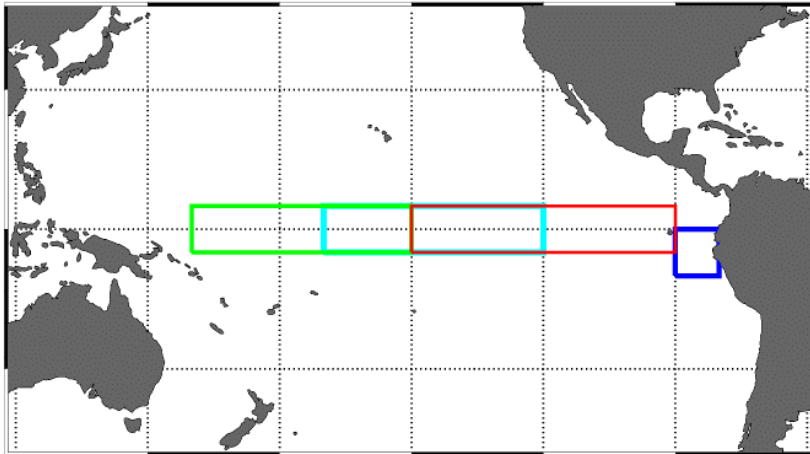
D20 (thermocline depth) anomaly (Eastern Pacific)



Ocean data assimilation also improves the forecast skill

(Alves et al 2003)

Nino3.4, Lon = [-170, -120], Lat = [-5, 5]
Nino12, Lon = [-90, -80], Lat = [-10, 0]
Nino4, Lon = [160, -150], Lat = [-5, 5]
Nino3, Lon = [-150, -90], Lat = [-5, 5]



January 1997: Seasonal Forecast System 1

- The Coupled model
 - Atmosphere (IFS-15r8) + Ocean (HOPE ~2x2)
- Ocean Analysis: OI
- Ensemble: “Lag average” (1 fc per day to 6 months)
- The Calibration Period: 1991-1996

January 2002: Seasonal Forecast System 2

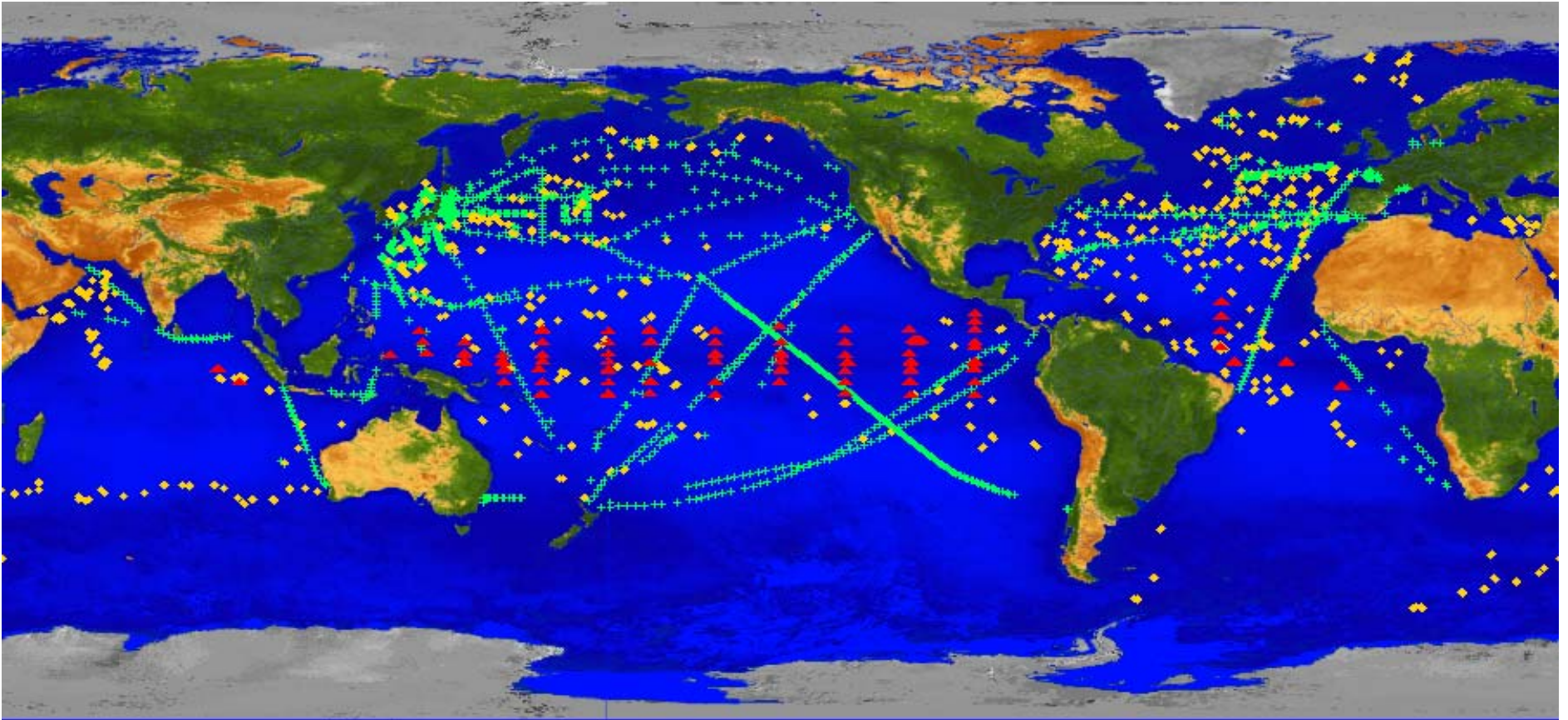
- The Coupled model
 - Atmosphere (IFS-23r4) + Ocean (HOPE 1x1)
- Ocean Analysis: (OI)
 - 5-member ensemble of ocean analysis (wind pertub.)
- Ensemble: “*BURST MODE*”
 - 40 forecasts made on 1st of month to 6 months
- The Calibration Period: From 1987-2001, using ERA15

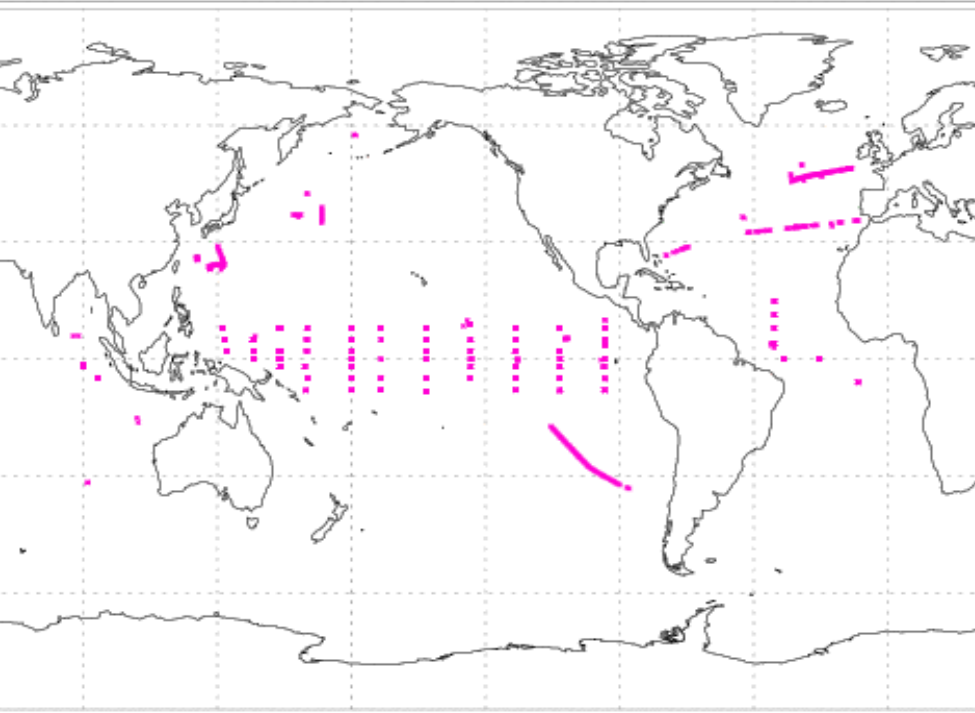
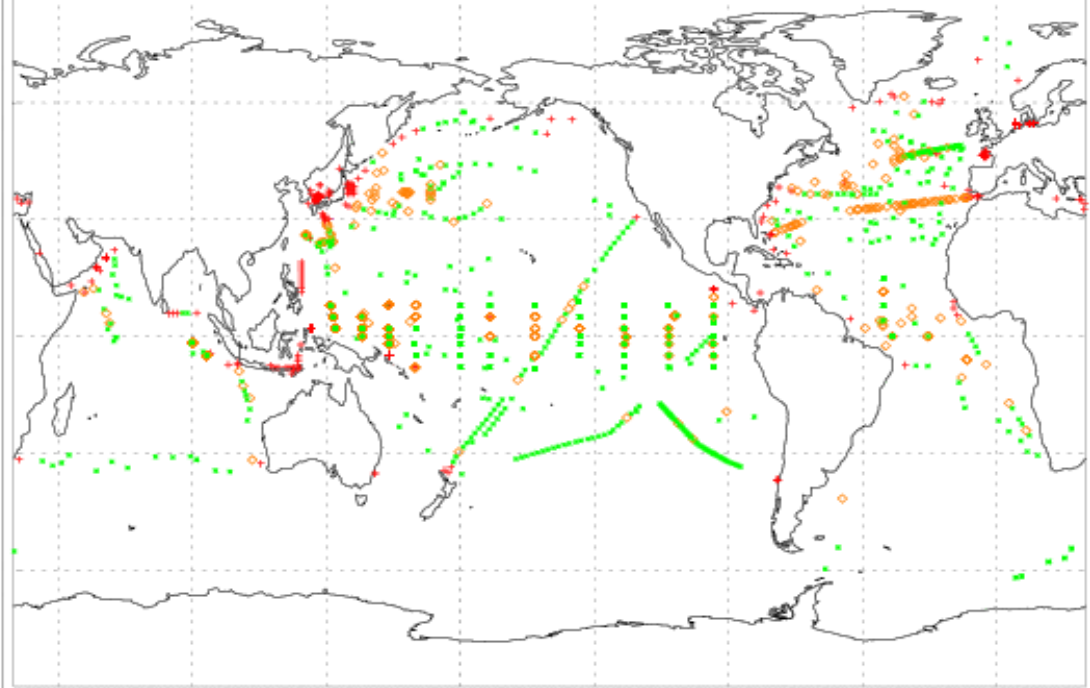
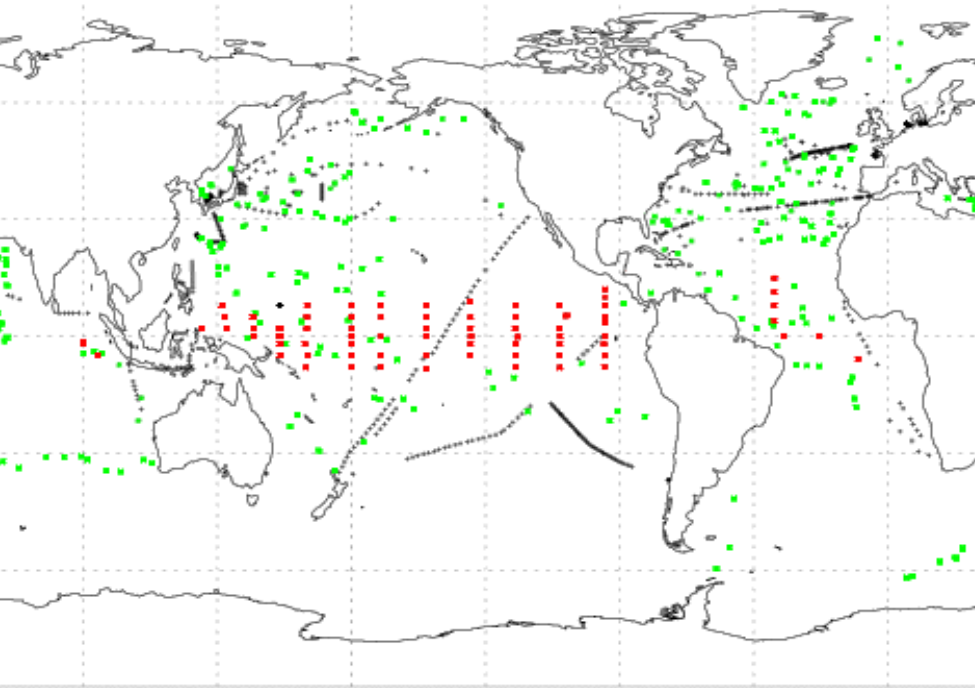
OCEAN INITIALIZATION

- **Strong Relaxation to Reynolds OIv2 SST (~2 days)**
- **OI of subsurface T, every 10 days**
- **10 days assimilation window**
- **Salinity Updates (T-S scheme) *New***
- **Velocity Updates (geostrophy) *New***
- **Subsurface 3D relaxation to T and S Levitus 98 (~18 months). *New***
- **Ensemble of 5 analysis *New***
- **Daily forcing for mass, momentum, and heat from NWP**
- **Wind perturbations (SOC-ERA, monthly values)**
- **11 days behind real time**

