

Theories of low-frequency variability over Europe

Brian Hoskins

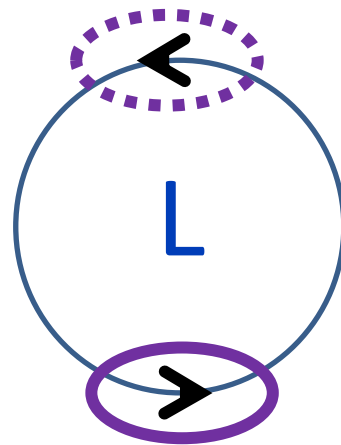
**Department of Meteorology, University of Reading
Grantham Institute for Climate Change, Imperial College**

The Agenda

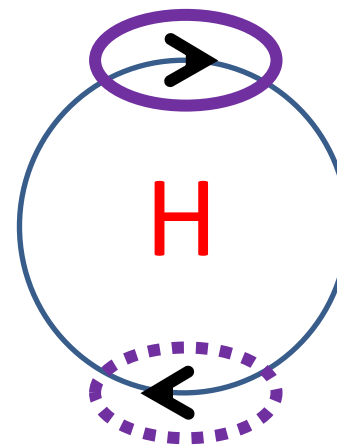
1. Synoptic eddy organisation & feedback
2. Rossby waves – horizontal propagation
3. Teleconnection patterns
4. Blocking
5. Events related to the Asian Summer Monsoon

1 Synoptic weather systems

For zonal flows and low frequency anomalies, synoptic weather systems
Are organised by the westerly flow
Feedback + on the barotropic motion, - on the baroclinic motion



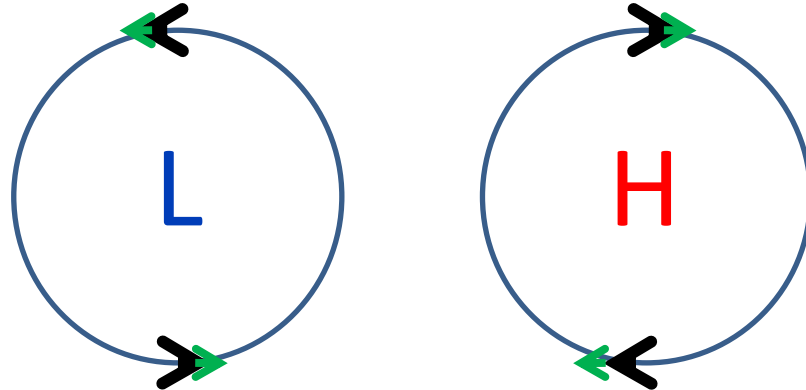
More synoptic
activity



Less synoptic
activity

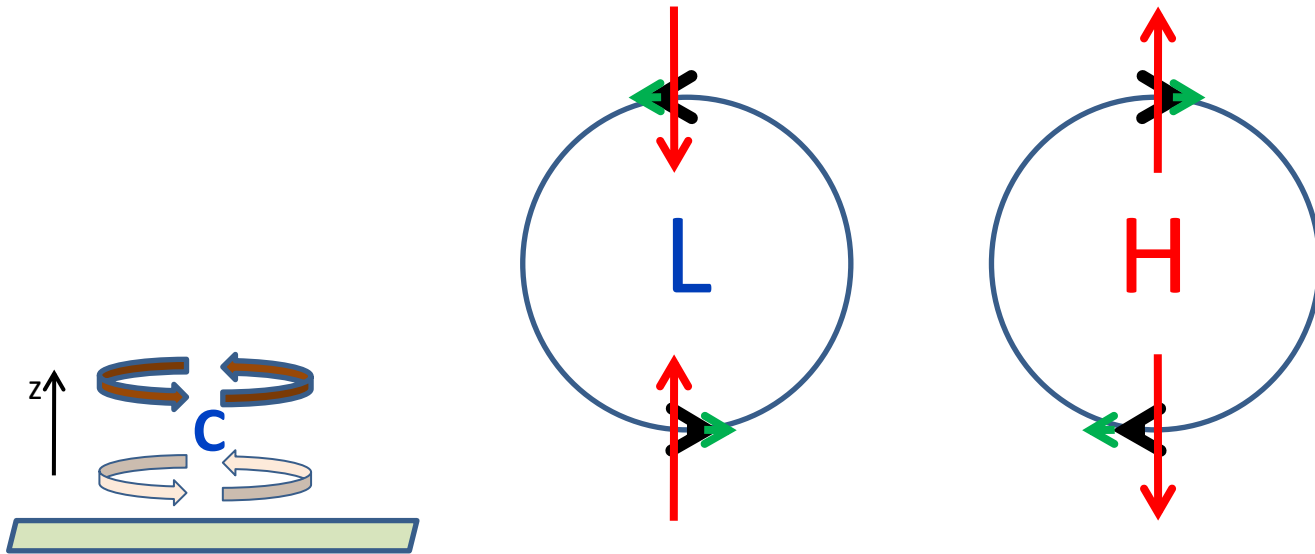
Synoptic weather systems

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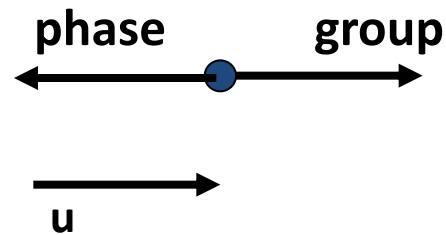
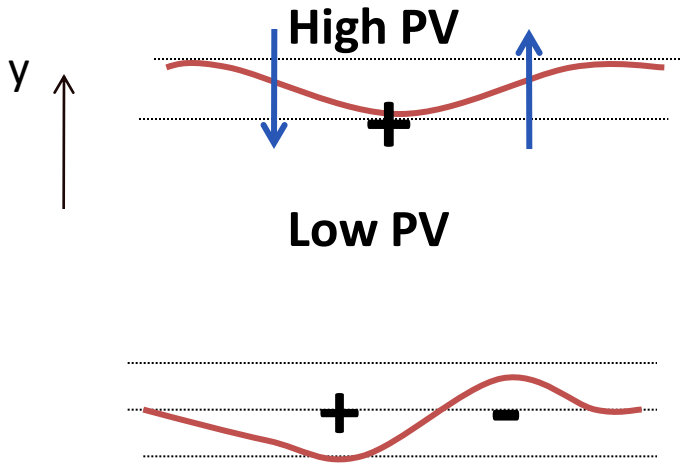
Synoptic weather systems

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Feedback + on the barotropic motion, - on the baroclinic motion



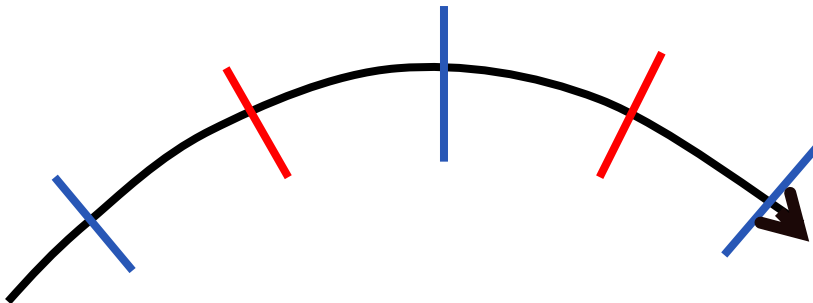
Therefore they act to counter the spin-down of the lower frequency anomalies due to surface friction

2. Rossby wave propagation - horizontal



Stationary waves possible on westerlies
Influence to the east

On the sphere and refracted by the ambient flow

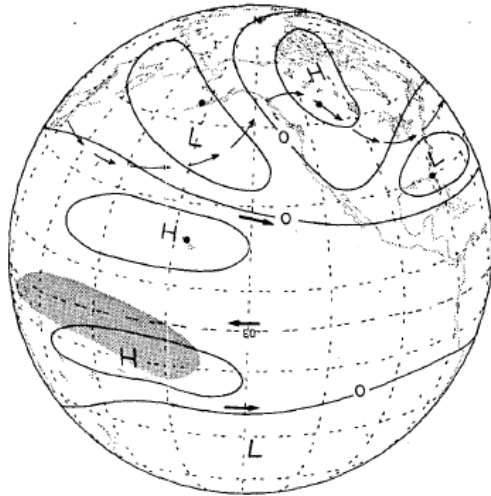


Strong jets can act as waveguides

Eddies can take barotropic energy
from the mean in jet exit regions

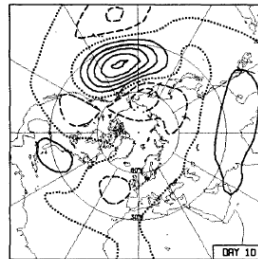
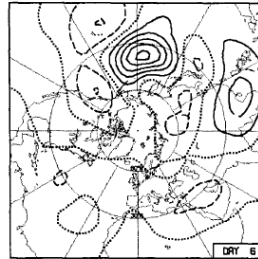
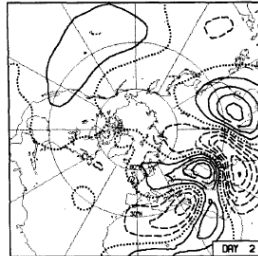
Forcing and Propagation of Rossby waves

Observations



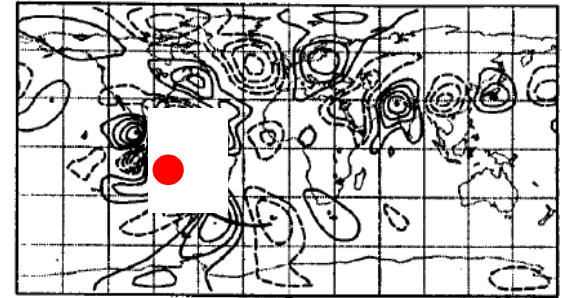
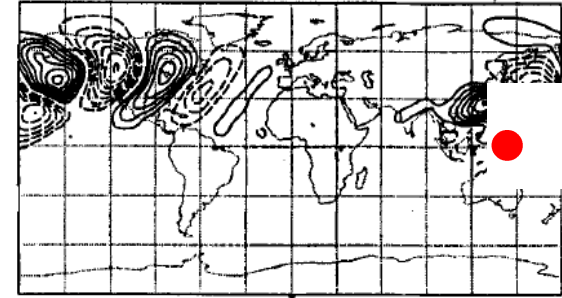
Horel and Wallace (1981)

Initial perturbation
barotropic model
2-D basic state



Simmons, Wallace
& Branstator (1983)

Forced baroclinic model
3-D basic state
After 9 days



Ambrizzi and Hoskins (1997)

$$(\partial_t + \mathbf{v}_\psi \cdot \text{grad}) \xi + \beta v_\psi = -\zeta D - \mathbf{v}_\chi \cdot \text{grad} \zeta$$