Data coverage for May 2002

- XBT, MOORINGS, ARGO floats





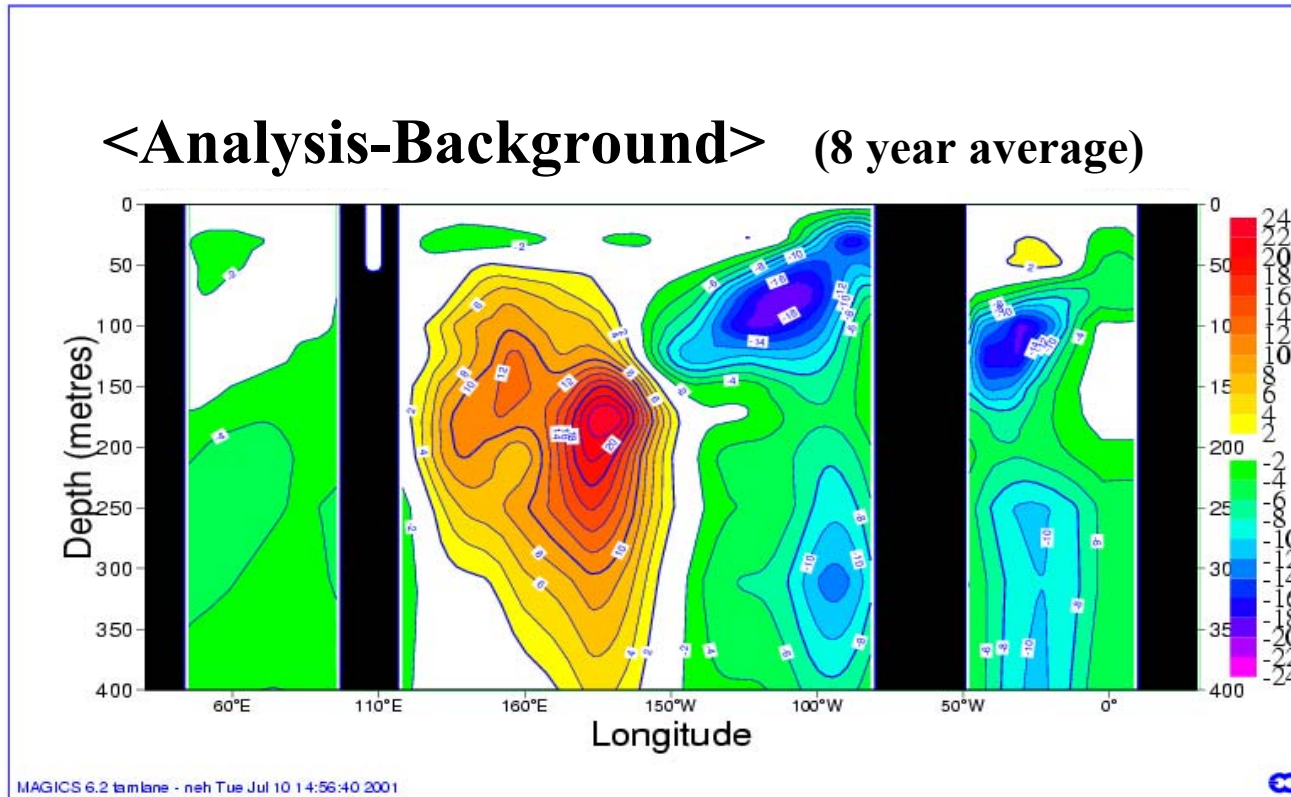
x TAO+TRITON+PIRATA (2255)
 x ARGO floats (1614)
 x other (65452)

x rejected profile (575)
 x full profile accepted (1778)
 o partially accepted profile (948)

x Superobs (2325)
 (at least one of the profile)

Position of in-situ observations that entered the ECMWF operational system from date 20020528 to 20020528

Assimilation Increments

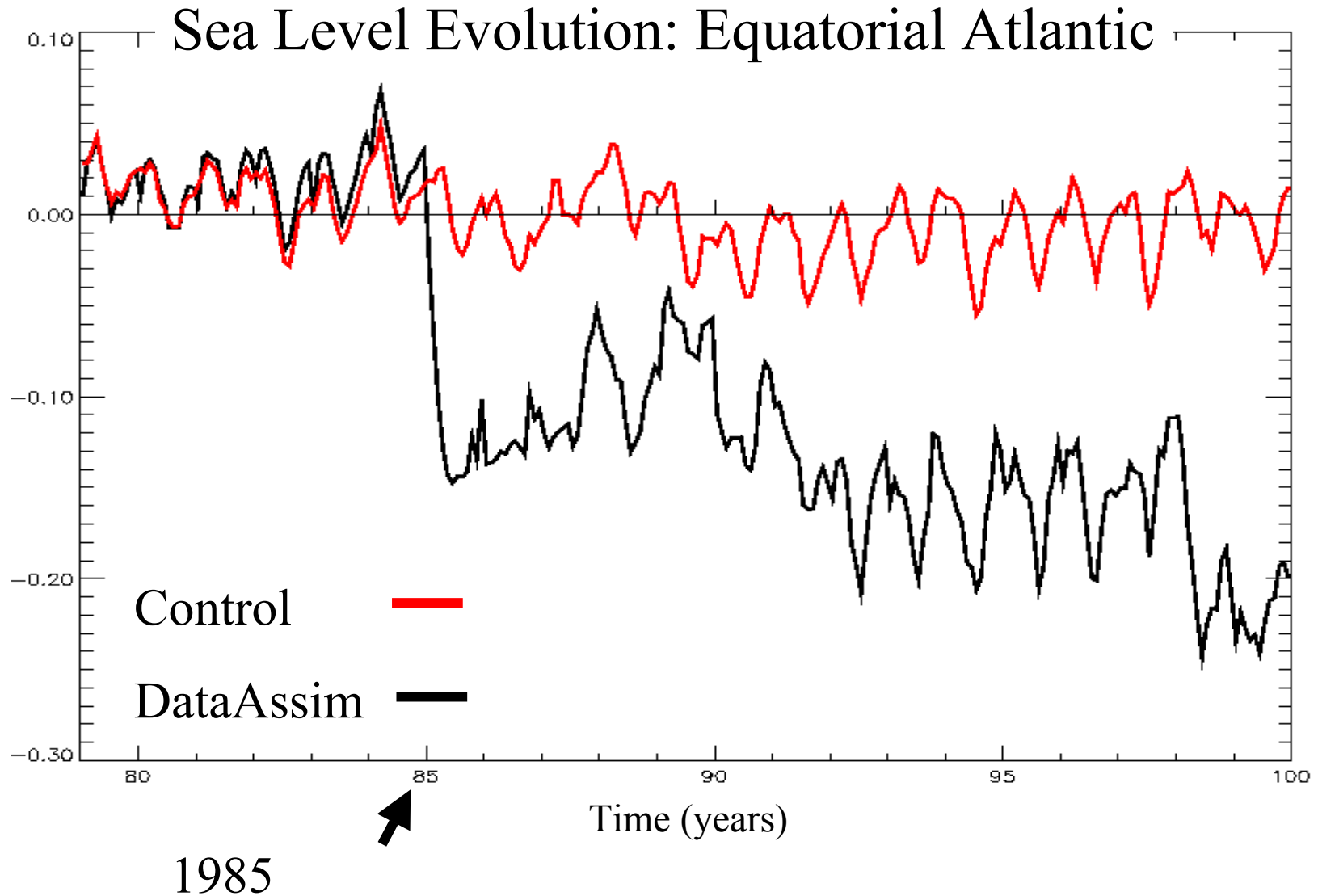


C.I: 2 C/year

Why worry about systematic error?

- Suboptimal use of observations:
 - Increased variance in the analysis, convergence, etc...
- Practical considerations:
 - spurious time variability: If observations are not homogeneous in time
 - Systematic error in “unconstrained” variables

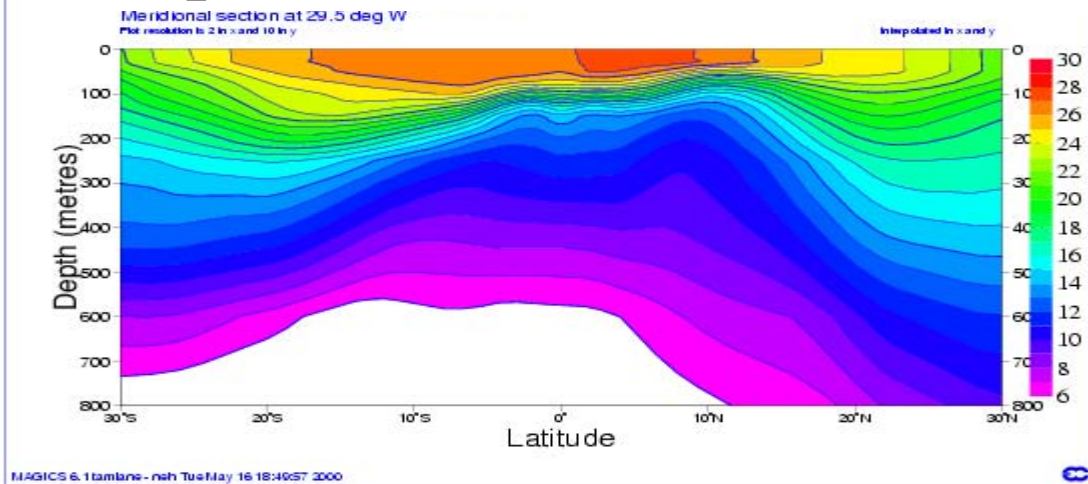
SYSTEM 1



OCEAN INITIALIZATION

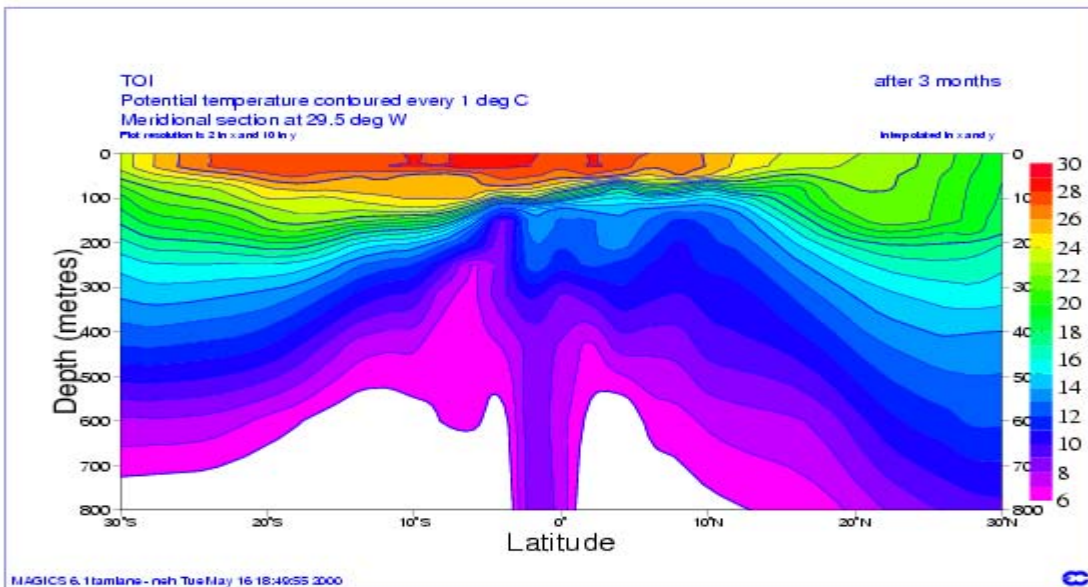
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Temperature Section 30 W



Initial Conditions :

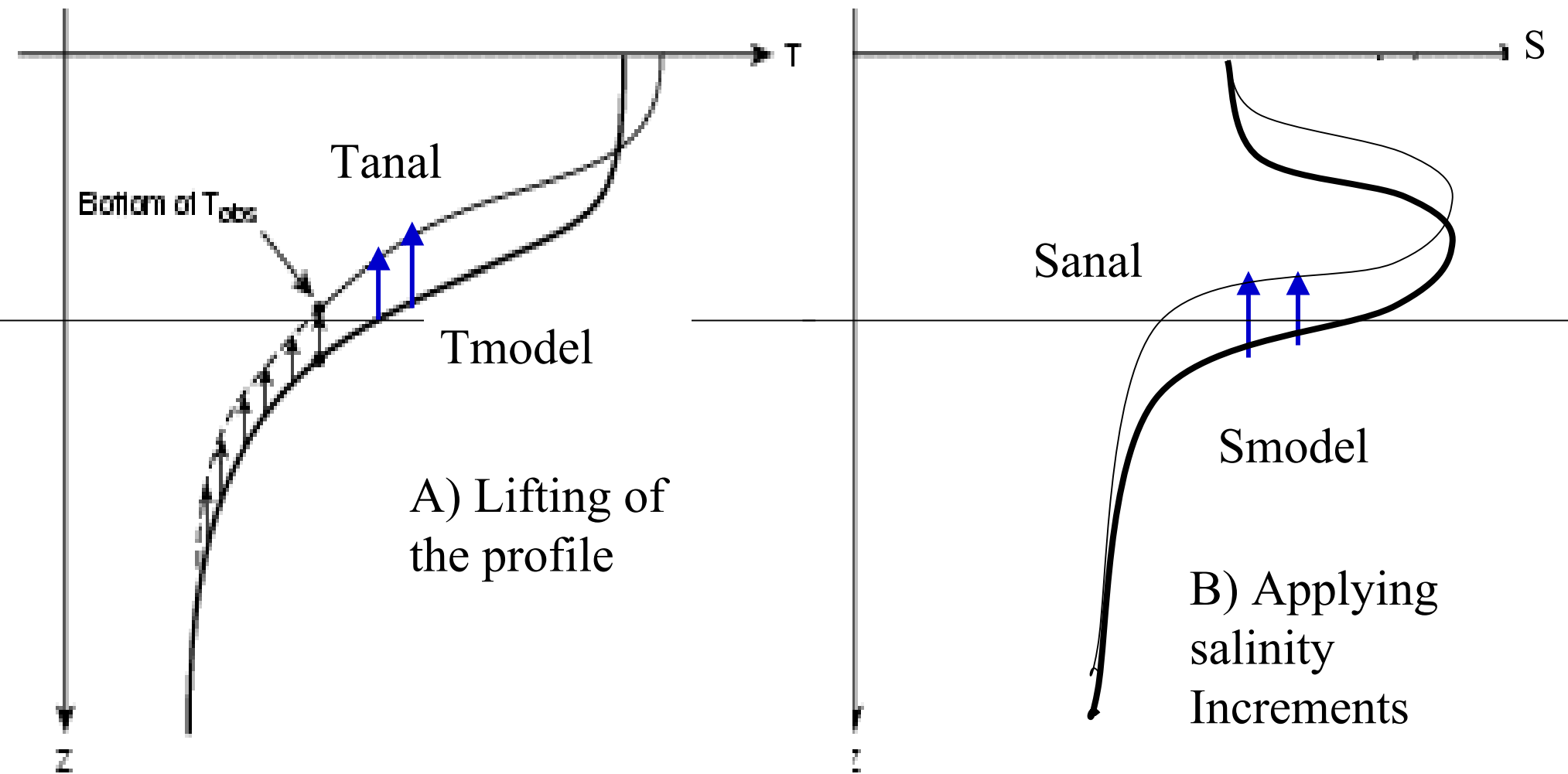
Stratified water column



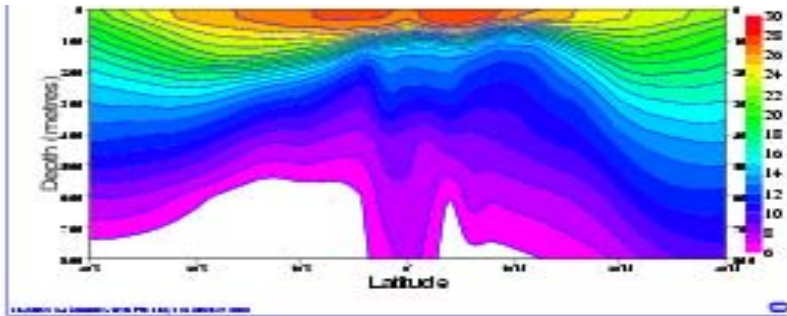
**3 months into the
assimilation of Subsurface
temperature:**

Convection occurs at the
Equator. Broken stratification

Updating Salinity: S(T) SCHEME



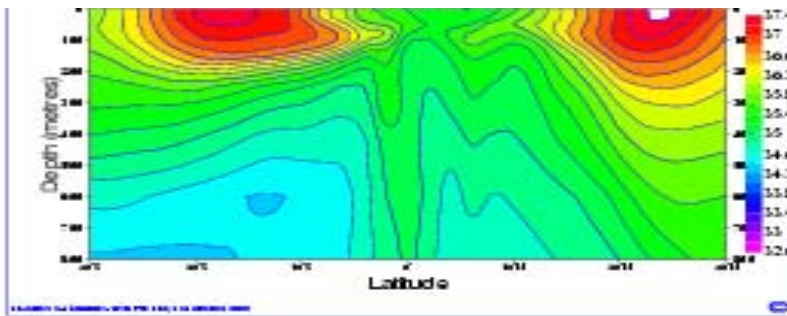
Temperature : No S(T)



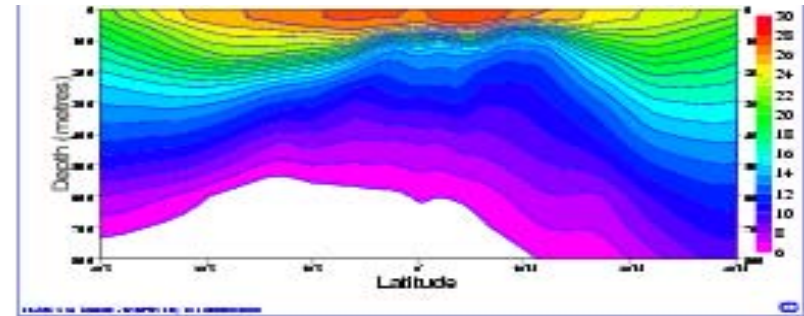
Spurious Convection



Salinity: No S(T)



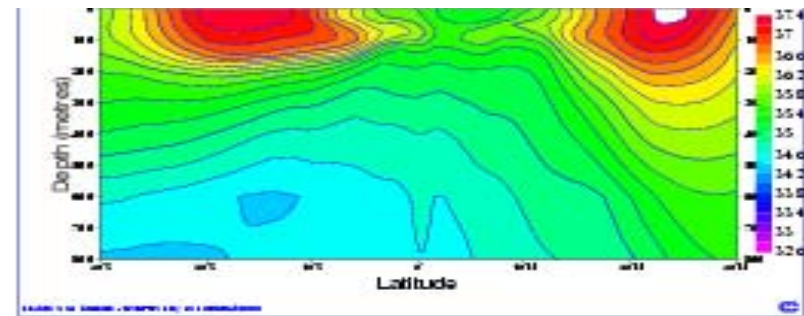
Temperature: S(T)



S(T) prevents convection



Salinity: S(T)

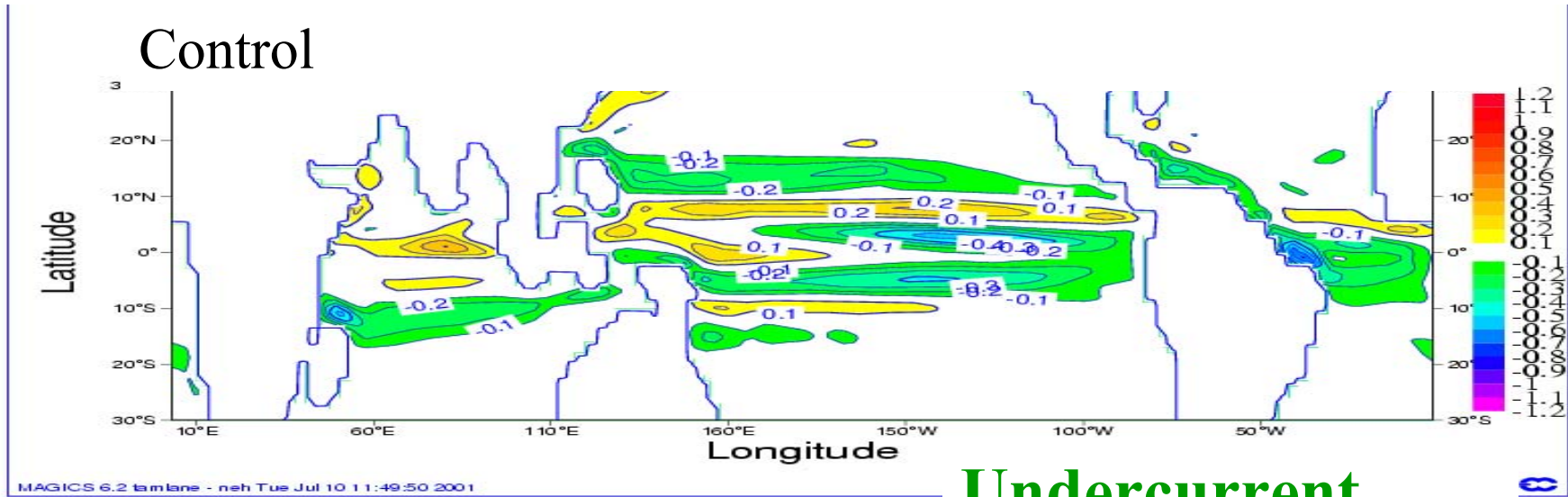


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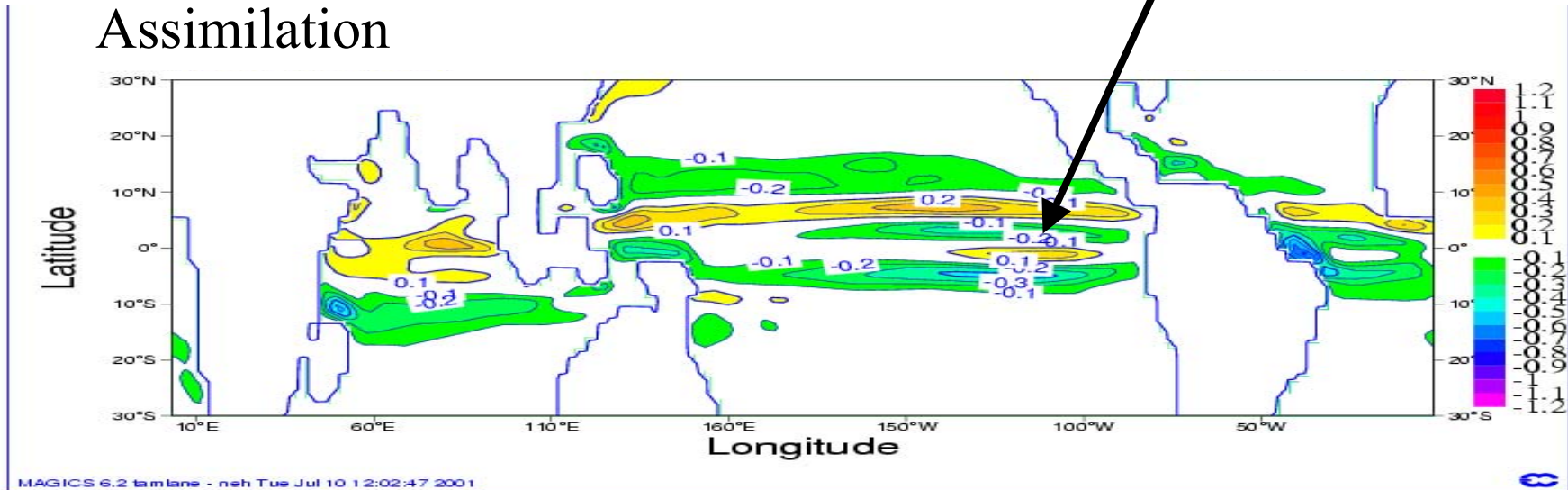
Surface Zonal Velocity

Control



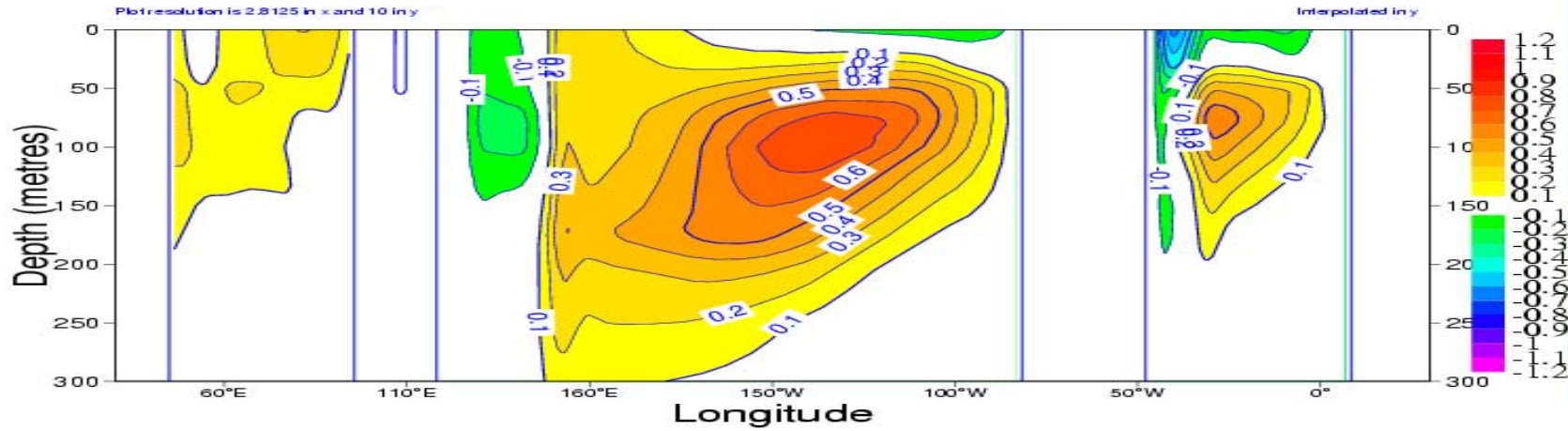
Undercurrent
Surfaces

Assimilation



Equatorial Zonal Velocity

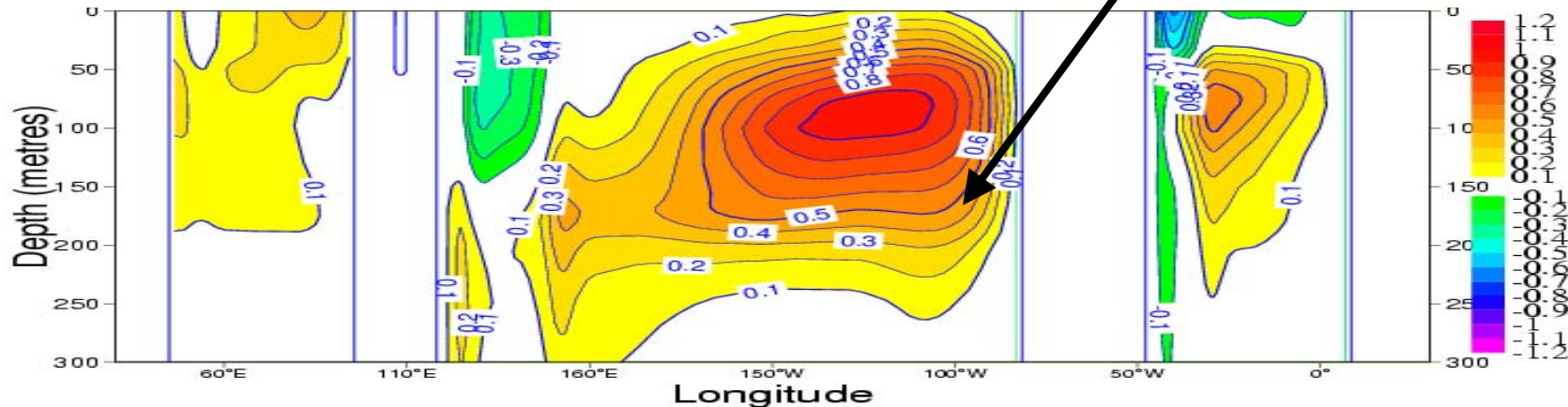
Control Run



MAGICS 6.2 tmlane - neh Tue Jul 10 11:49:50 2001

Undercurrent too strong and deep

Assimilation Run



MAGICS 6.2 tmlane - neh Tue Jul 10 12:02:50 2001



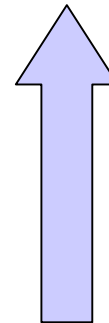
Problem with Equatorial Currents

**Equatorial currents
degrade when assimilating
density information
(temperature /salinity)**

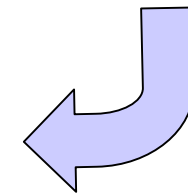
**The currents get worse
as the constraint to
the density data is
increased**

- Possibility of feedback :

error in density



density increment



degradation of
currents

- **Strategy 1: to introduce dynamical constraints** (*Burgers et al, JPO 2002*)
 - Error Induced by the assimilation process by disruption of dynamical balances

- **Strategy 2: to estimate and correct error** (*Bell, Martin, Nichols, 2001*)
 - It assumes that the source of error lies on the momentum equation

Balanced Currents Method

- To update currents / the velocity increment is partially in geostrophic balance with the density increments:

$$\eta_a = \eta_b + \delta\eta ; \vec{u}_a = \vec{u}_b + \delta\vec{u}$$

$$\delta\eta = \delta\eta_n + \delta\eta_e ; \delta\eta_e = \alpha\delta\eta ; 0 \leq \alpha \leq 1$$

$$\delta u = -\frac{g}{f} \frac{\alpha \partial \delta\eta}{\partial y} ; \delta v = \frac{g}{f} \frac{\alpha \partial \delta\eta}{\partial x}$$

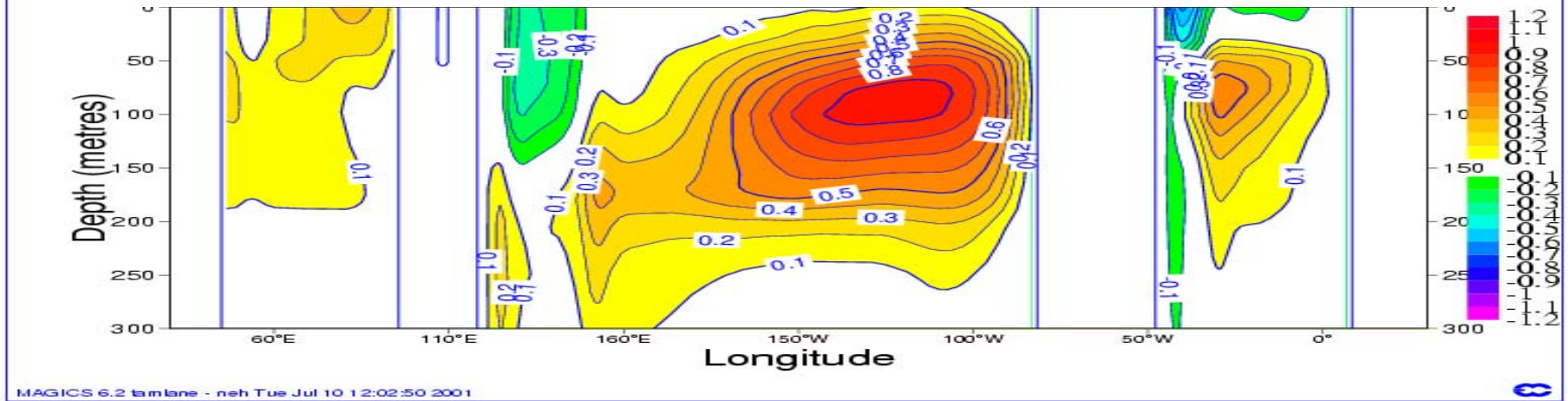
- At the Equator:

$$\delta u = -\frac{g}{\beta} \frac{\alpha \partial^2 \delta\eta}{\partial y^2} ; \delta v = 0$$