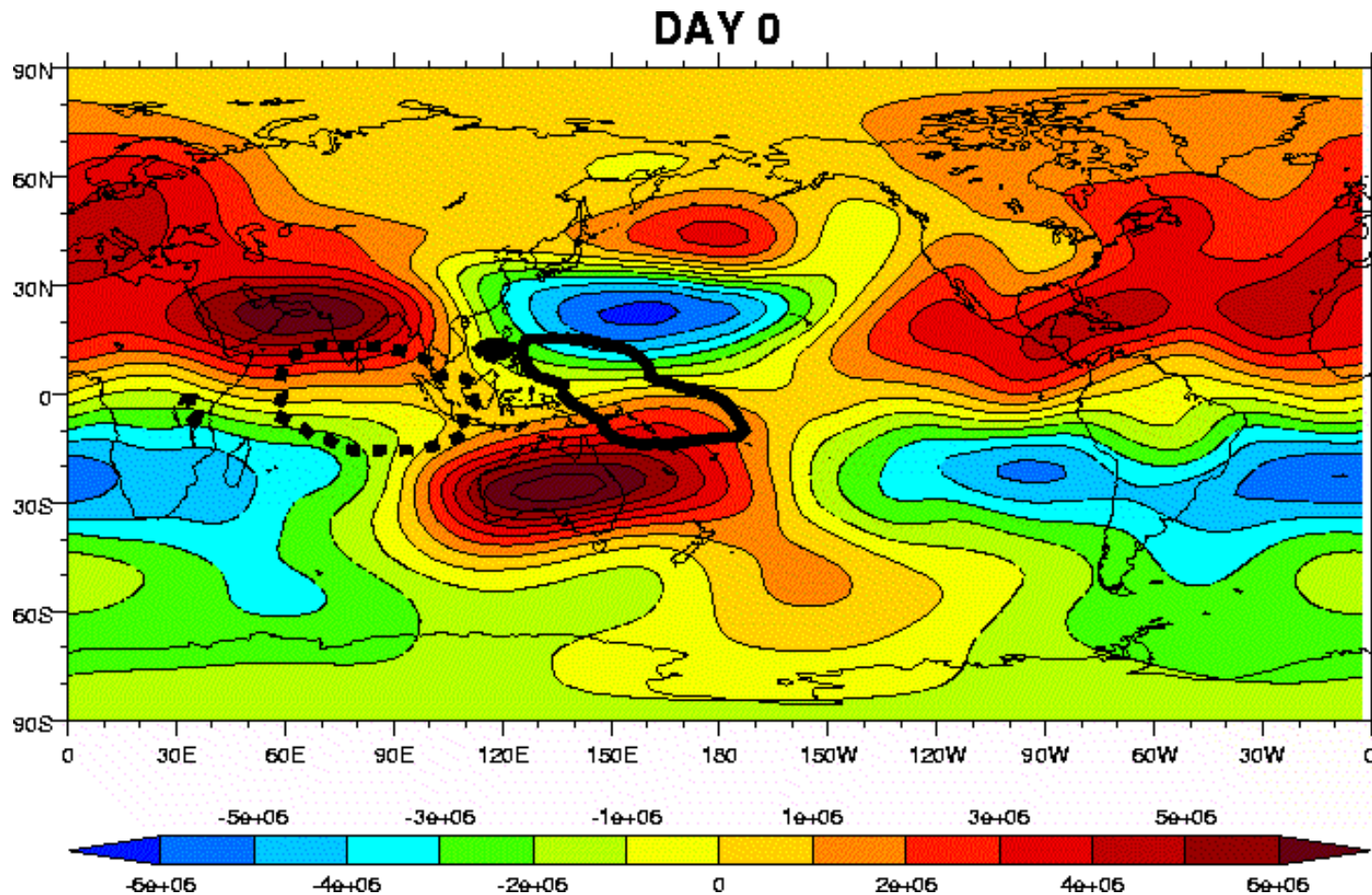
Rossby wave source

Sardeshmukh & Hoskins (1988)

DJF global circulation anomalies associated with an MJO cycle

Cycle from 2 EOFs of 20-200 day filtered OLR: heavy contours

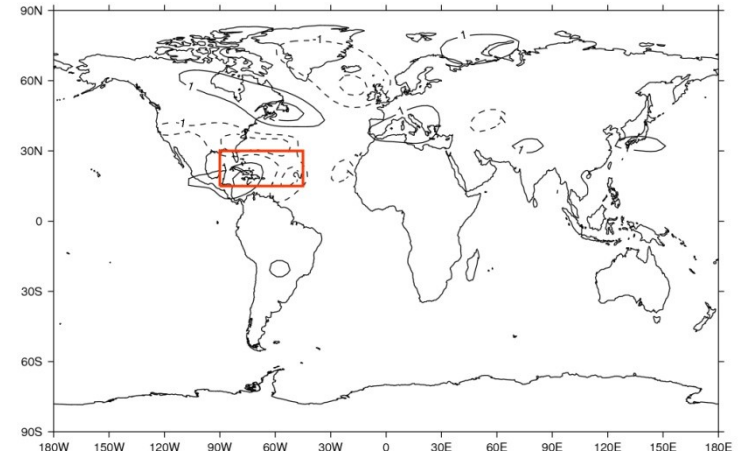
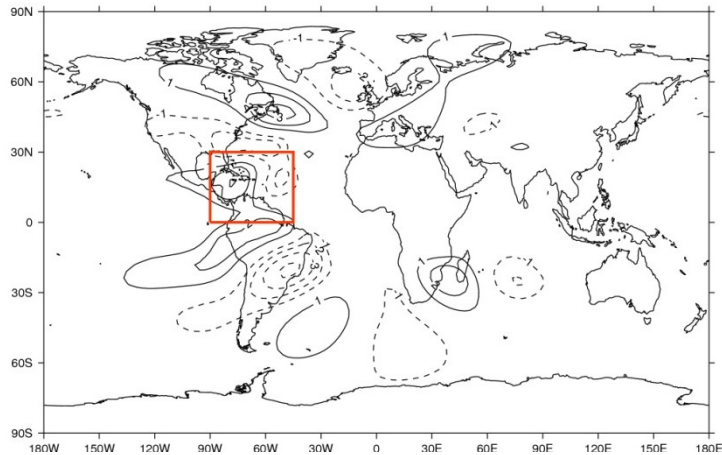
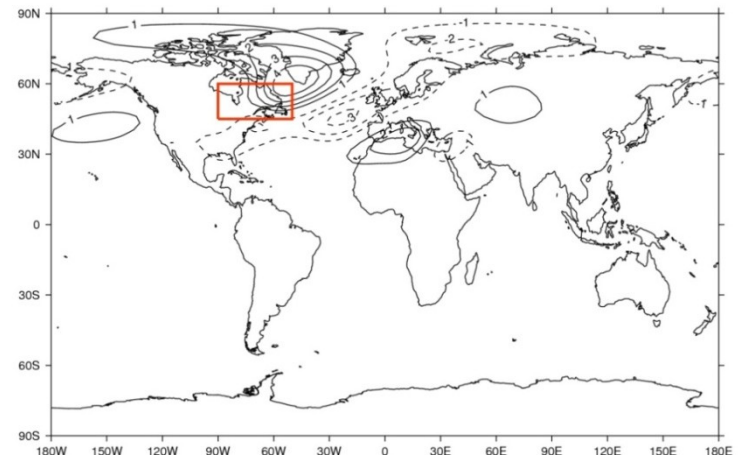
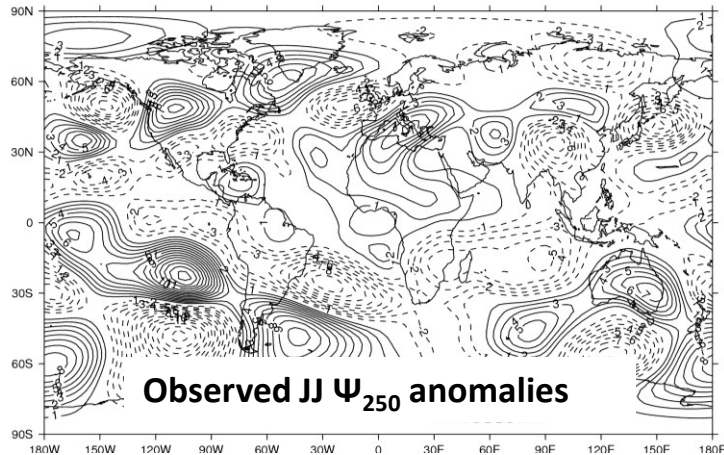
Regressed ψ_{200} : colours



Summer 2007: experiments with a time dependent 3-D baroclinic model

Ricardo Fonseca

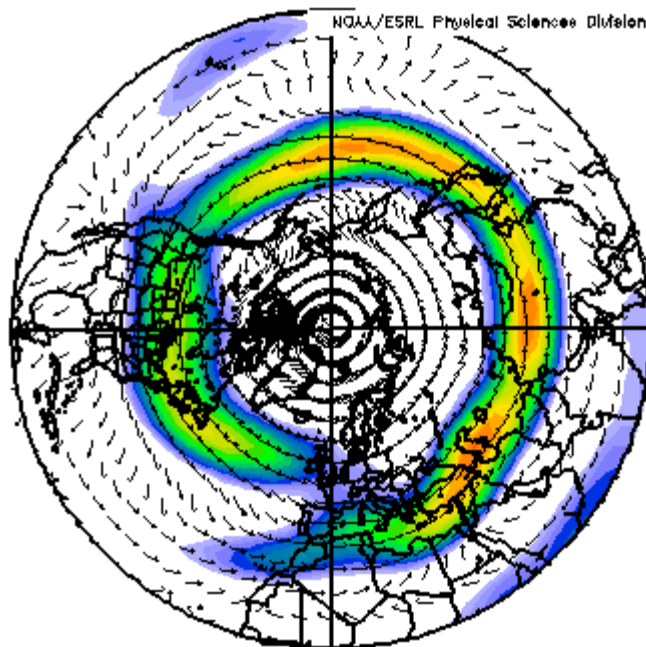
Damping only. Climatological flow specified and balanced by a forcing.
Relaxation to observed anomaly in specified regions: show Ψ_{250} anomaly at day 30