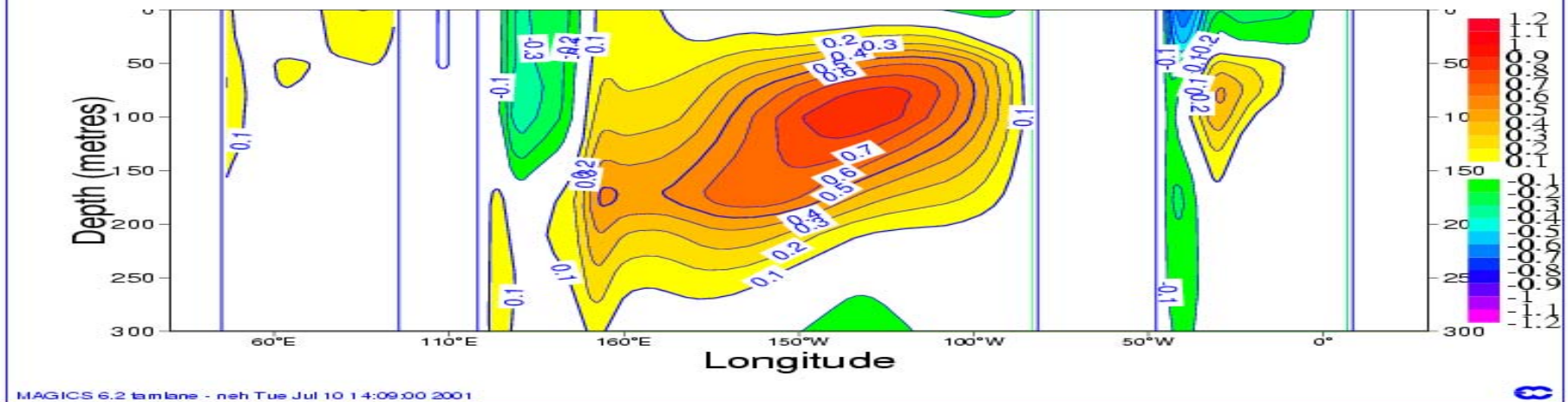
α may
depend
on z

Equatorial Section: Zonal Velocity

Assimilation

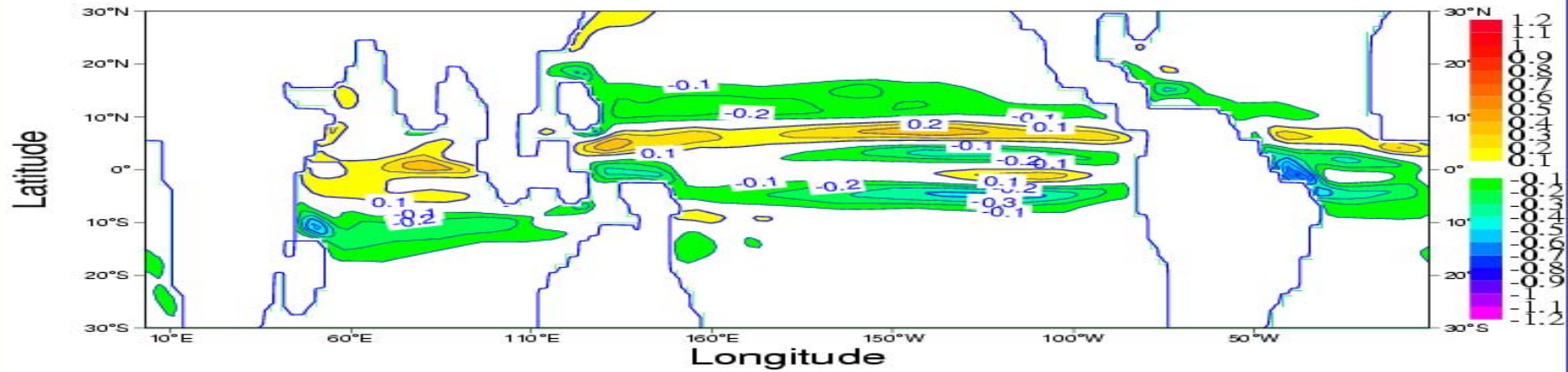


Assimilation + Geos_Balanced_currents

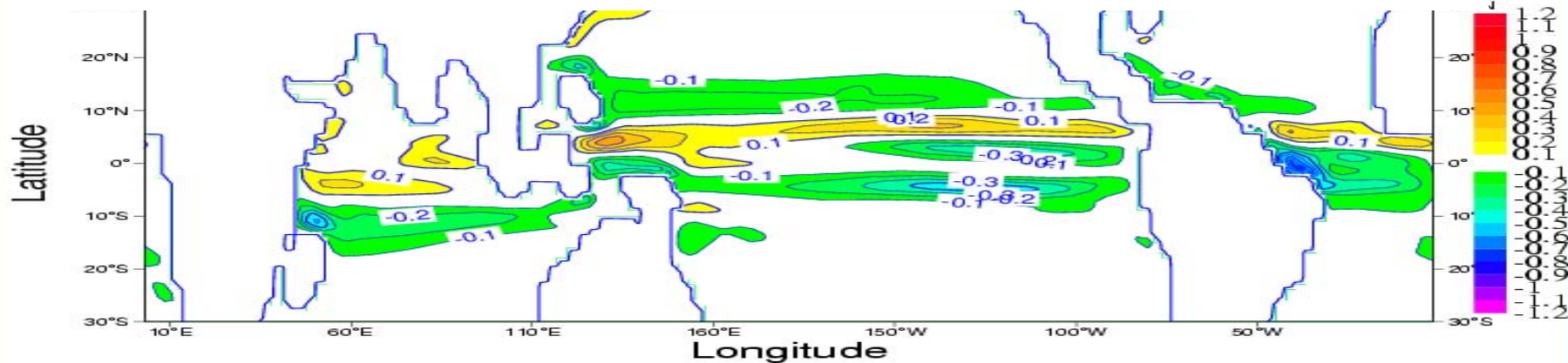


Surface Zonal Velocity

Assimilation

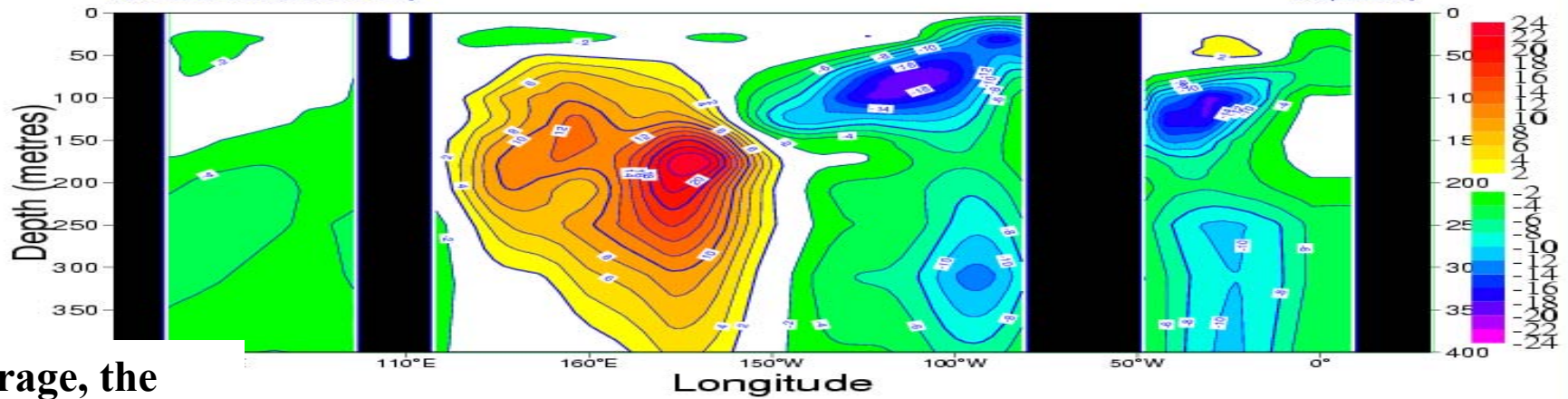


Assimilation + Geos_balanced



Temperature Increments

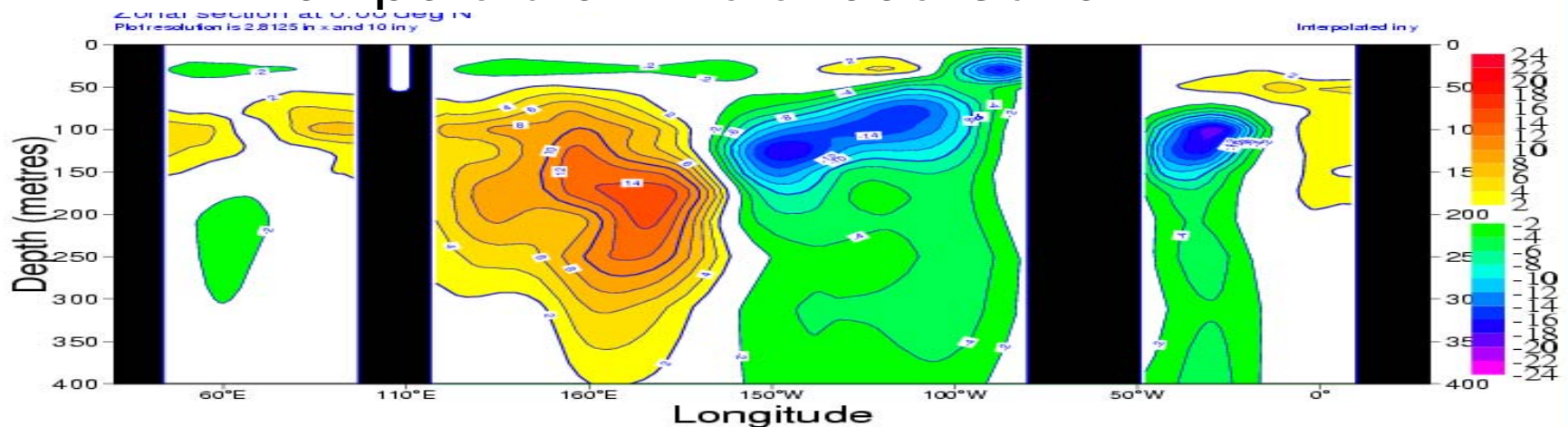
Temperature only



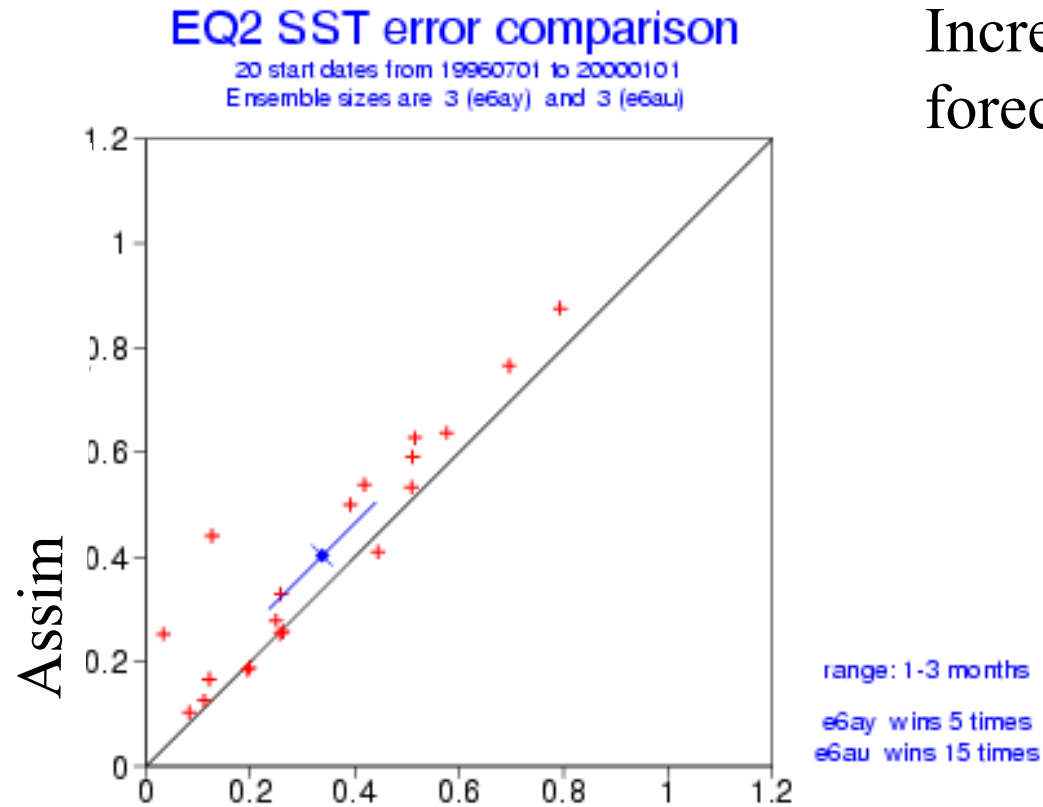
On average, the averaged assim increment

decreases by ~20%

Temperature + Balanced Current



Velocity
Increments help
forecast statistics



Dealing with Model Error

- More general solution
- Estimation problem:
 - Time dependency?
 - Which equations are in error?
 - Uncertainty in the bias estimate
- Sequential Methods (Dee and Da Silva, 1998, Bell et al, 2001 ...):
they use only past information
- Adjoint methods (Derber 1989, Bonekamp et al, 2001, Vidard et al, 2003): they use also future information

Bias Correction Methods

- Data assim system:

$$x(t_0) = x_b + \varepsilon(t_0)$$

$$\frac{dx}{dt} = \mathbf{M}(x(t)) + \eta(t)$$

$$\mathbf{H}(x(t_0)) = y_0 + \delta(t_0)$$

with

$$\langle \varepsilon(t) \rangle = 0; \dots \langle \varepsilon | \varepsilon \rangle = \mathbf{B}$$

$$\langle \delta(t) \rangle = 0; \dots \langle \delta | \delta \rangle = \mathbf{R}$$

$$\langle \eta(t) \rangle = \mathbf{b}; \dots \langle \eta | \eta \rangle = \mathbf{Q}$$

$$d = (\mathbf{H}x_b - y)$$

$$x_a = x_b + \mathbf{K}d$$

State Augmentation

$$x(t_0) = x_b + \varepsilon(t_0)$$

$$\frac{dx}{dt} = \mathbf{M}(x(t)) + \mathbf{T}(\eta(t))$$

$$\eta(t_0) = \eta_b$$

$$\frac{d\eta}{dt} = \mathbf{\Phi}(\eta(t)) + \mu(t)$$

$$x_a = x_b + \mathbf{K}_b d$$

$$\eta_a = \eta_b - \mathbf{K}_\eta d_\eta$$

T : Which model variables are biased?

M : how does the model error evolve in time?

K_η : what is the gain matrix of the bias term?

Bell et al, 2001: systematic error in momentum equations

M: how does the model error evolve in time?

\mathbf{K}_η : what is the gain matrix of the bias term?

T : Which model variables are biased?

$$\Phi = \mathbf{I} ; \text{constant}$$

$$\eta_b(t+1) = \eta_b(t)$$

$$\mathbf{K}_\eta = \alpha \mathbf{K}_b$$

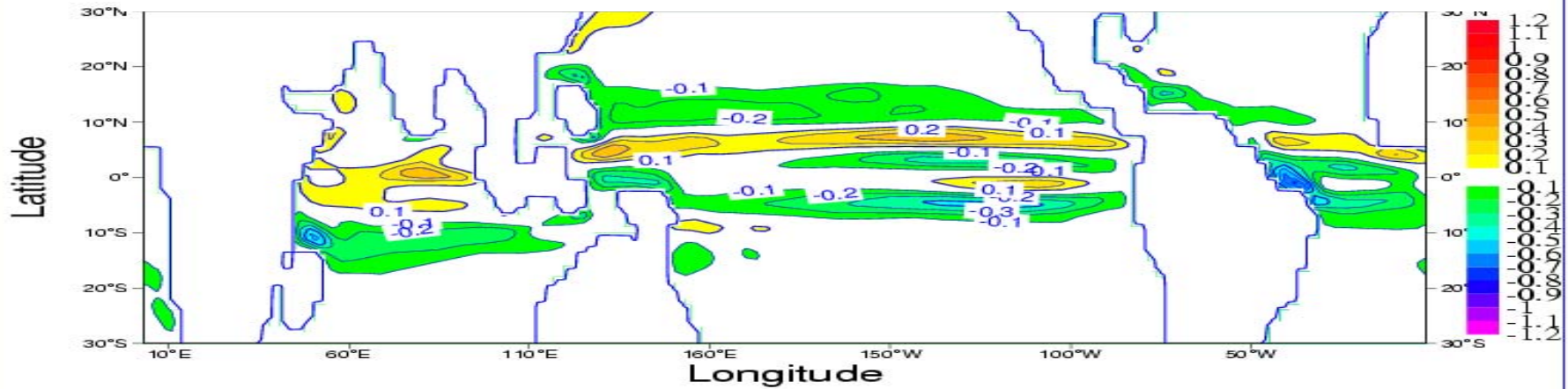
$$\eta_a = \eta_b - \alpha \mathbf{K}_\eta d$$

$$T = (X, Y)$$

$$X = -g \frac{\partial(\eta)}{\partial x} ; Y = -g \frac{\partial(\eta)}{\partial y}$$