Summer 2007 UK floods - 250hPa v

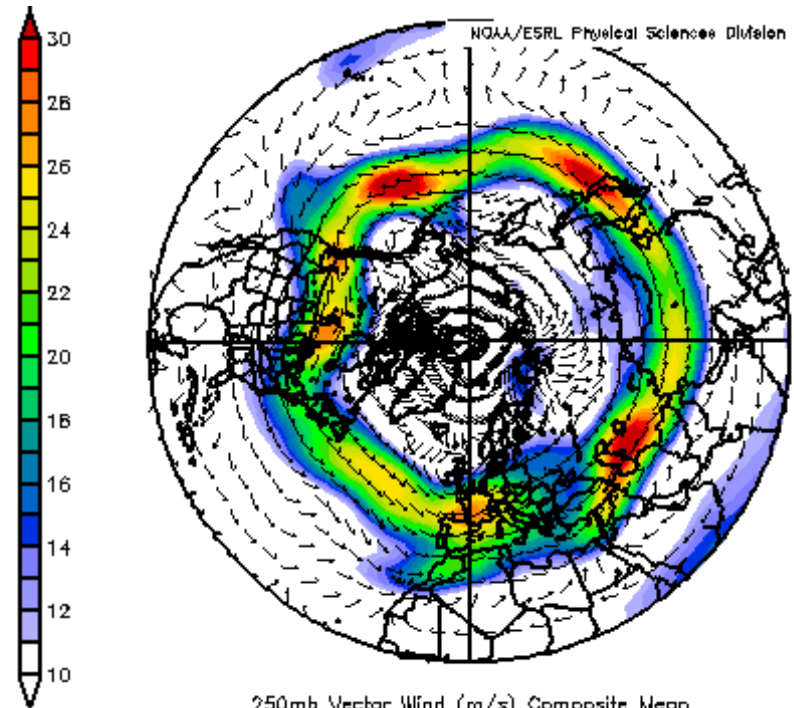
Average from 12 June to 25 July

Climatology



250mb Vector Wind (m/s) Climatology (1968-1996 Climatology)
6/12 to 7/25

2007



250mb Vector Wind (m/s) Composite Mean
6/12/07 to 7/25/07

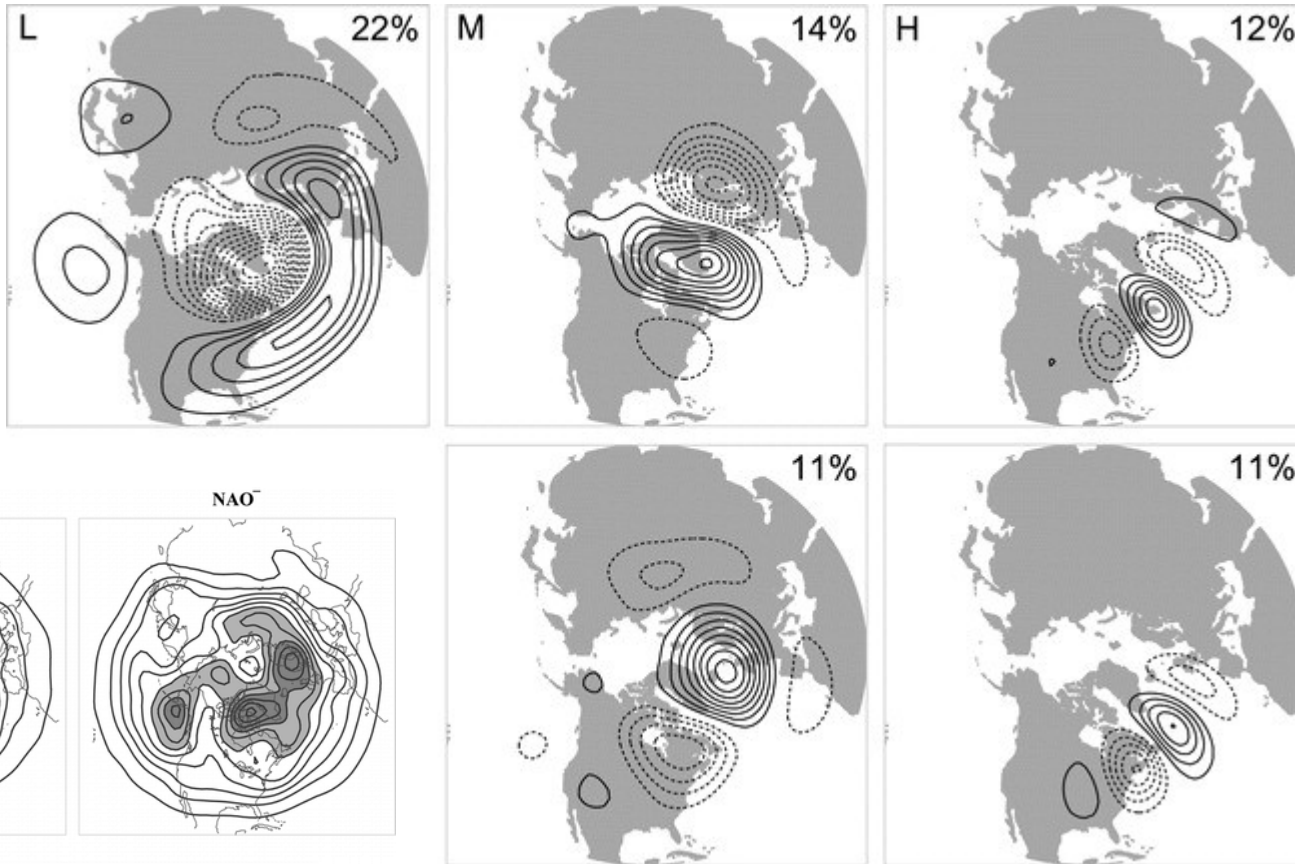
3. Leading DJF Z_{500} EOFs in the Atlantic Sector

Rennert & Wallace (2009)

> 30d

6-30 d

<6d

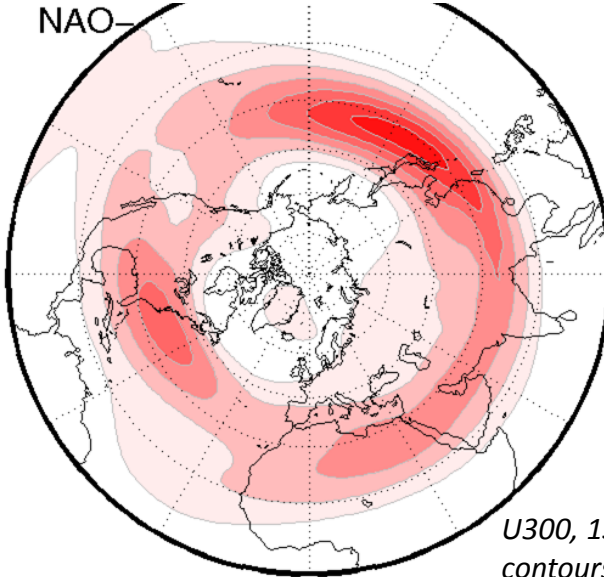
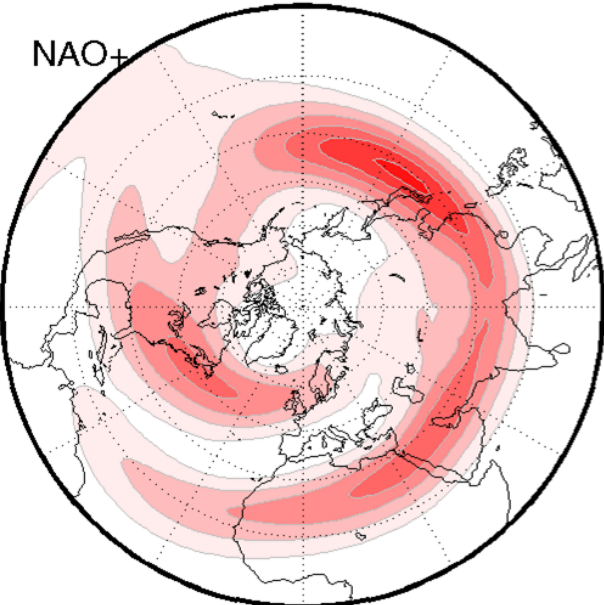


NAO⁺

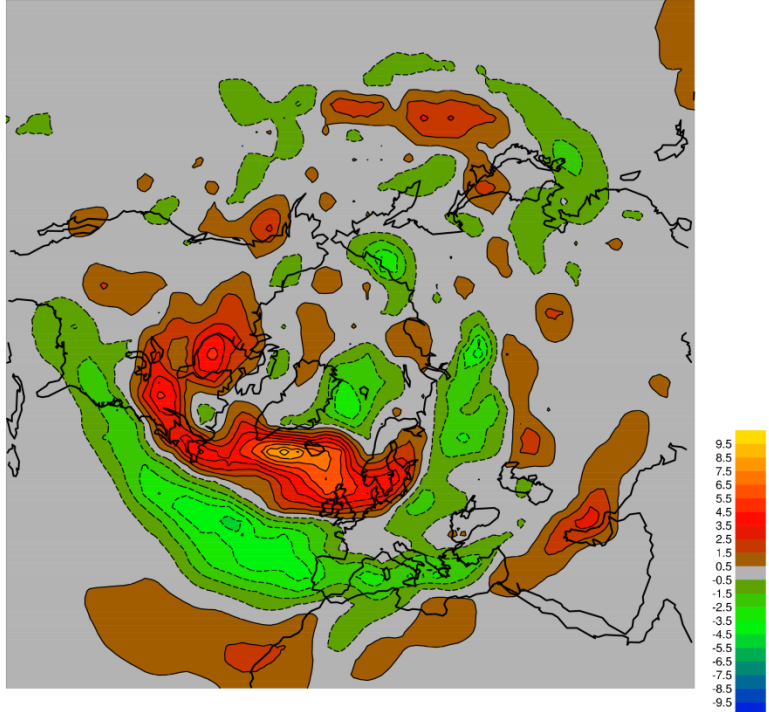
NAO⁻

7000 8500 10000 11500 13000 14500 16000 17500

The North Atlantic Oscillation (NAO) describes synchronous variations in the strength and orientation of the jet and storm track.



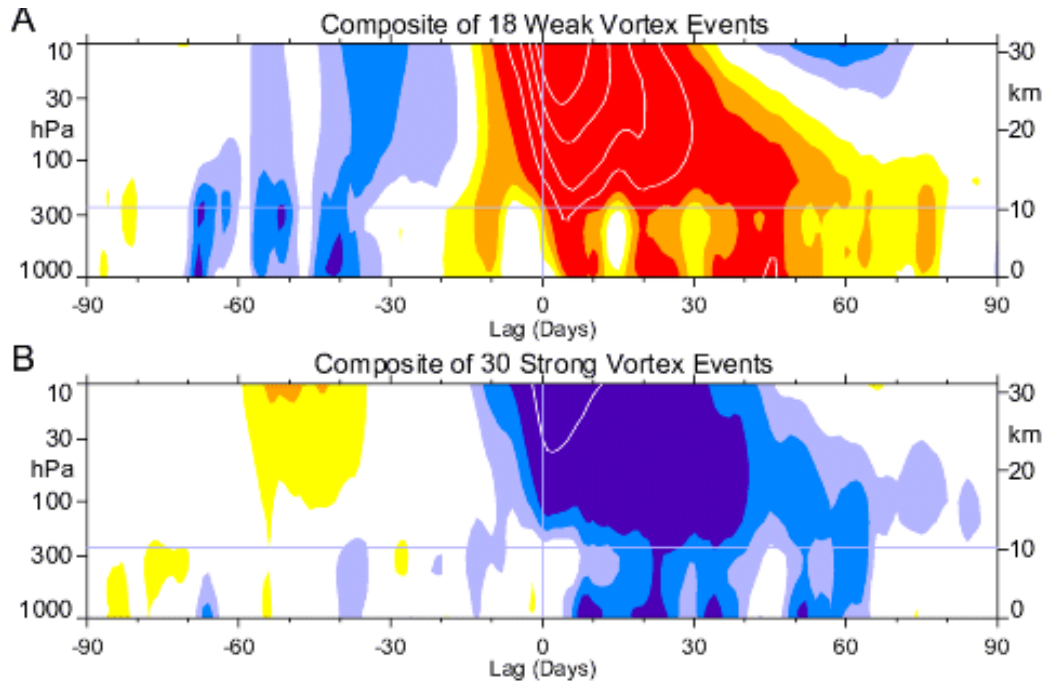
U300, 1SD daily comps, contours 10m/s



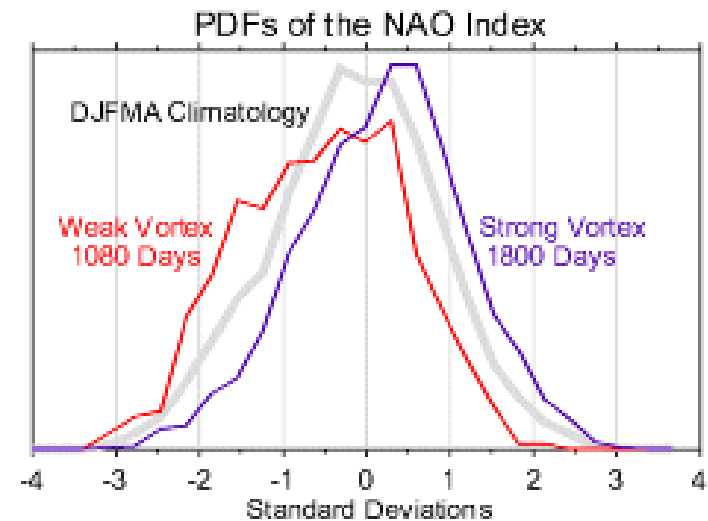
Track density regressed on NAO index; Hoskins & Hodges ?