Surface Zonal velocity

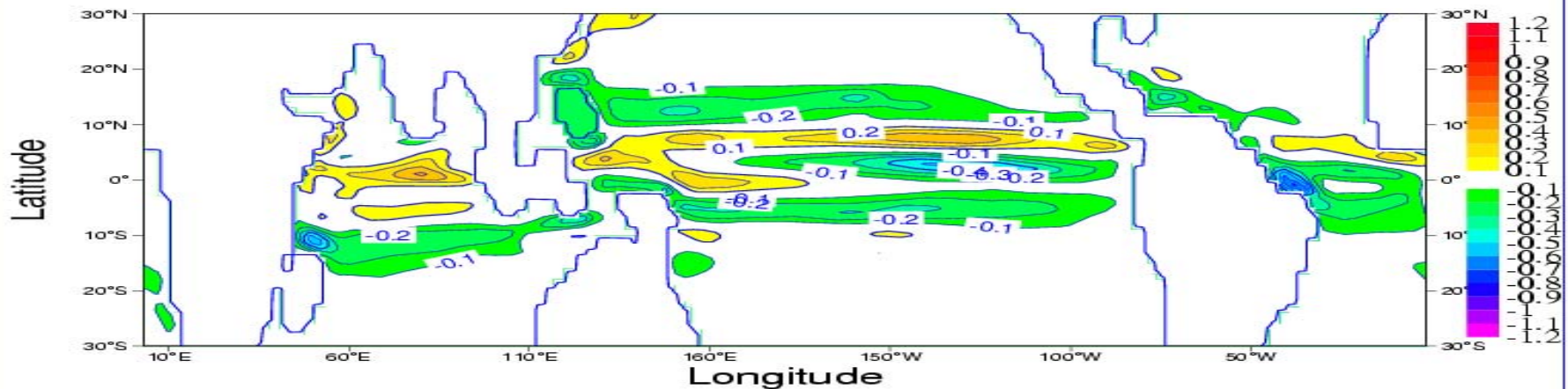
Assimilation



MAGICS 6.2 tam lane - neh Tue Jul 10 12:02:47 2001



Assimilation + Bias correction

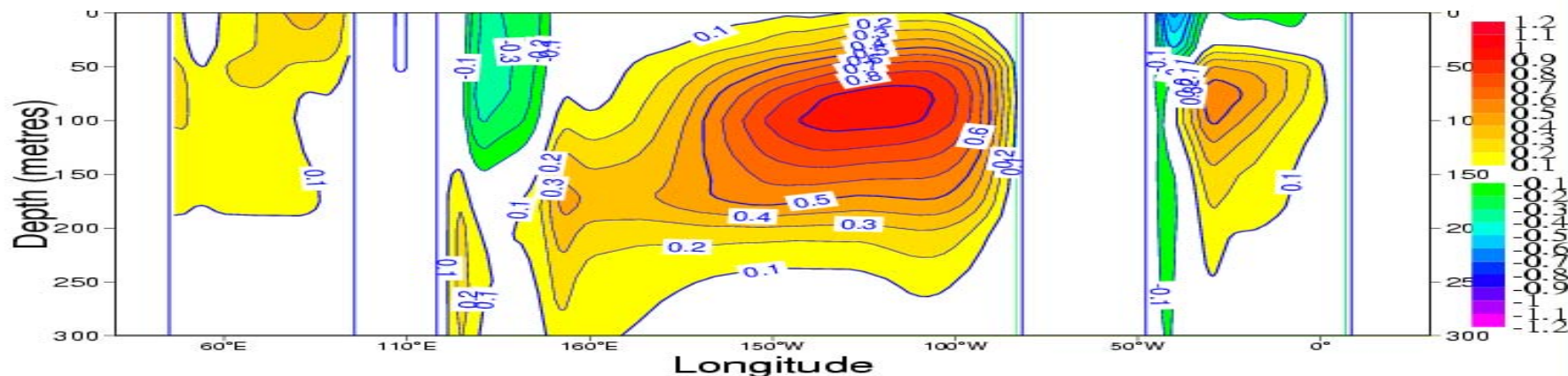


MAGICS 6.2 tam lane - neh Tue Jul 10 13:47:42 2001

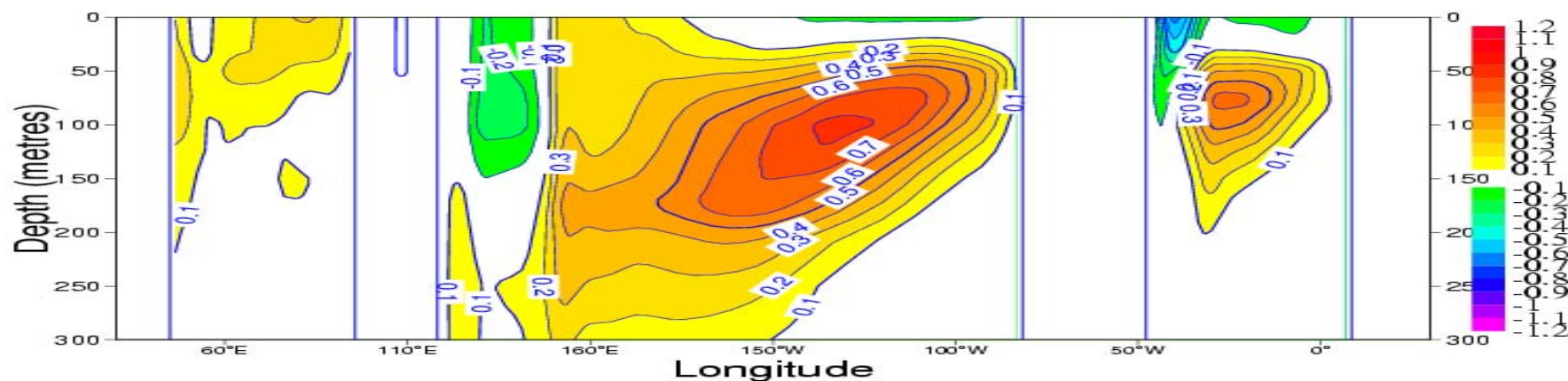


Equatorial Section of Zonal velocity

Assimilation



Assimilation + Bias correction

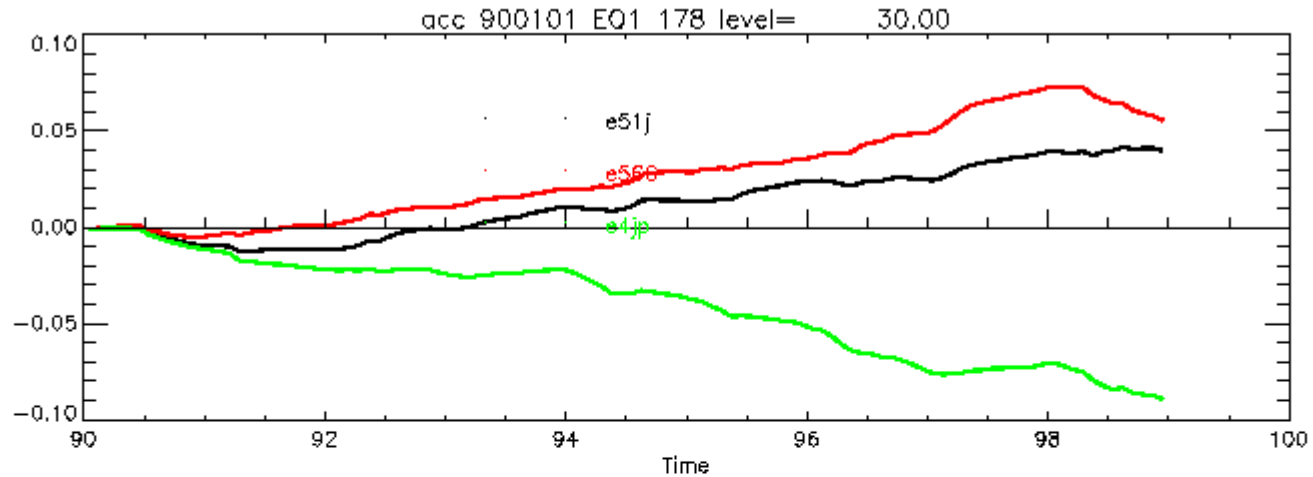


Accumulated Assimilation Increment

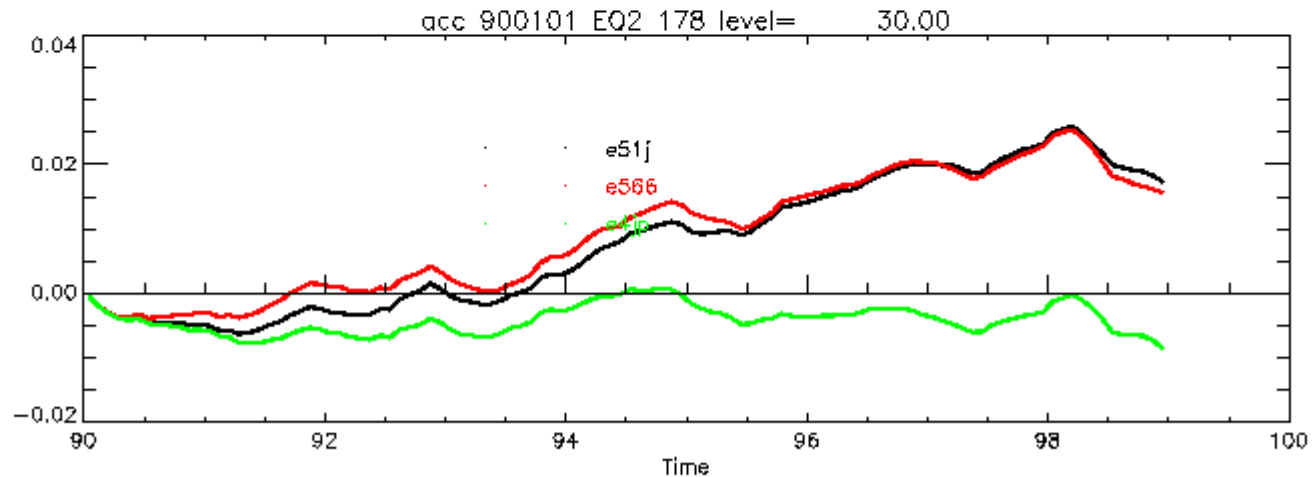
Standard

Bias $\beta = 0.1$

Bias $\beta = 0.3$



Eastern
Pacific
Bias
correction
works

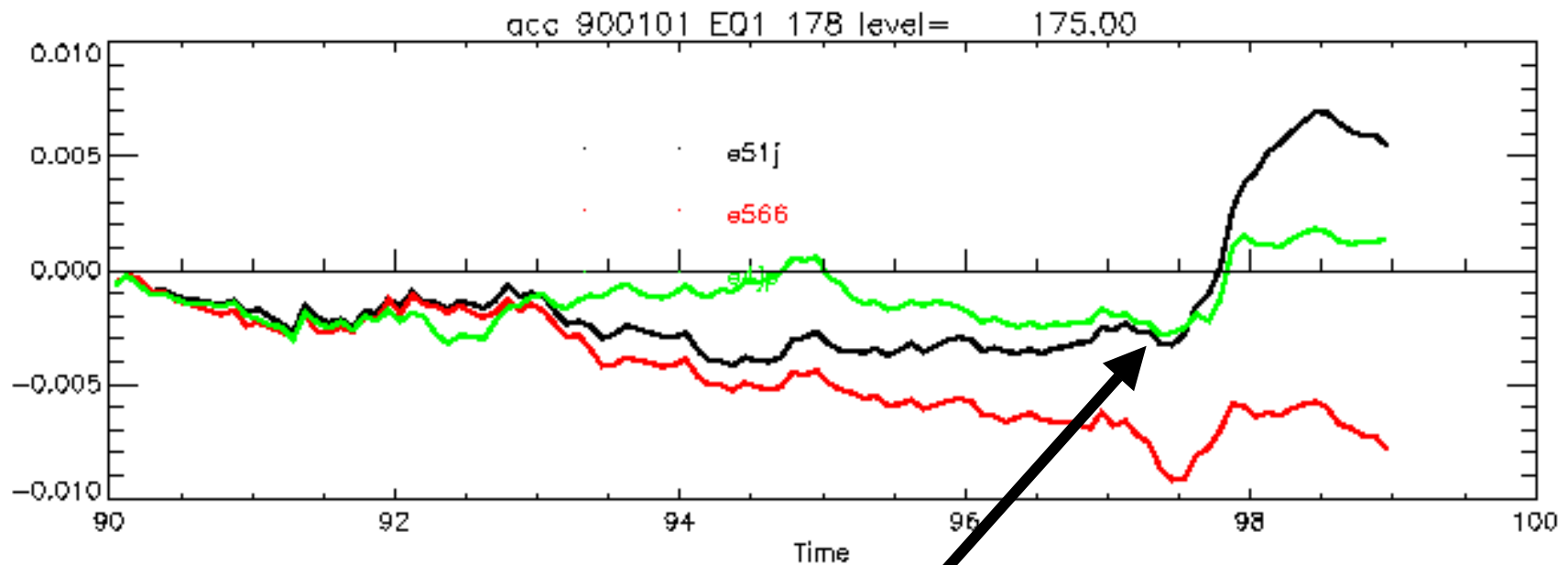


Central
Pacific

Standard

Bias $\beta = 0.1$

Bias $\beta = 0.3$



Bias is not constant

There is Interannual variability

Bias correction method

- It prevents the degradation of the velocity
By maintaining the balance between wind and pressure gradient
- Results are sensitive to the prescribe time evolution β
Errors may be flow dependent.
Introduce memory term? Or different slow evolving error terms (Dee lecture)
- Useful methodology for error diagnosis (passive)
- Online error correction requires robust estimates of model error
 - Adjoint methods may do better than sequential methods

In S2, it was decided to “control” the bias by a weak relaxation to Levitus climatology (T and S). It is not a definitive solution, and more work is needed on bias detection/correction.

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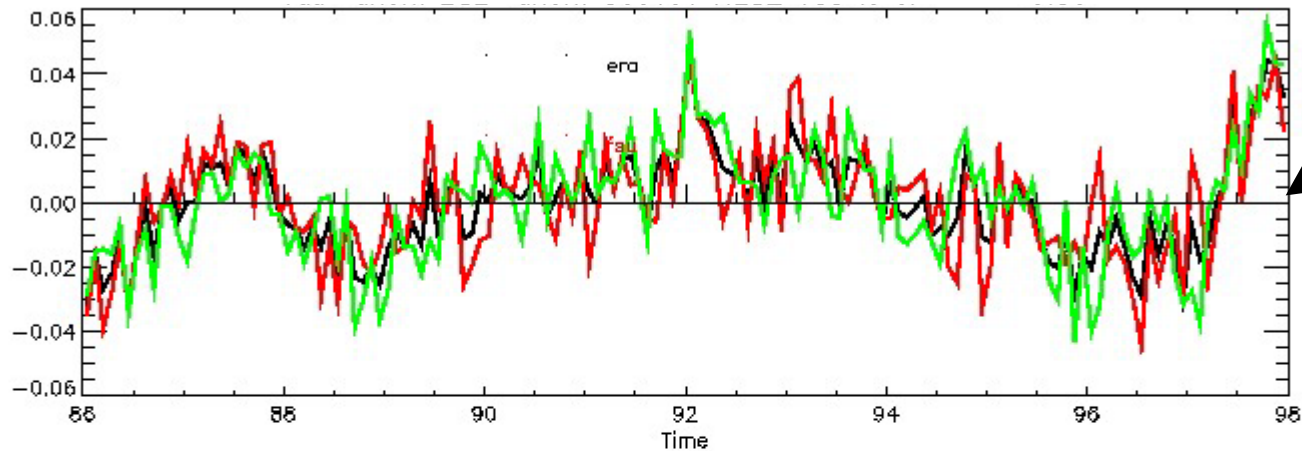
Ensemble of Ocean Analysis

The ensemble of forecasts is created by sampling the uncertainty on the Ocean Initial Conditions + uncertainty in sub-grid processes

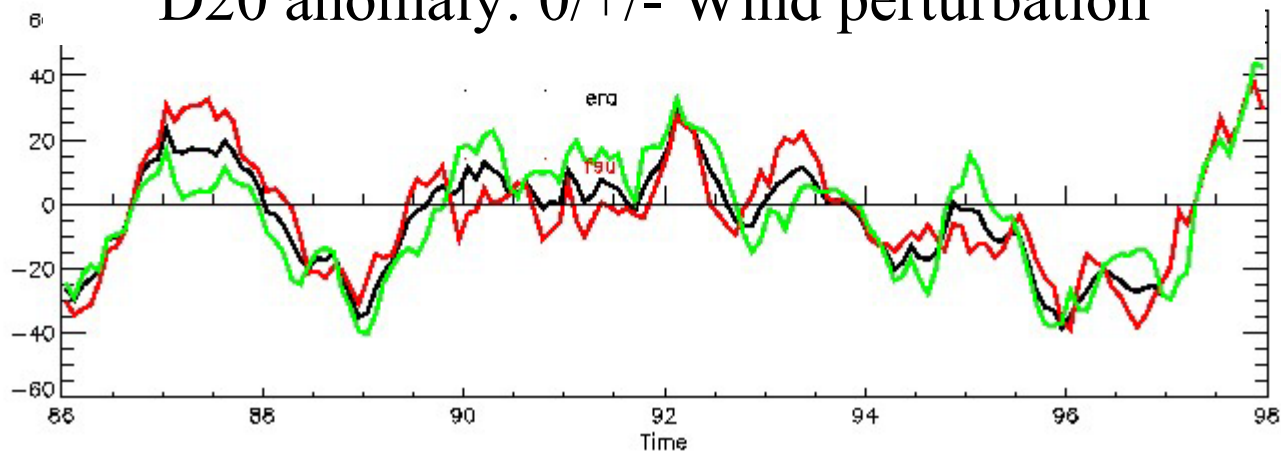
Ocean Initial Conditions: Surface + Subsurface

- The ensemble of ocean analysis is created to represent uncertainties in the subsurface, by adding perturbations to the wind stress forcing:
 - SOC - ECMWF monthly means (1980-2000)
 - 1 month de-correlation scale
- Applied during ocean analysis
 - They affect the subsurface structure of the ocean
- Not actively used in the estimation of error covariances

Taux anom: 0/+/- Wind perturbation

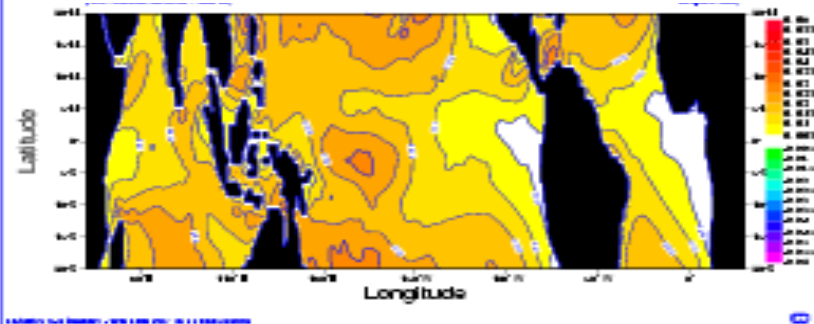


D20 anomaly: 0/+/- Wind perturbation

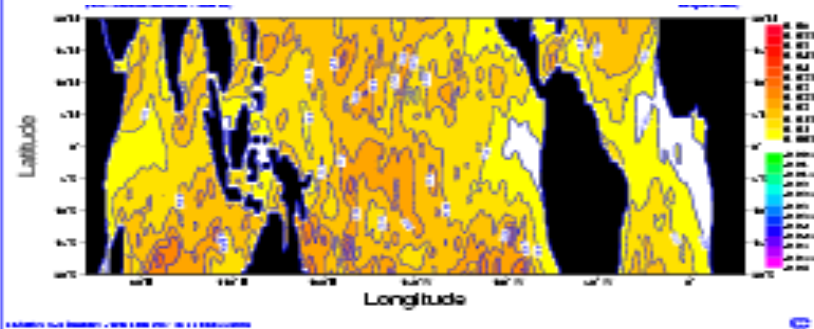


Ensemble spread
~ Interannual
variability

Taux Interannual C.I.=0.005Pa

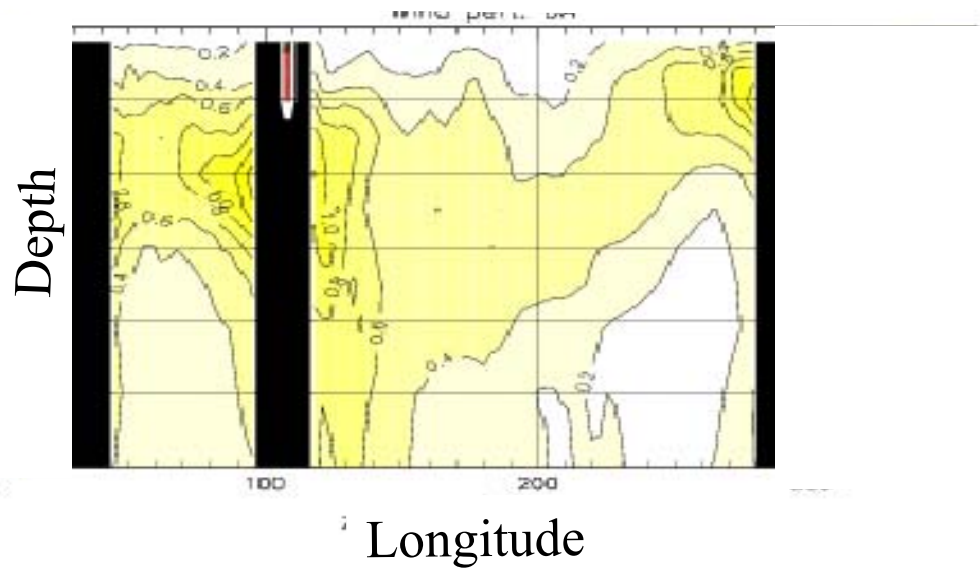


Taux Ensemble C.I.=0.005Pa

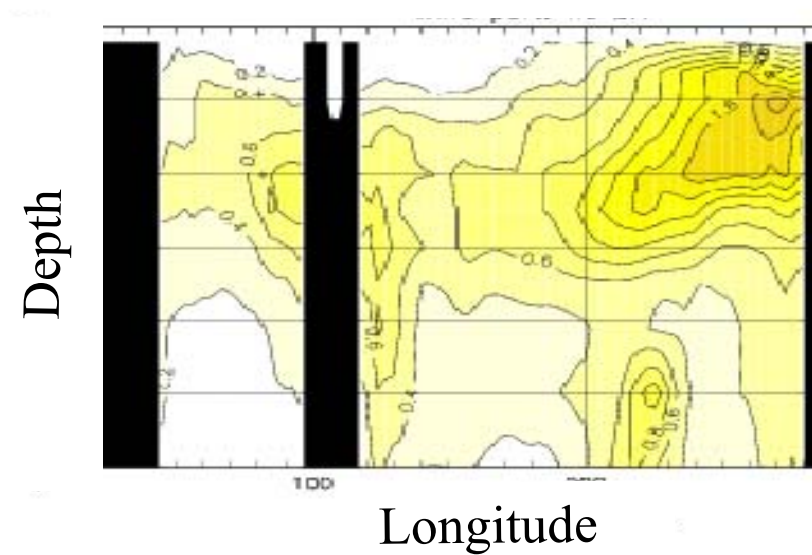


Uncertainties in ocean initial conditions due to wind stress errors

With
data assimilation



Without
data assimilation



Quality of the Analysis

A) Data assimilation

versus

No Data assimilation

B) New System (S2)

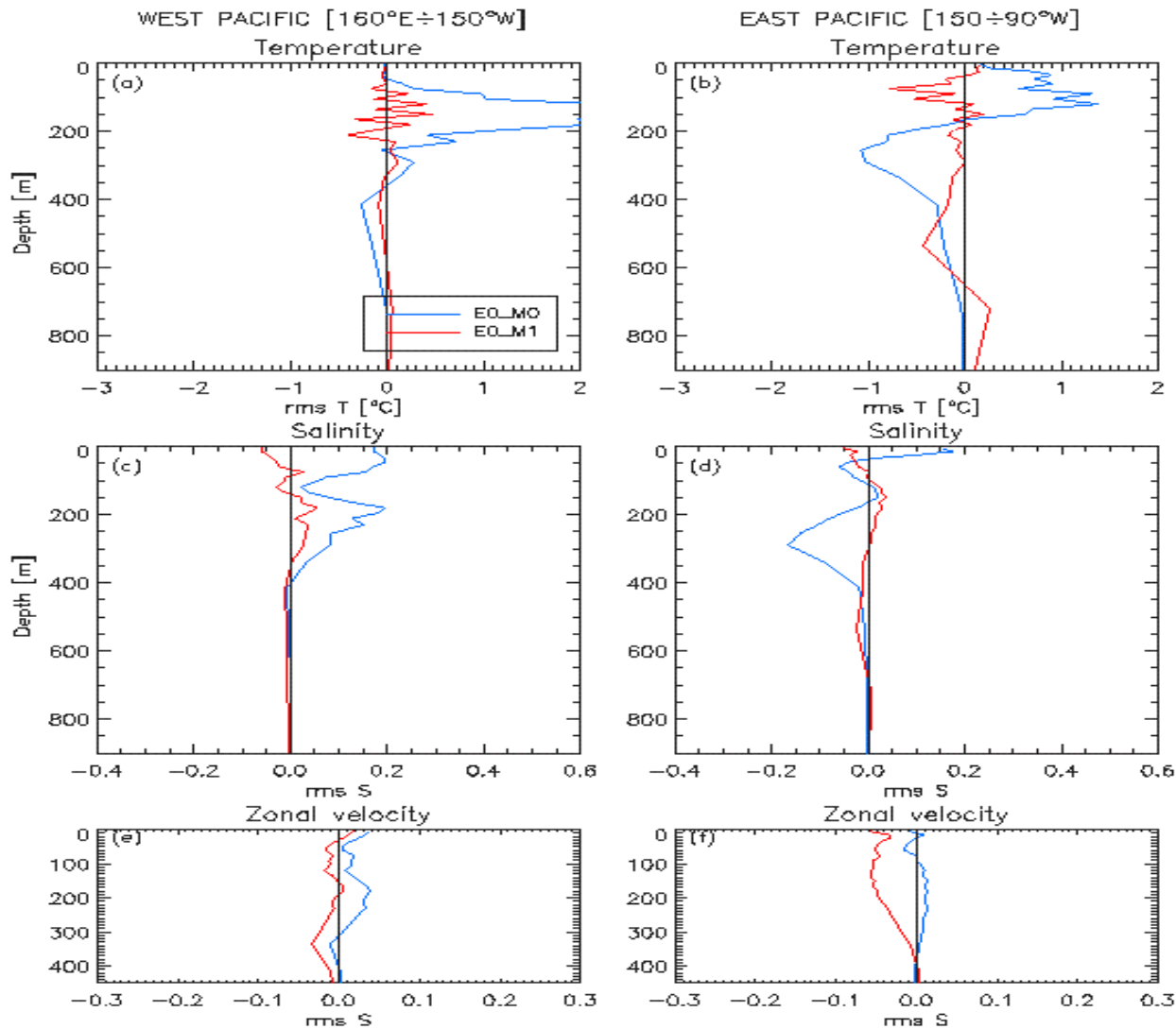
versus

Old System (S1)

State Estimation

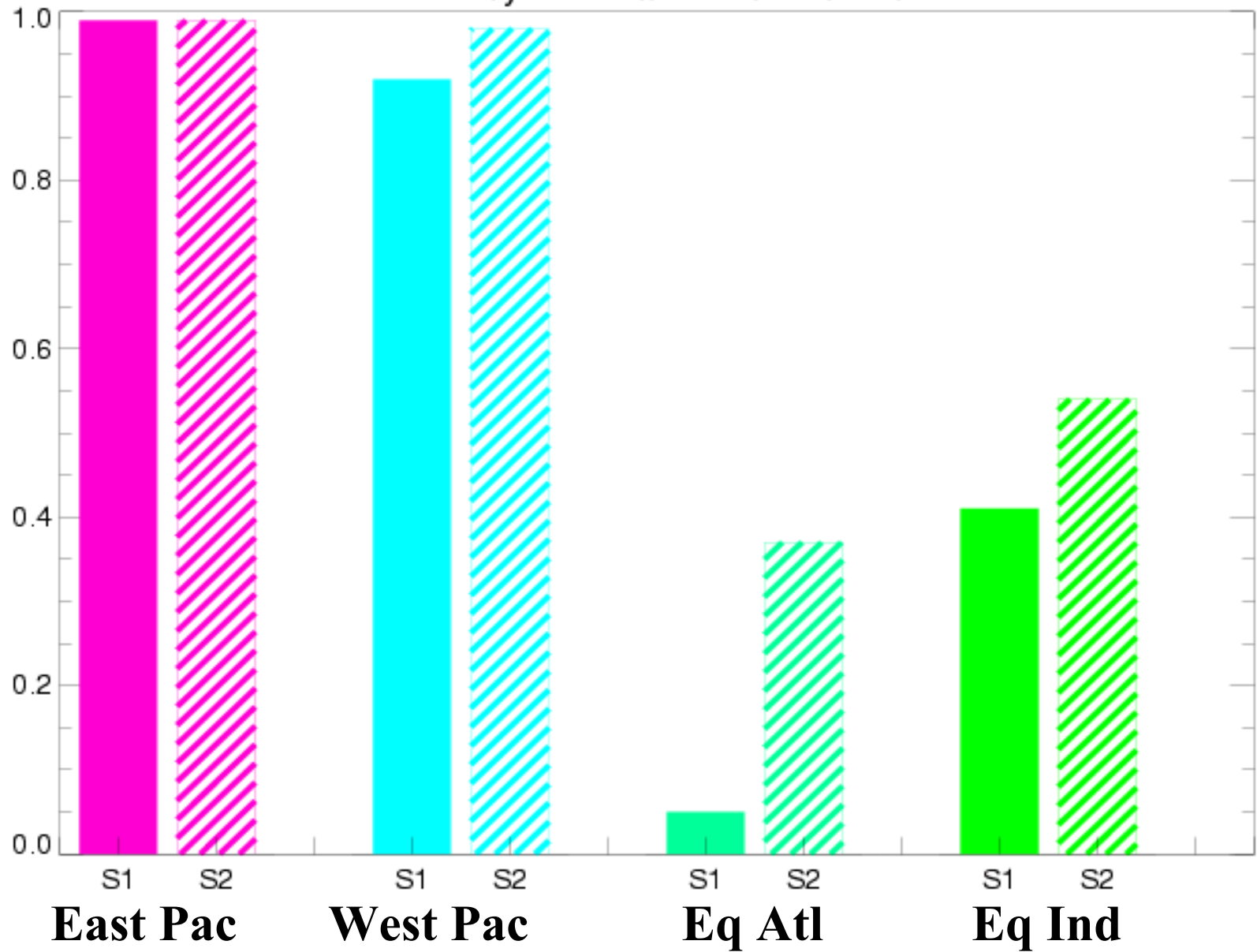
Comparison with Observations: mean Obs - Anal

Johnson et al, JGR (2000), Johnson et al, Progr. Ocean. (2002),



Produced
by Alberto
Troccoli

Anomaly correlation with Altimeter



Quality of the Analysis

A) Data assimilation

versus

No Data assimilation

**As Initial conditions for
seasonal forecasts**

B) New System (S2)

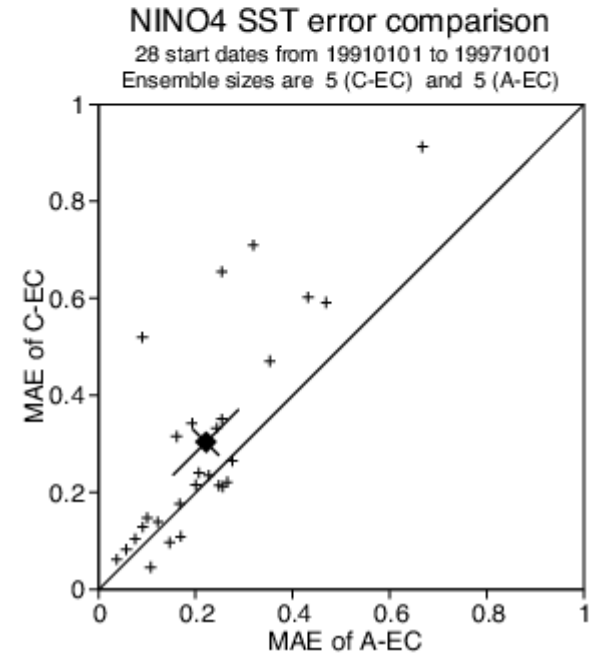
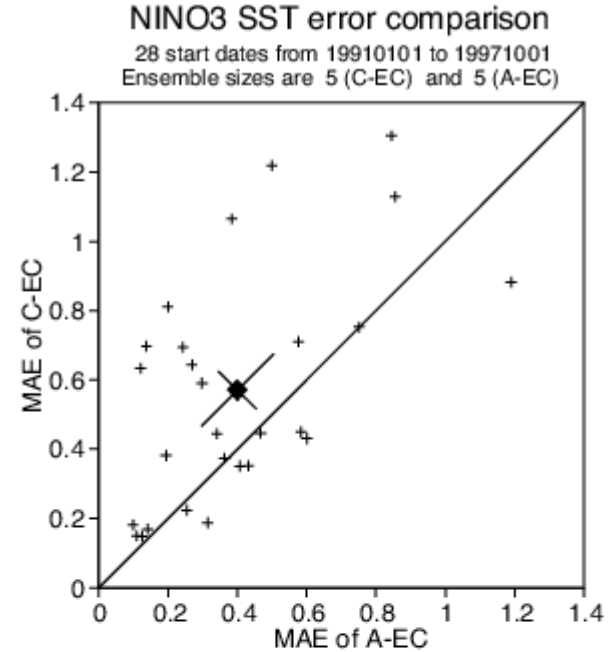
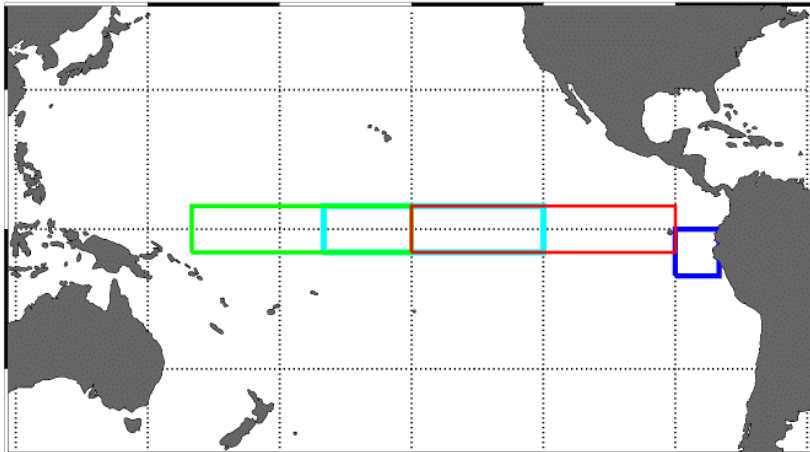
versus

Old System (S1)

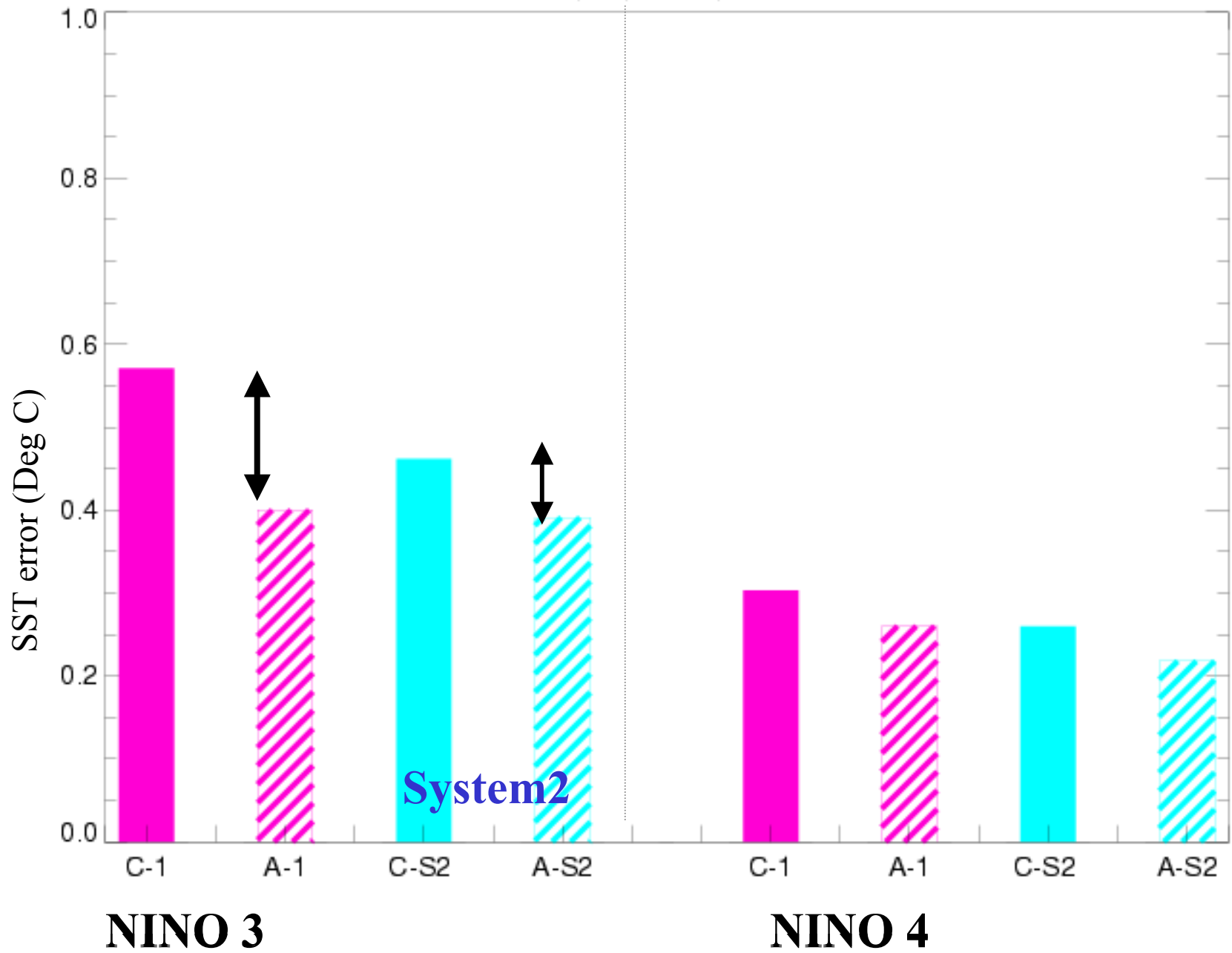
**Previously, it was shown
that ocean data
assimilation improves the
forecast skill**

(Alves et al 2003)

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 Nino12, Lon = [-90, -80], Lat = [-10, 0]
 Nino4, Lon = [160, -150], Lat = [-5, 5]
 Nino3, Lon = [-150, -90], Lat = [-5, 5]



Mean absolute error



NINO 3

NINO 4



Some results

State Estimation

- Data assimilation corrects the mean state of the ocean model
 - Temperature
 - Currents?
 - Salinity
- S2 ocean analysis shows better interannual variability than S1

Initial Conditions

- S2 forecasts have smaller errors than previous system
- Data assimilation improves the forecasts skill
- The impact of data assimilation is reduced
 - Need for better DA methods?
 - Need for better coupled models?