The stratospheric connection

Composites of EOF1/annular behaviour for weak and strong 10 hPa vortex events



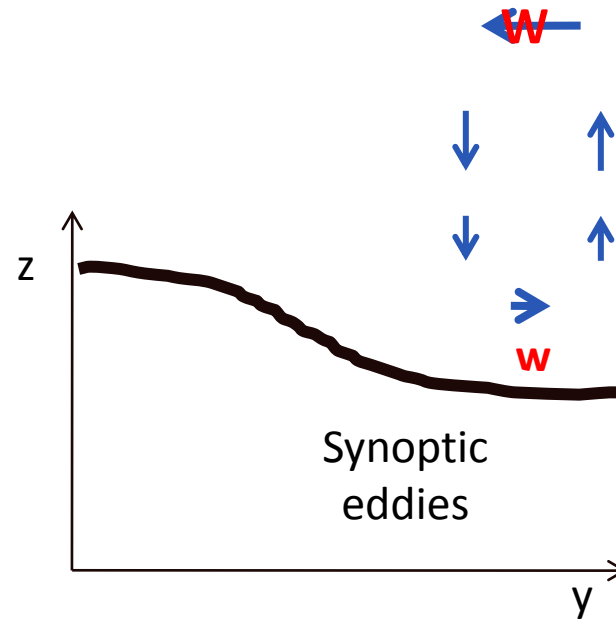
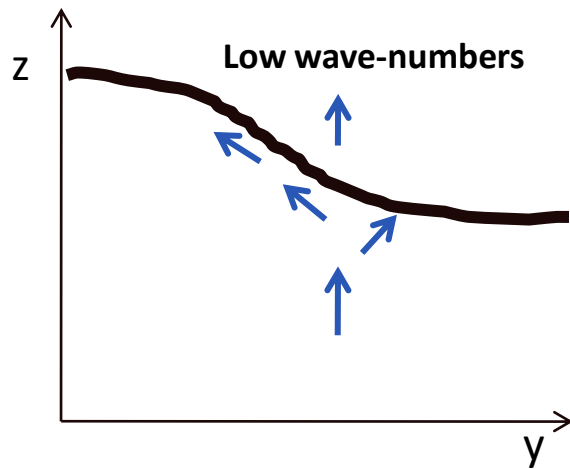
PDFs of the NAO index

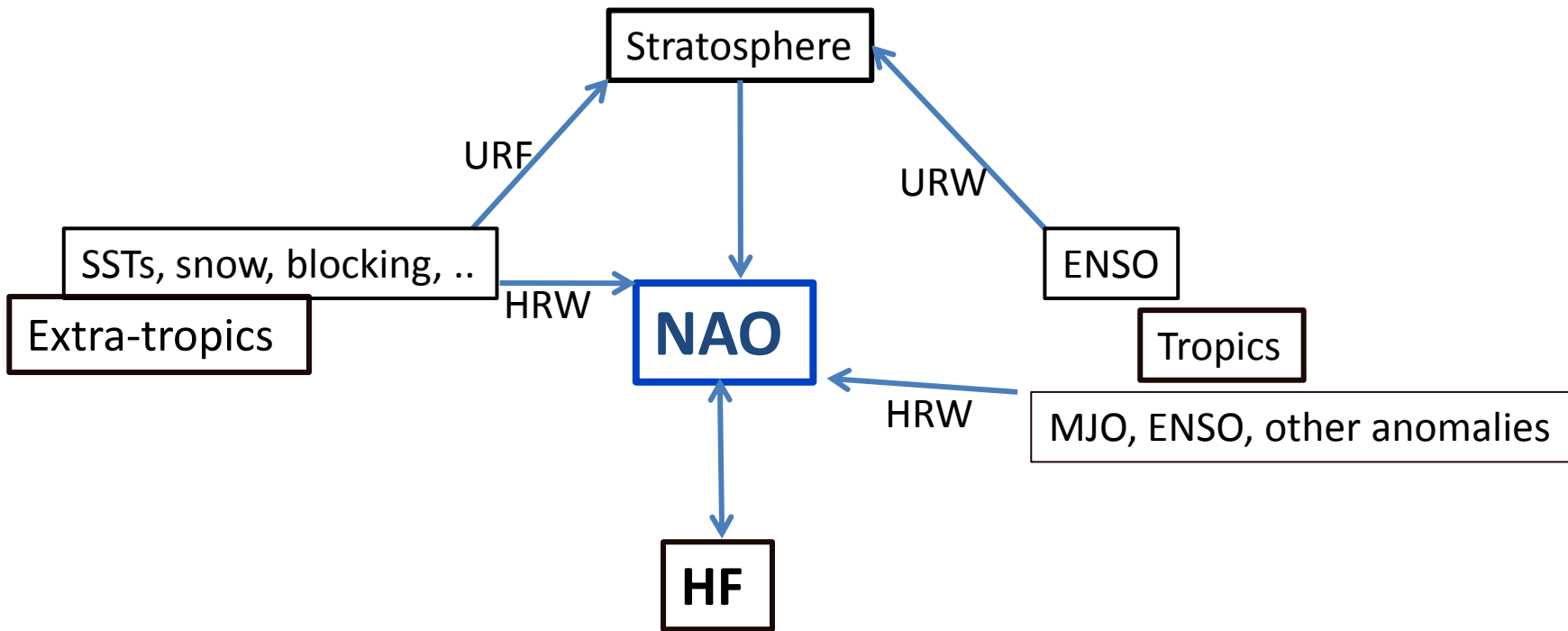


Baldwin & Dunkerton (2001)

Troposphere-Stratosphere Interaction

Vertical propagation
of Rossby Waves





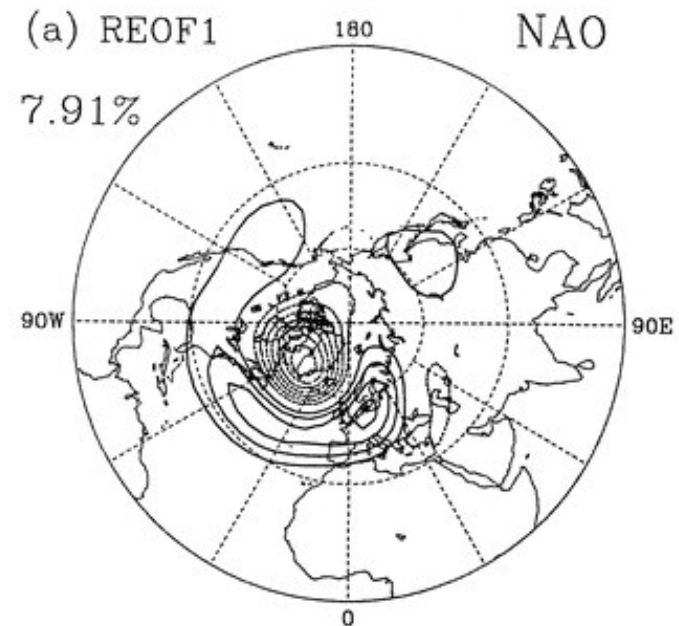
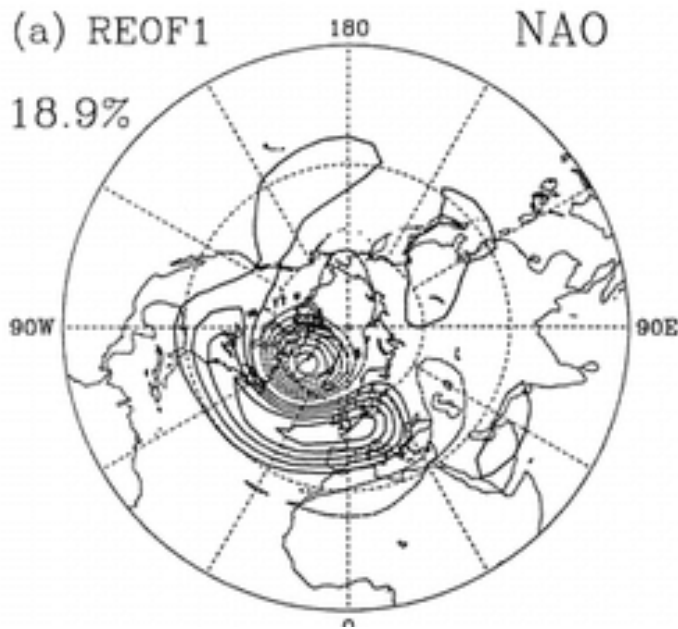
Do monthly-seasonal teleconnection patterns reflect dynamics on these time-scales?

Feldstein (2000)

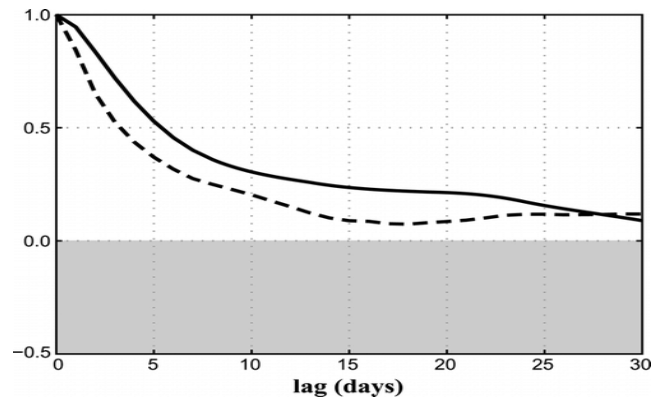
Rotated EOFs

Seasonal data

Daily data



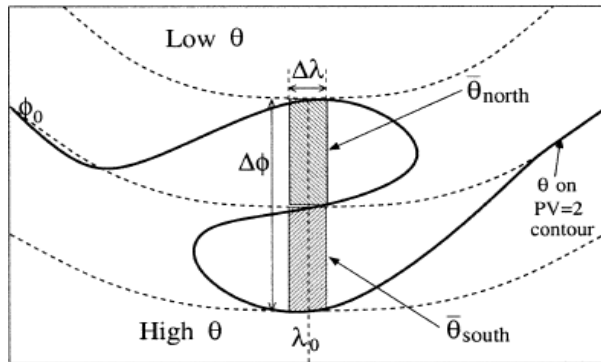
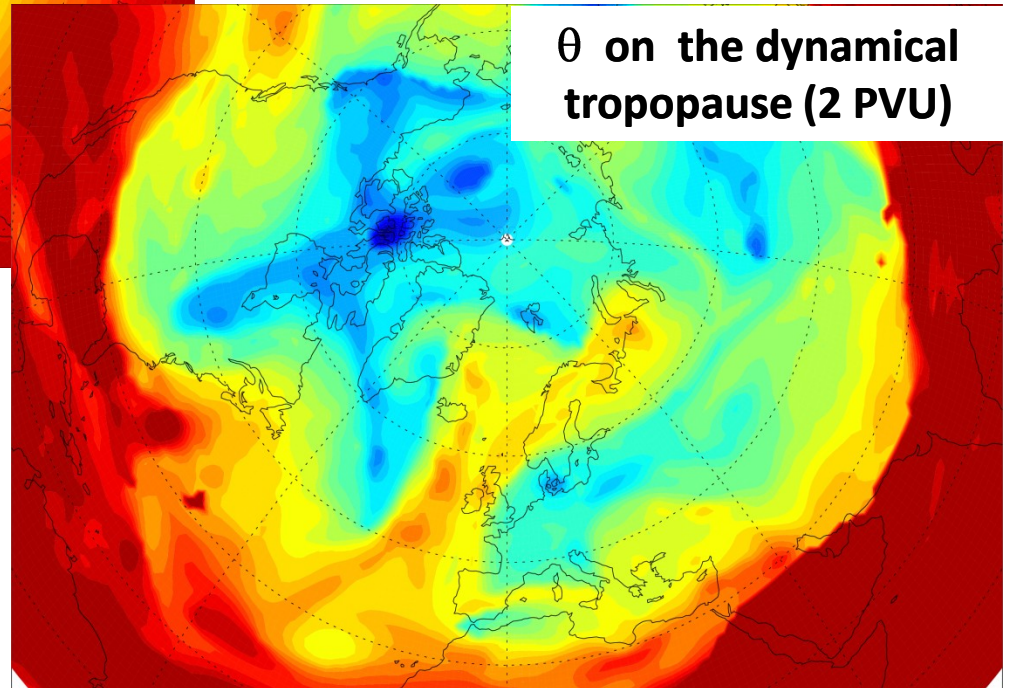
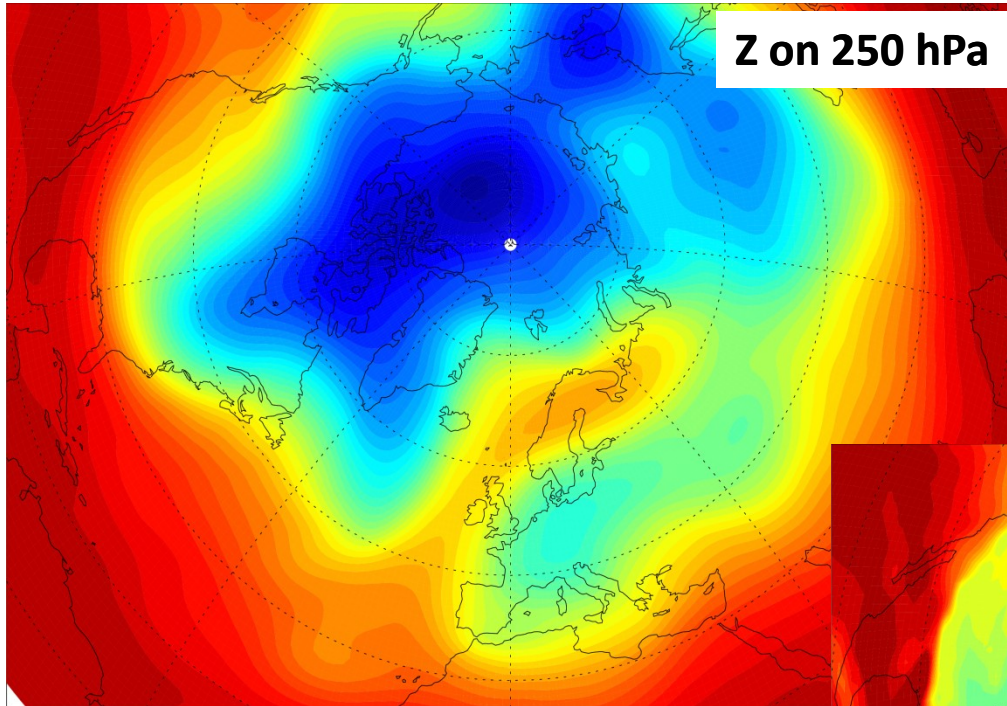
Autocorrelation
of daily NAO