Developments

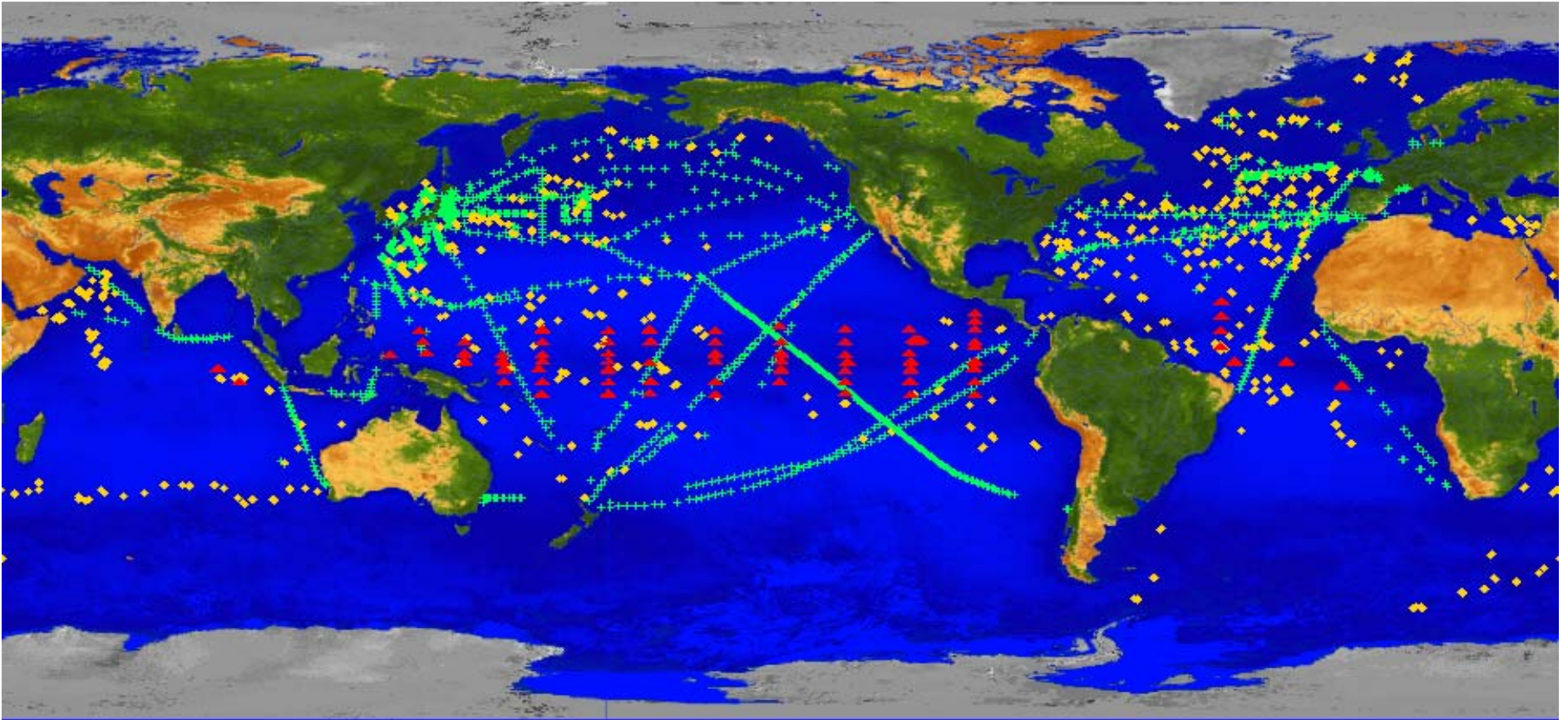
HOPE + OI

- **Different methods for combining SL with InSitu Data**
- **OSE experiments**
- **Assimilation of Salinity data**
- **Dealing with model error**

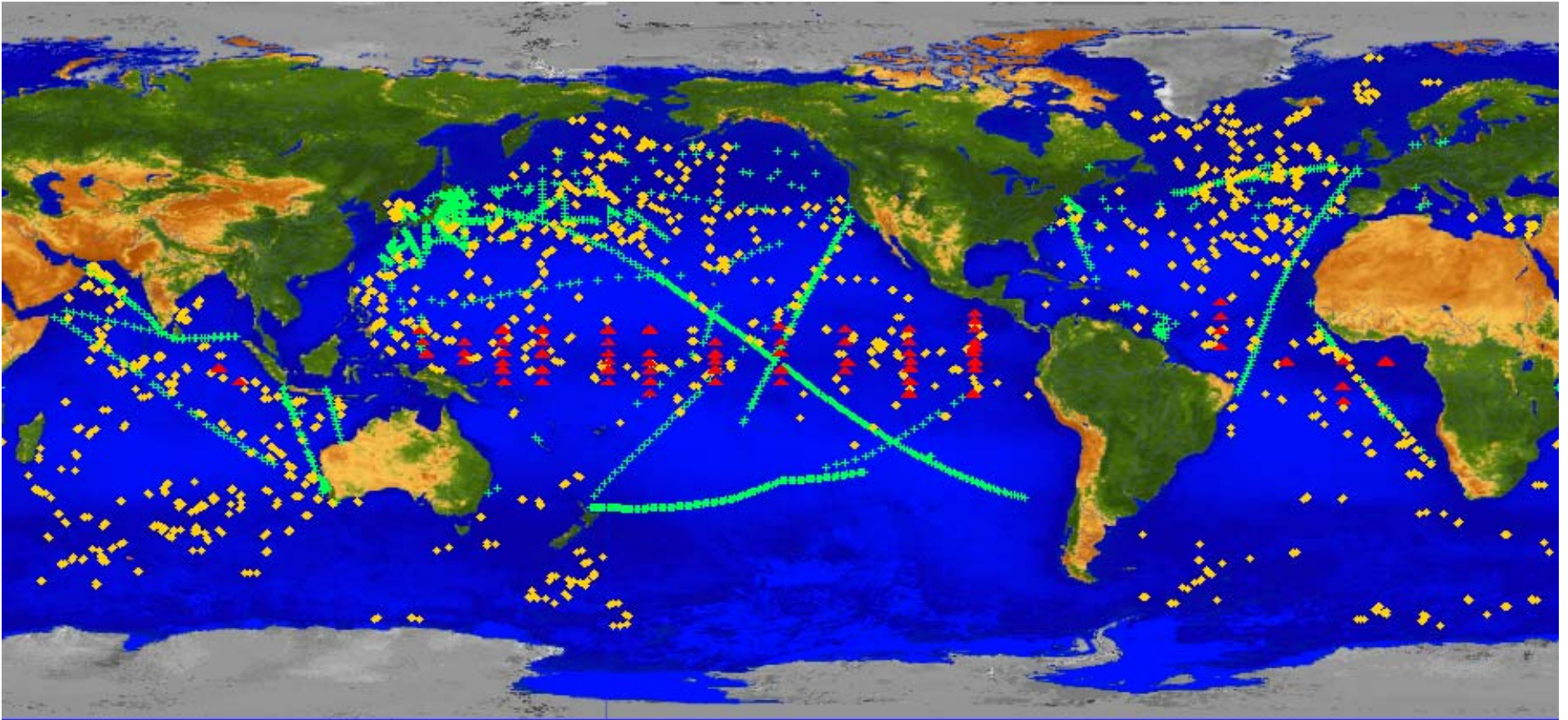
OPA-VAR

- **Installing the ocean model (ongoing)**
- **Installing (& integrating) the variational system (Weaver lecture)**

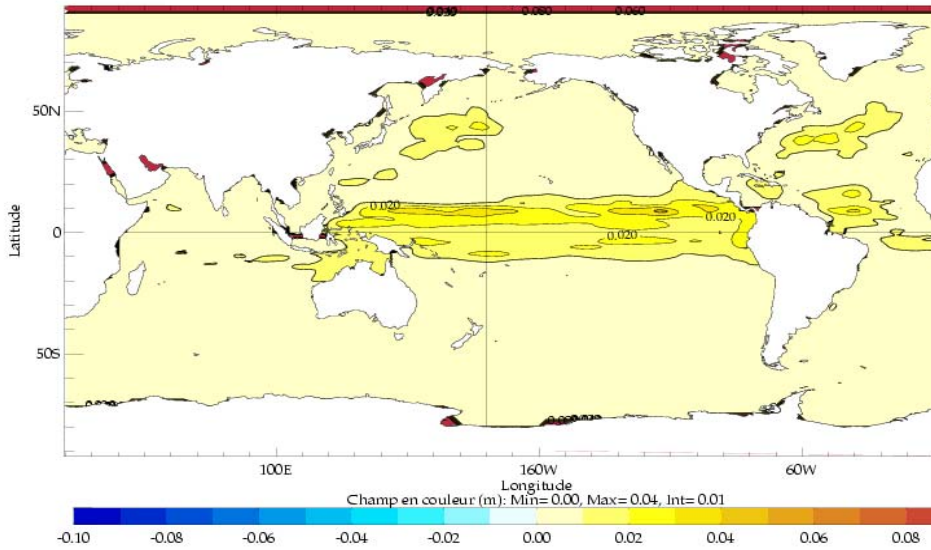
Data coverage for May 2002



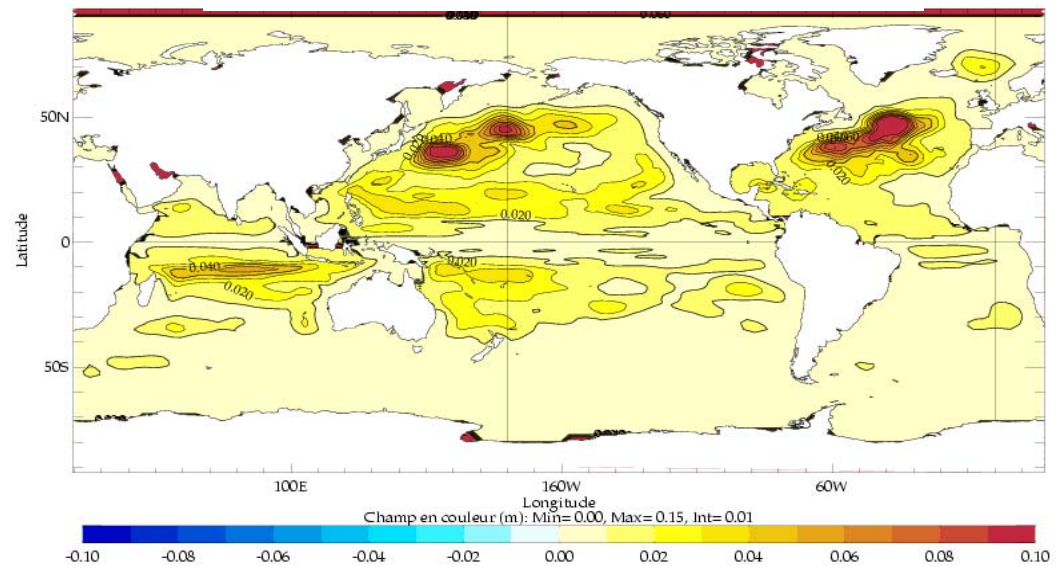
Data coverage for May 2003



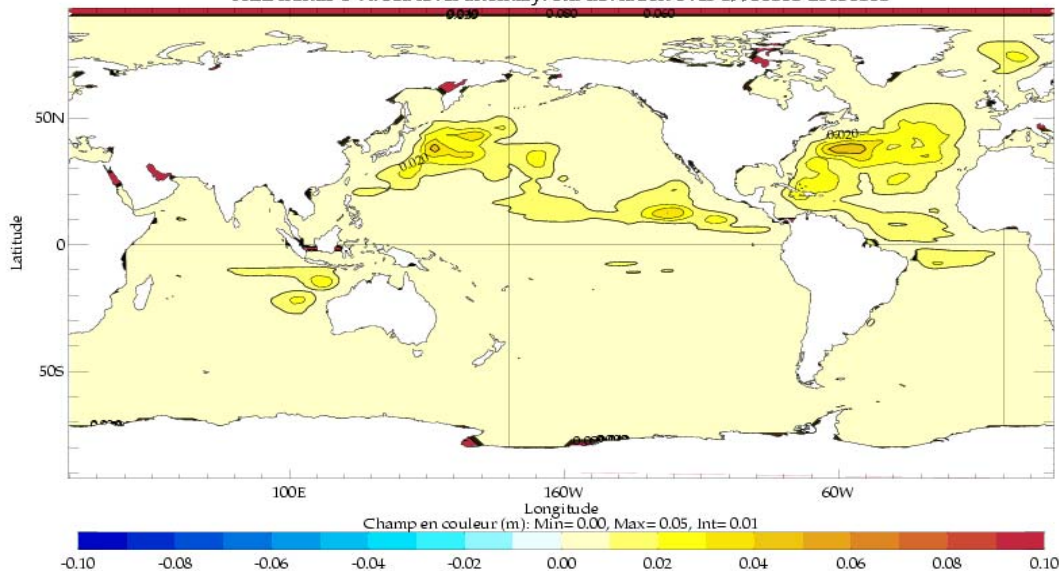
RMS: ALL - Moorings



RMS: ALL - XBT



RMS: All-ARGO



Impact of Different Data types

Variable: Surface Elevation

Time Period: 1998-2003

From A. Vidard

What Next?

Longer time scales

- Era40 offers the possibility of longer ocean Reanalysis:
 - a) Calibration for Seasonal Forecasts
 - b) Decadal Predictions
- Decadal variability is an issue:
 - Controlling model error
 - Assimilation in the deep ocean

Shorter time scales

- Monthly Forecasts already requires a more timely ocean analysis
Accelerated ocean analysis
- Possibility of using a coupled model for Medium Range Forecast
How to initialize the ocean component?

ENACT:

ENhance ocean data Assimilation and ClimaTe

- (A). Improve and extend ocean data assimilation systems, and apply them to produce global ocean analyses over a multi-decadal period.
- B). Quantify the benefits of the enhanced assimilation systems through retrospective seasonal and multi annual climate forecasts and through analysis of ocean behaviour in a multi-model framework

ENACT

- Forcing fields from ERA40
- Comprehensive Ocean Obs data set + QC procedure (provided by the UKMO)
- Development and Implementation of different Ocean Data assimilation Methods (3D-var, 4-Dvar, EnKF...)

It will allow the comparison of assimilation methods

- Development of diagnostics for Ocean D-A
- First trial of decadal forecasts with coupled models

ENACT: Framework of ODA systems

<u>OGCM</u>	<u>scheme</u>	<u>partner</u>
OPA	4d-var 3d-var	CERFACS/LODYC
OPA	generic EnKF	NERSC/ECMWF
OPA	Ol-type (SOFA)	INGV
C-HOPE	4d-var	MPIM
C-HOPE E-HOPE	generic EnKF 4d-var	KNMI
E-HOPE	Ol-type	U-READING/ECMWF
GLOSEA	Ol-type	Met Office

Summary

- Ocean data assimilation improves the state estimate of the tropical ocean
- Initialization of ocean models by means of data assimilation improves the forecast of ENSO
- The existence of model/forcing error in ocean data assimilation can not be ignored.
- The importance of balance relationships has been illustrated
 - The operational environment is ideal to spot problems
- The choice of a data assimilation method can not be independent of the application

Future Directions

Method:

- Implementation of OPA-VAR at ECMWF
 - Treatment of model error
 - Comparison of 4Dvar/EnKF
 - Development of internal diagnostics

Observations:

- **Assimilate in Situ Temperature/ SL /Salinity**
- **OSE experiments**
- **Further diagnostics**

Questions:

1. **Feedback to the atmospheric community?**
Should the atmospheric reanalysis include an ocean model in the future?
2. **Different methods for different problems?**
Initializing the ocean for NWP versus Seasonal