4. Blocking

A Typical European Block: 20 November 1993 12 UTC

Tyrlis & Hoskins JAS 2008

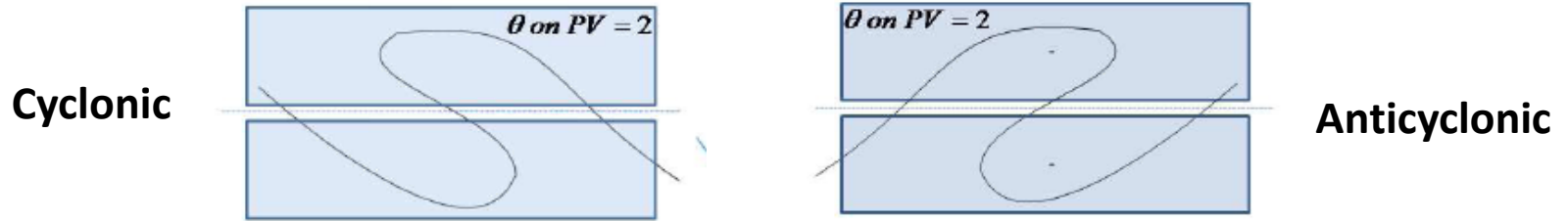


$$B = \theta_{\text{north}} - \theta_{\text{south}}$$

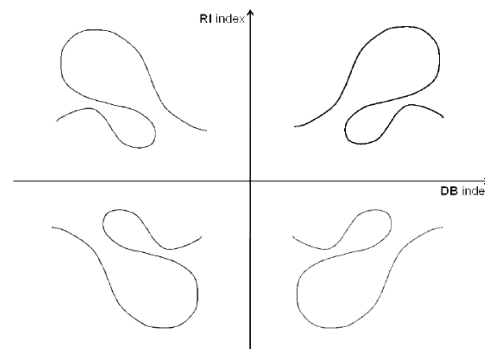
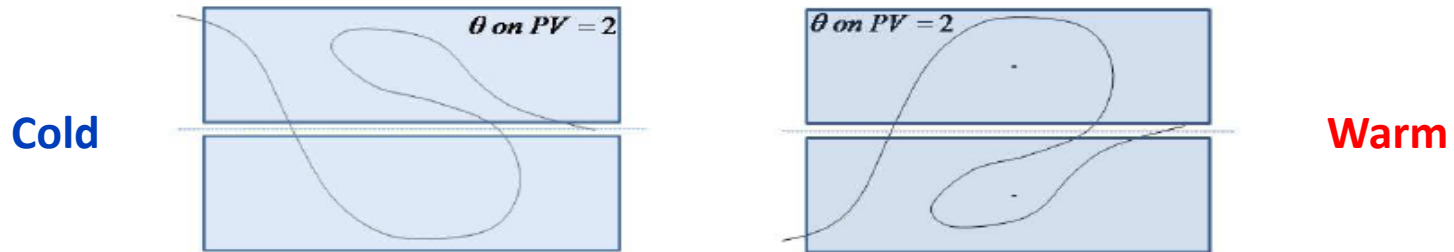
Further Blocking Diagnostics

Giacomo Masato

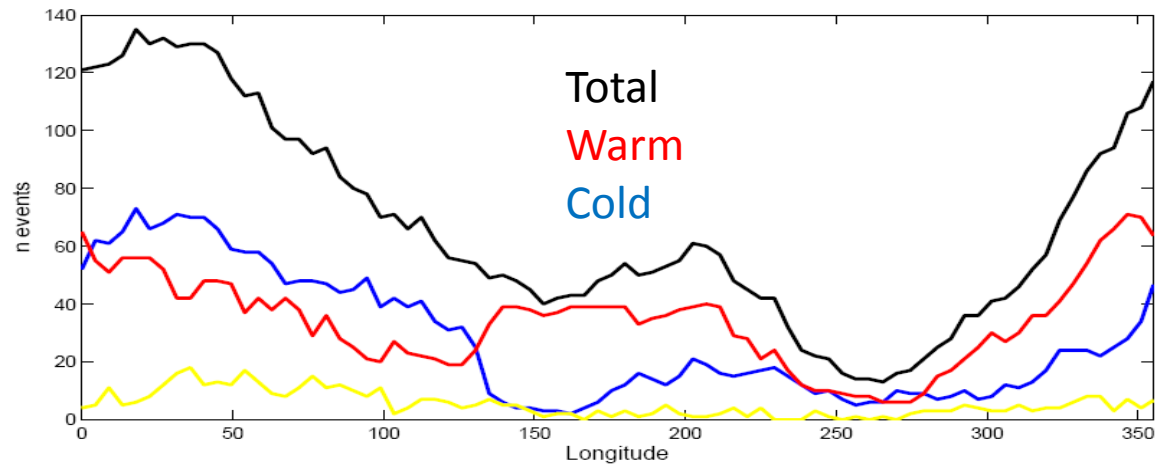
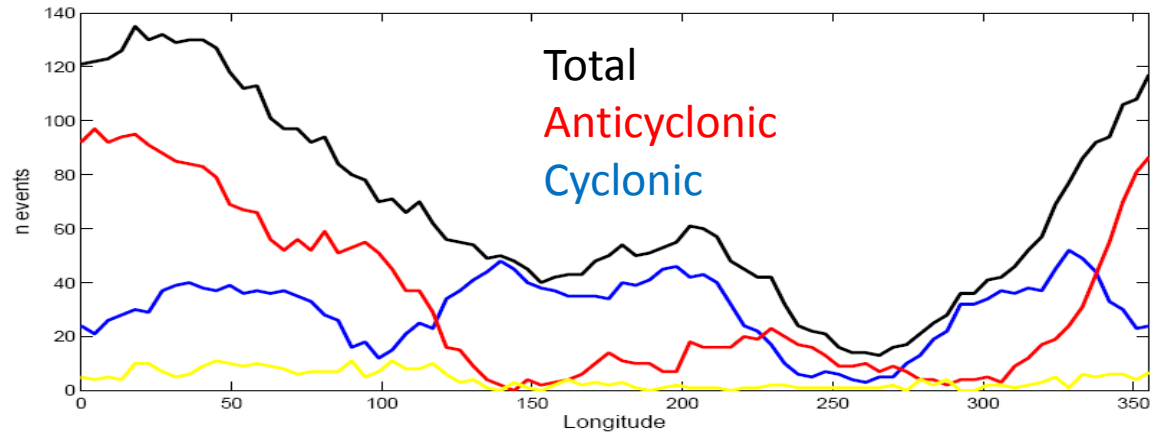
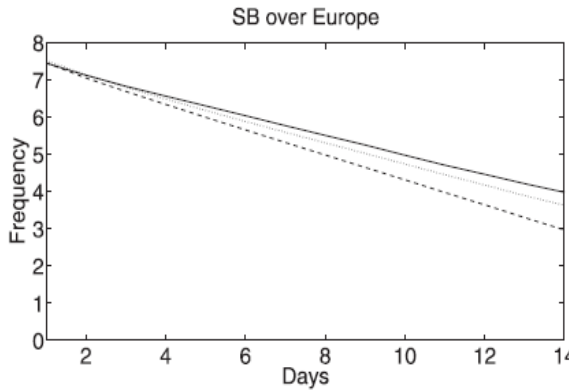
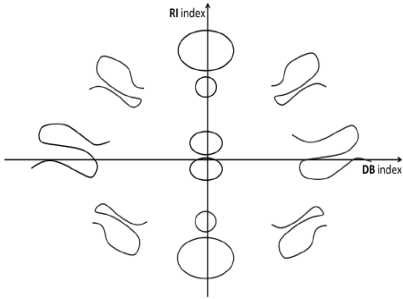
Direction of wave-breaking



Relative Intensity



Blocking frequencies with longitude



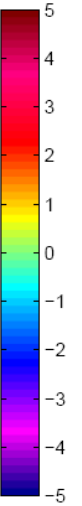
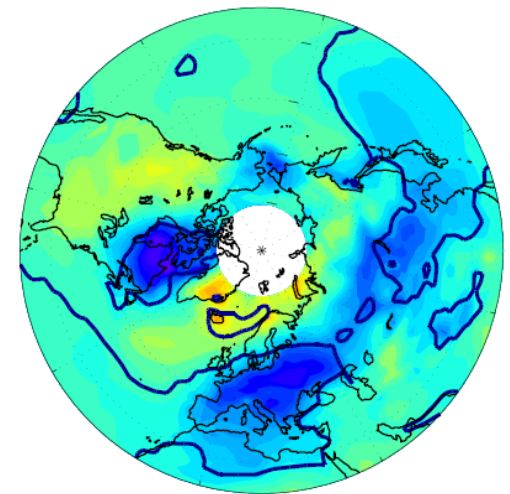
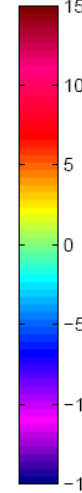
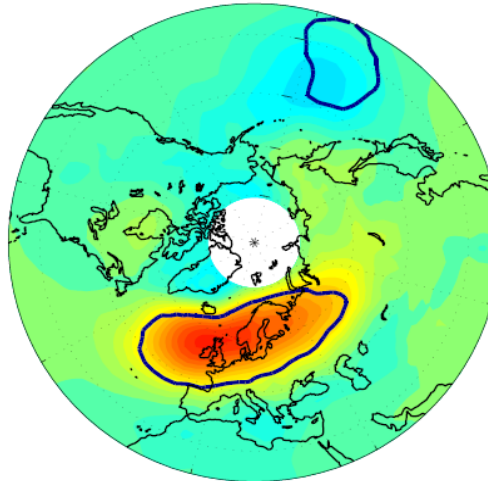
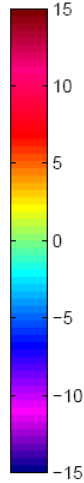
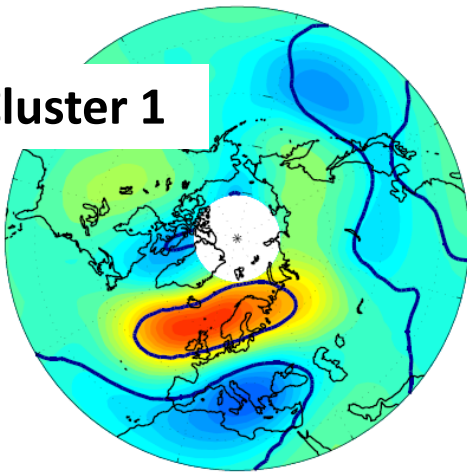
Cluster analysis for European Winter Blocking

Θ on PV2

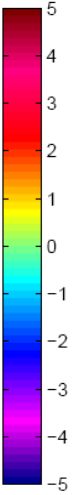
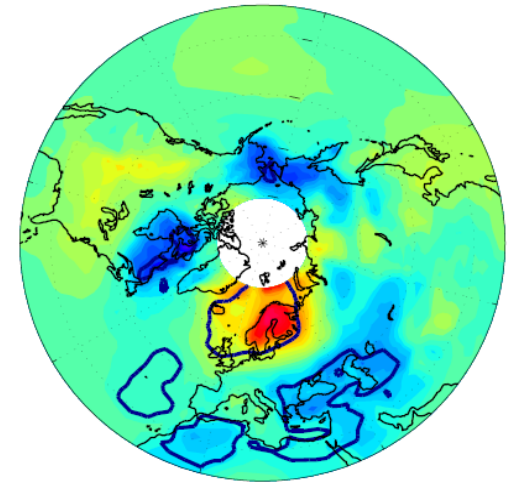
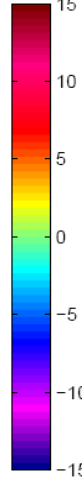
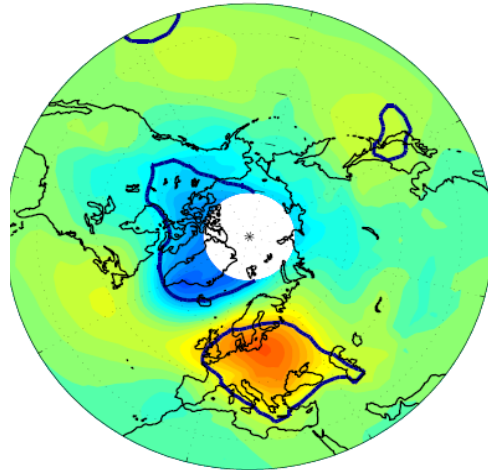
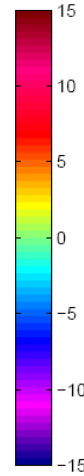
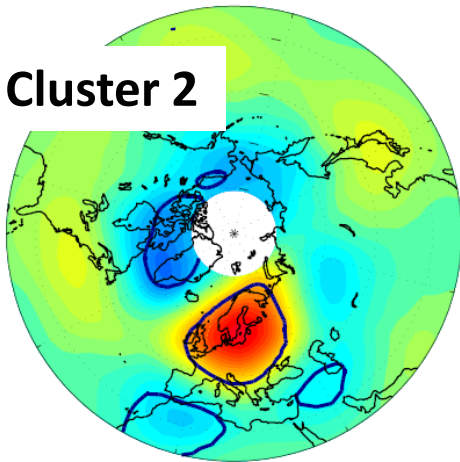
p_{msl}

T_{2m}

Cluster 1

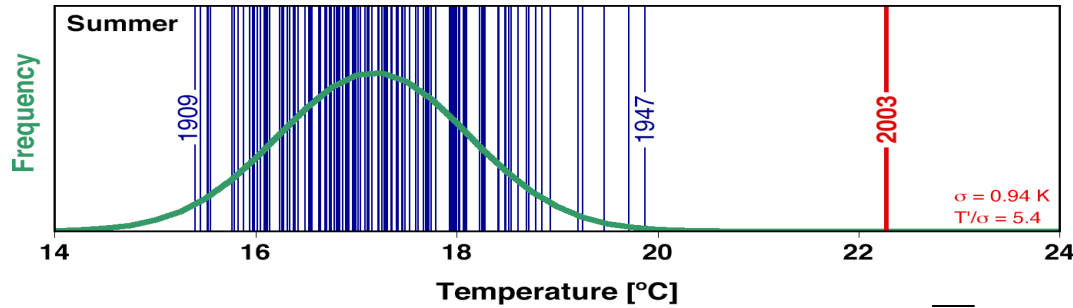


Cluster 2



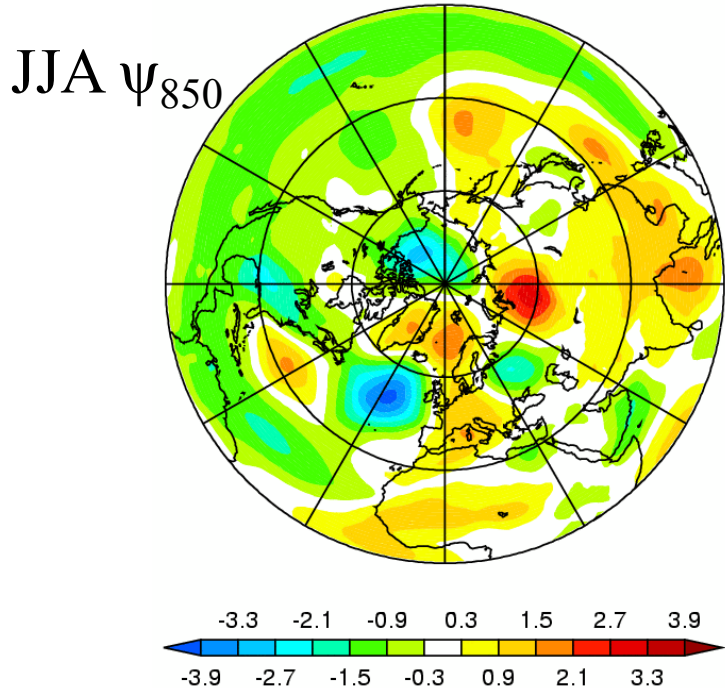
Giacomo Masato

Summer 2003: record European warmth



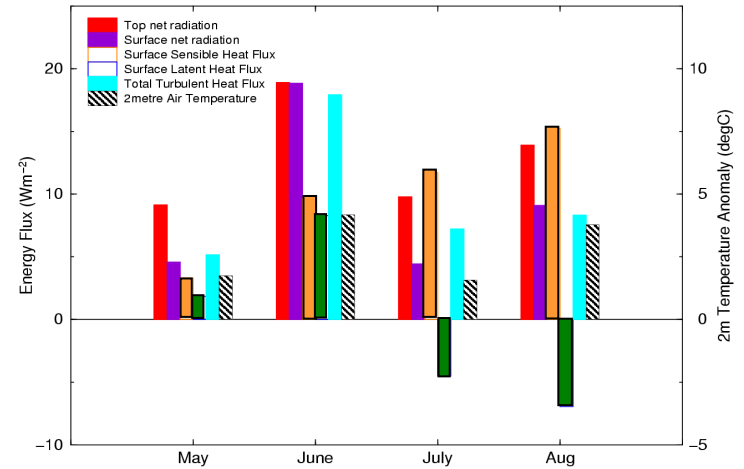
Swiss Temperature Series 1864-2003
(mean of 4 stations)

Schär et al. 2004, Nature, **427**, 332-336



Anomalies in Heat Budget over Europe

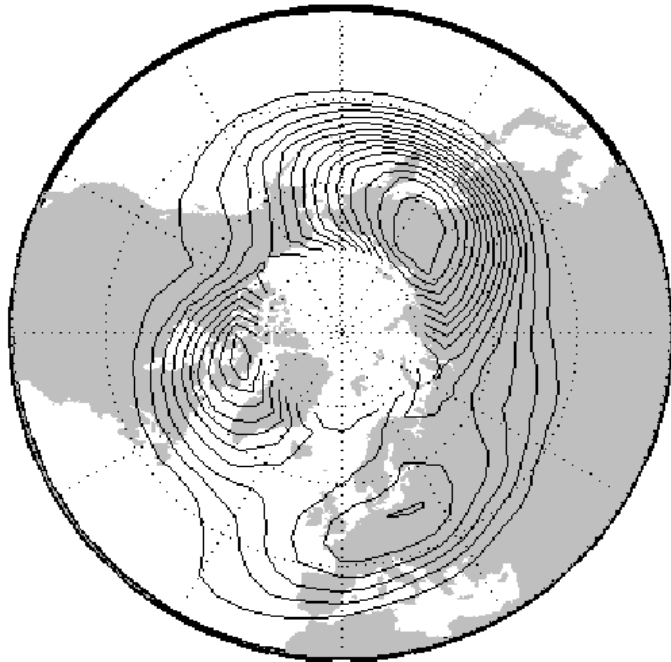
surface sensible heat flux surface latent heat flux



Black, Blackburn, Harrison, Methven & Hoskins (2004)

Wave-Breaking Frequencies Dec-Feb

Mean Episode Frequency



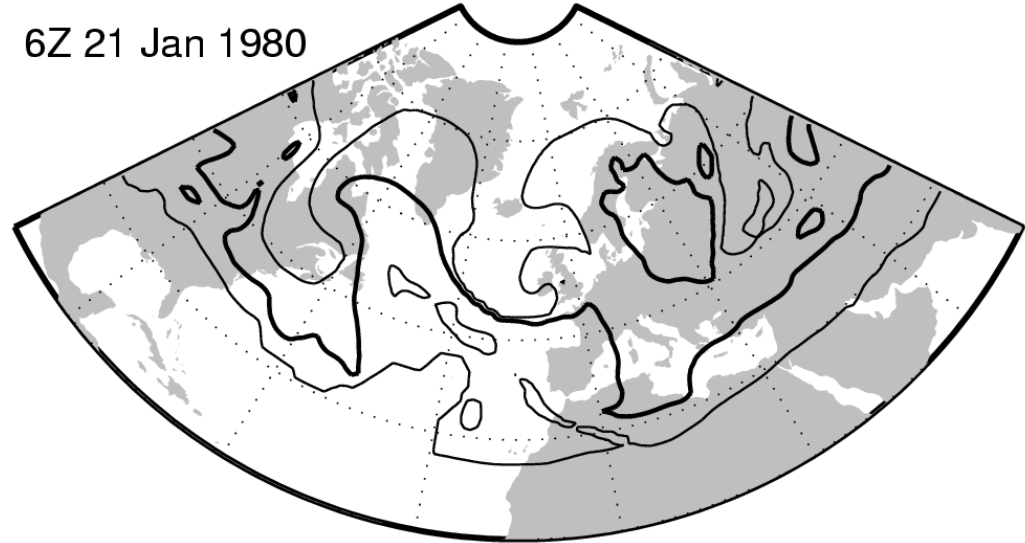
Mean

(c.i. 0.05)

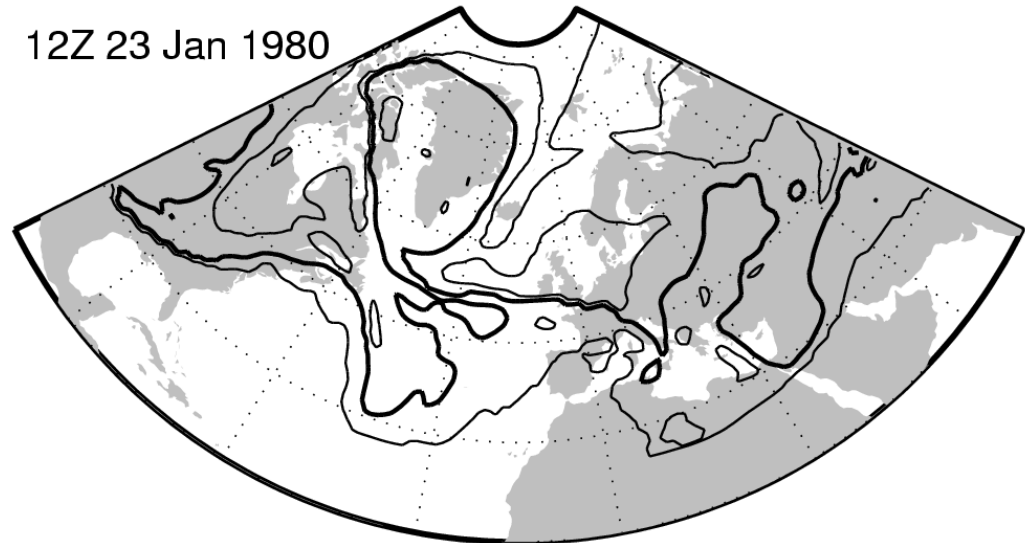
Woollings et al JAS 2008

The onset of a
typical Rossby
wave-breaking
event

6Z 21 Jan 1980



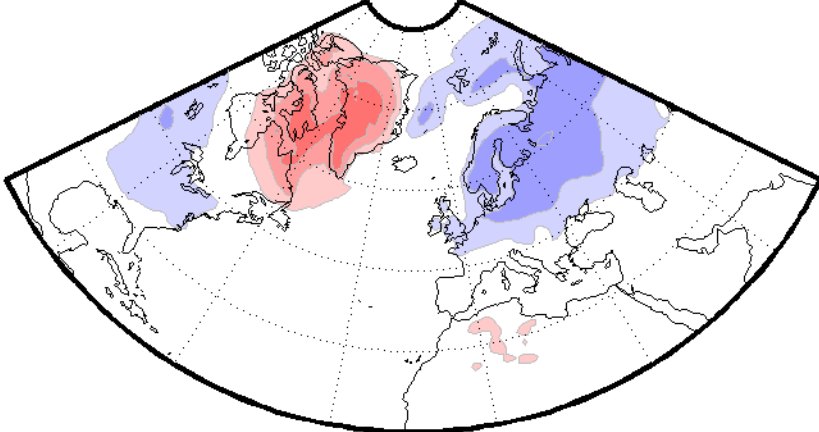
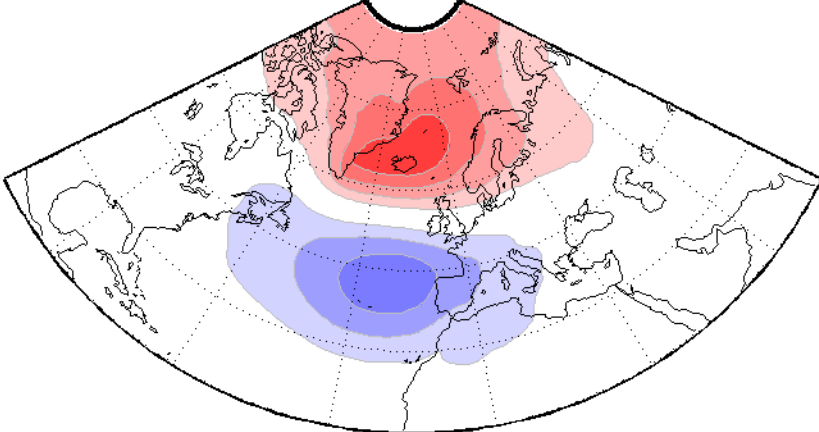
12Z 23 Jan 1980



Contribution of Wave-Breaking to NAO- Anomalies

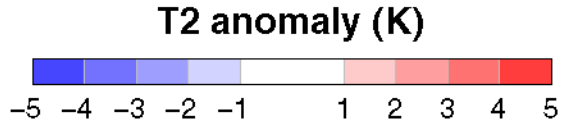
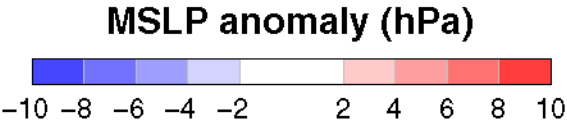
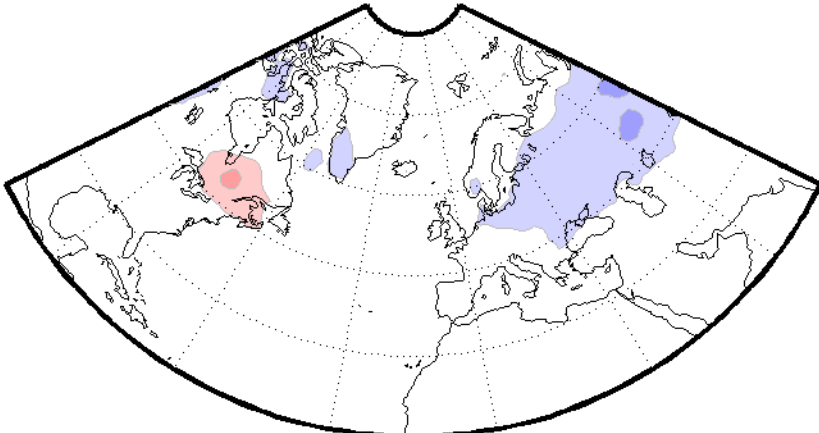
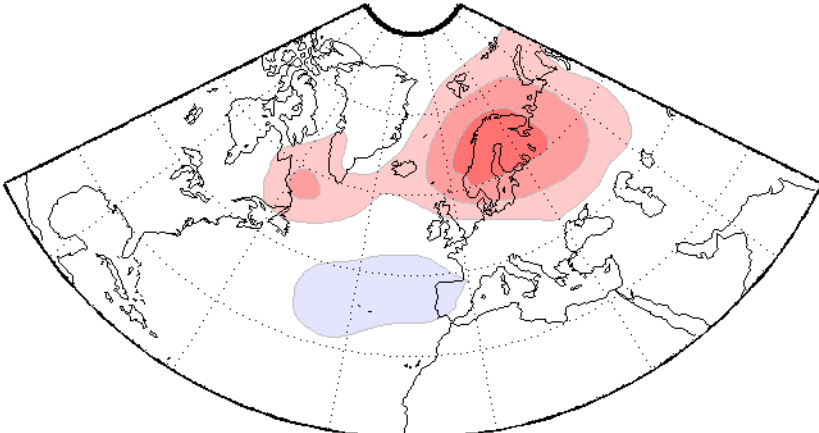
45 NAO- DJF months

45 NAO- DJF months

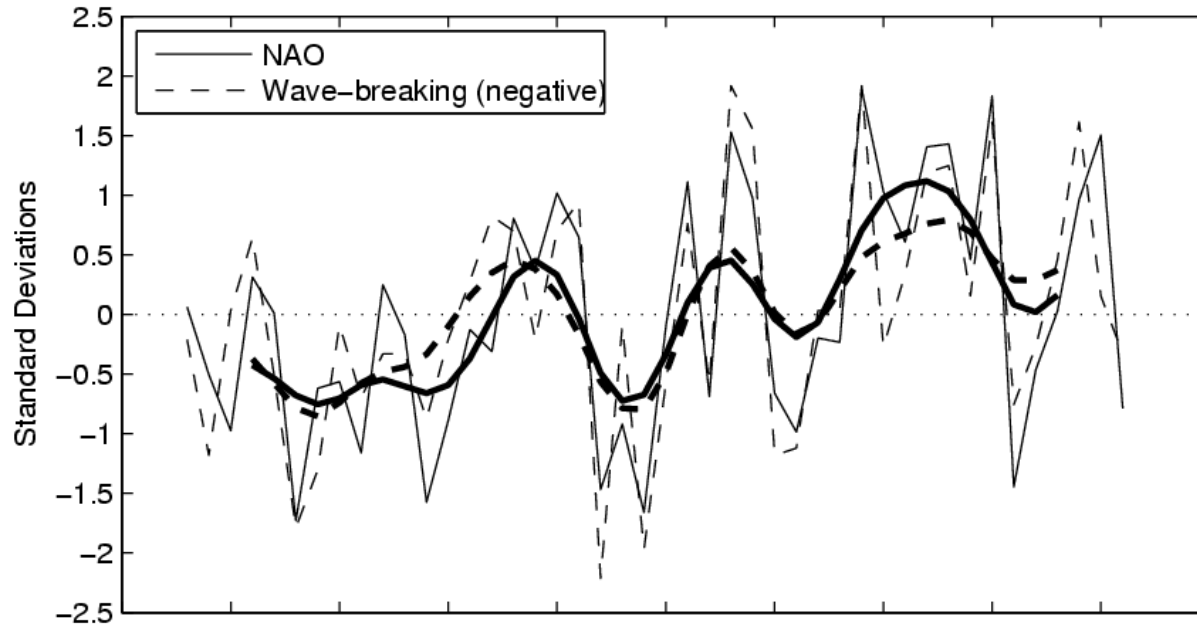


460 non-WB NAO- DJF days

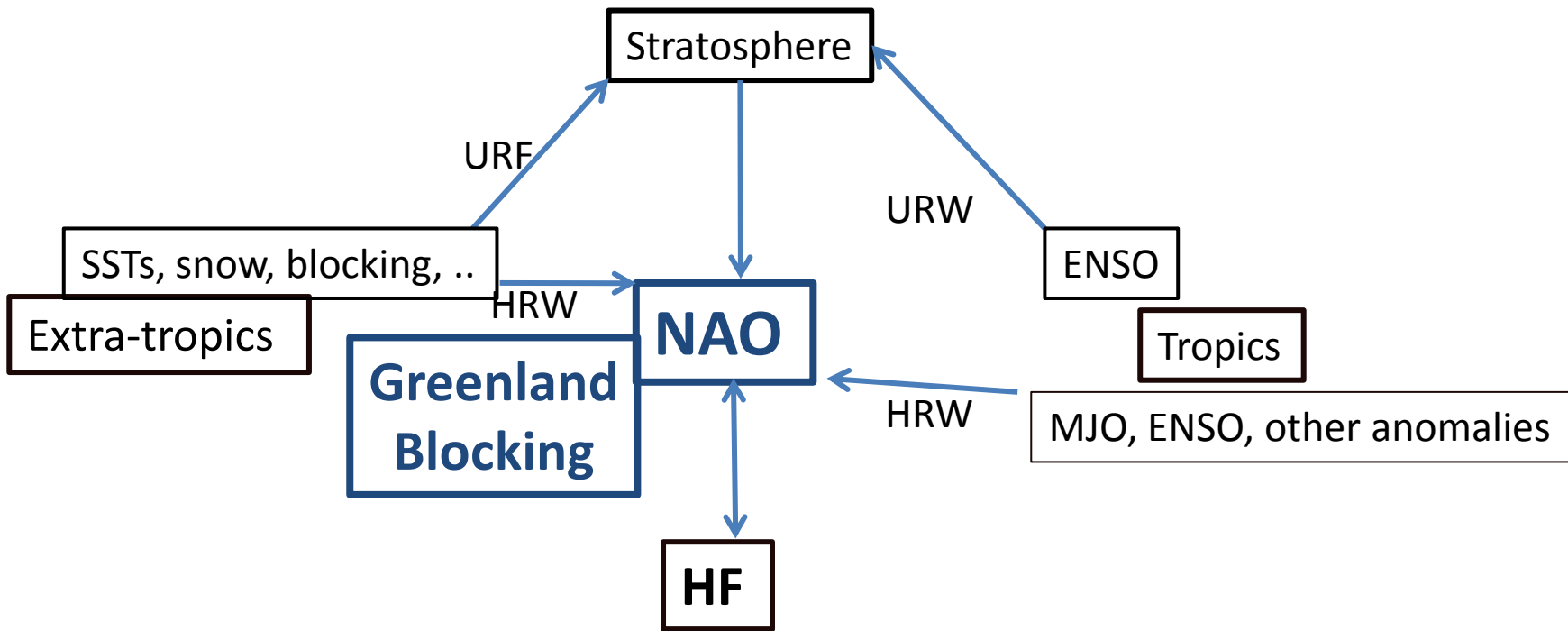
460 non-WB NAO- DJF days



Seasonal NAO and wave-breaking frequency time series

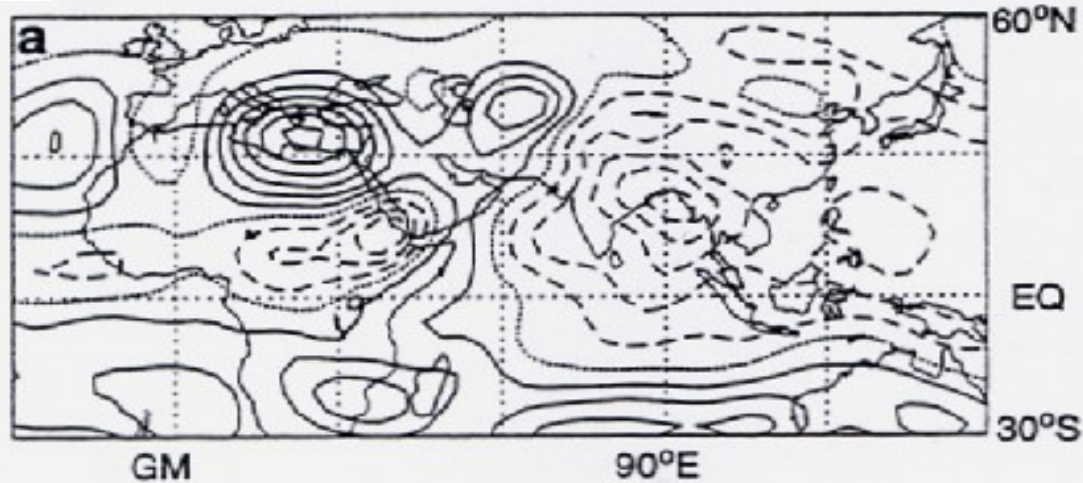
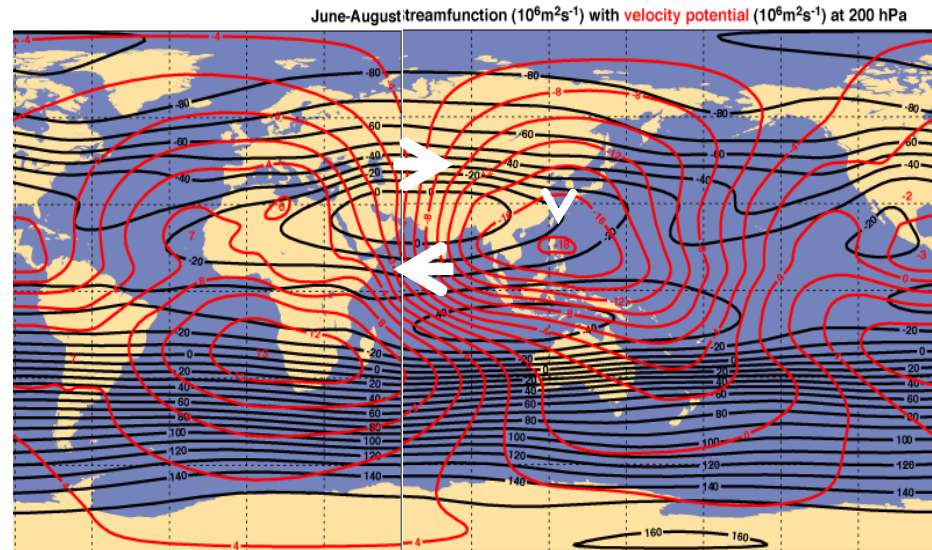


Correlations 0.84 & 0.93



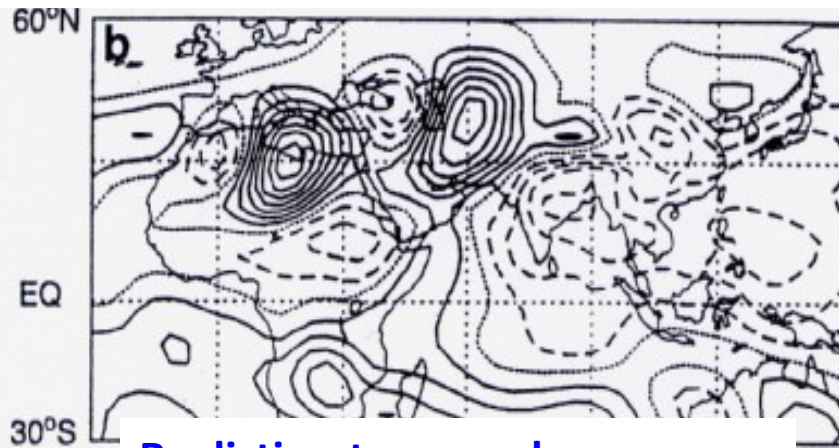
5. Events related to the Asian Summer Monsoon

Observed JJA upper ψ , χ and mid w

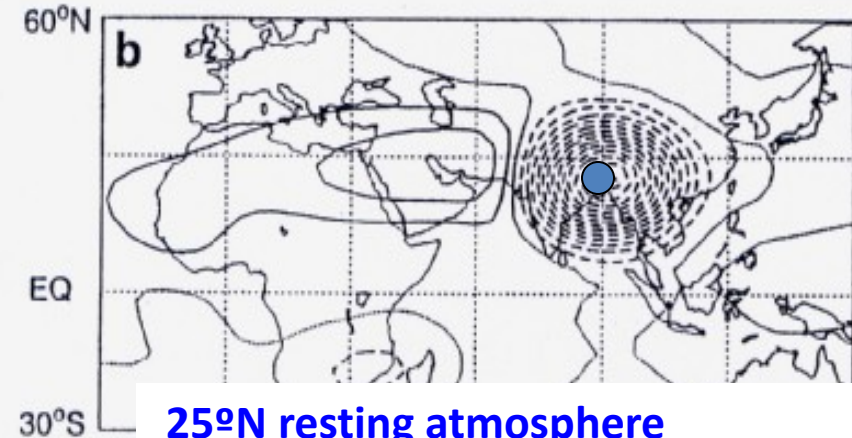


Compensating descent in a range of situations

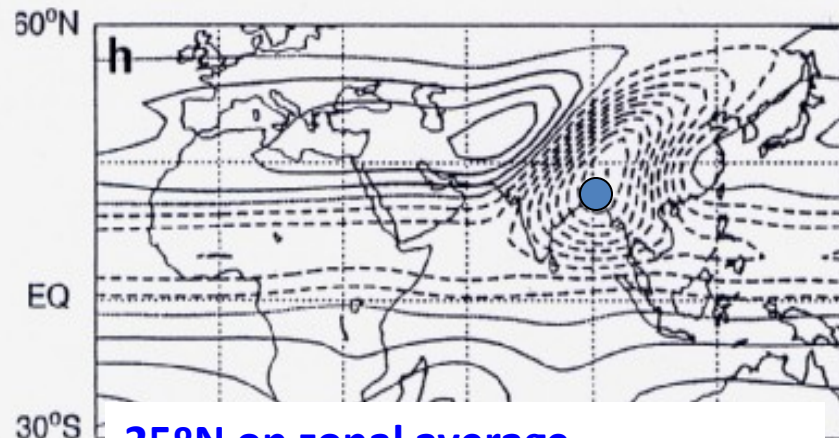
Rodwell & Hoskins (1996)



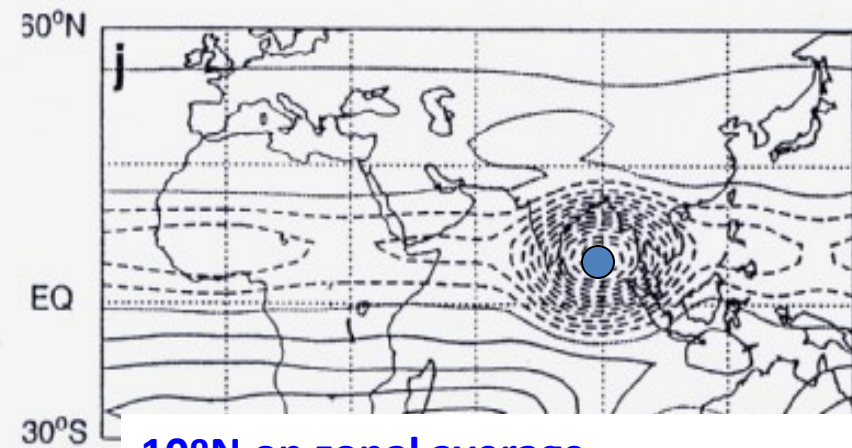
Realistic + topography



25°N resting atmosphere



25°N on zonal average



10°N on zonal average



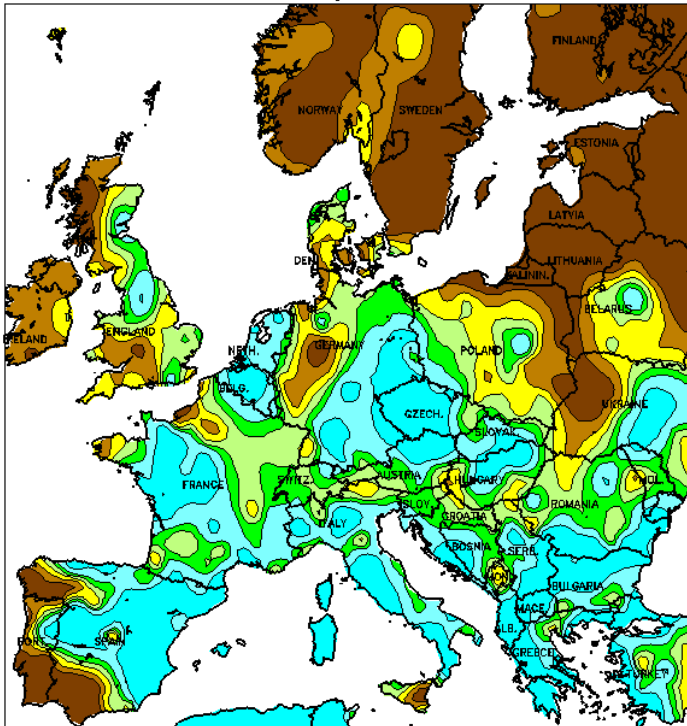
AFP

Summer 2002



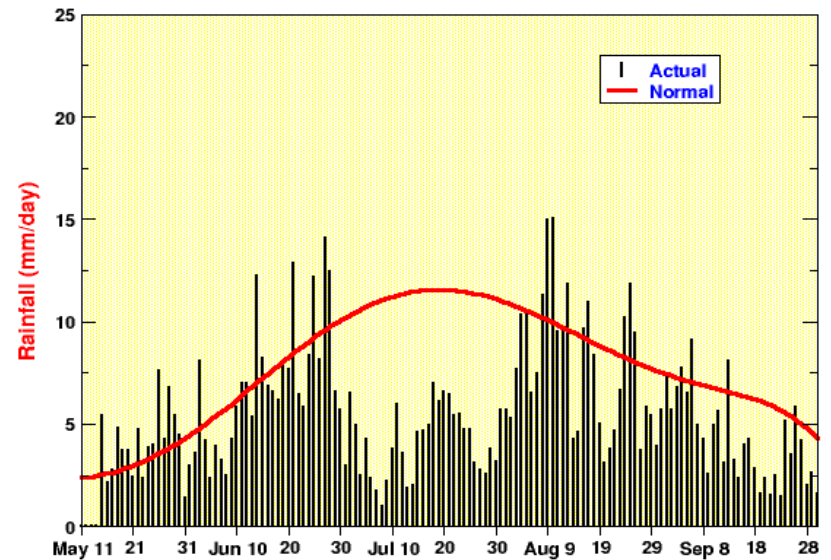
Flooding in Central Europe

EUROPE
Percent of Normal Precipitation
August 2002



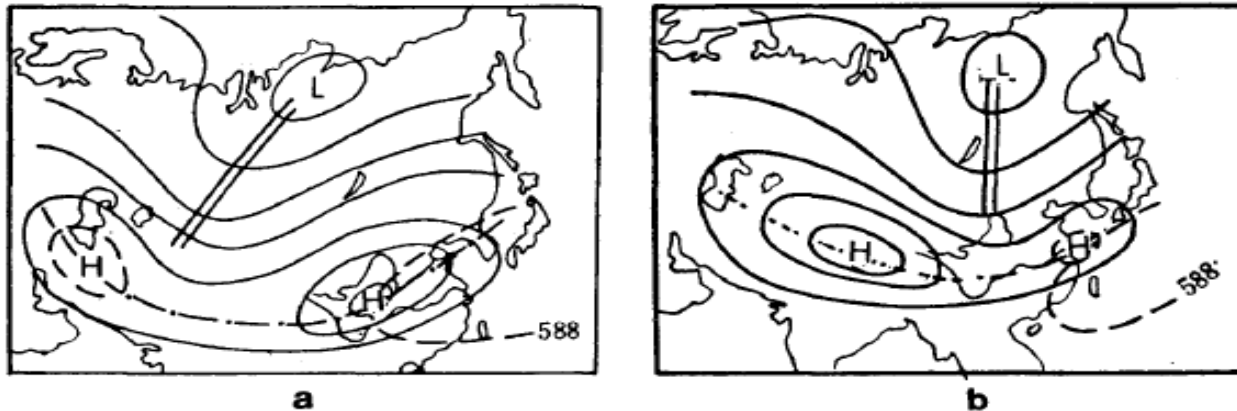
Blackburn & Hoskins (2006)

Drought in India



Summer : Quasi-biweekly oscillation of Tibetan High

Krishnamurti et al. ,1973; Krishnamurti and Bhalme, 1976; Shun, 1979; Tao and Ding, 1981

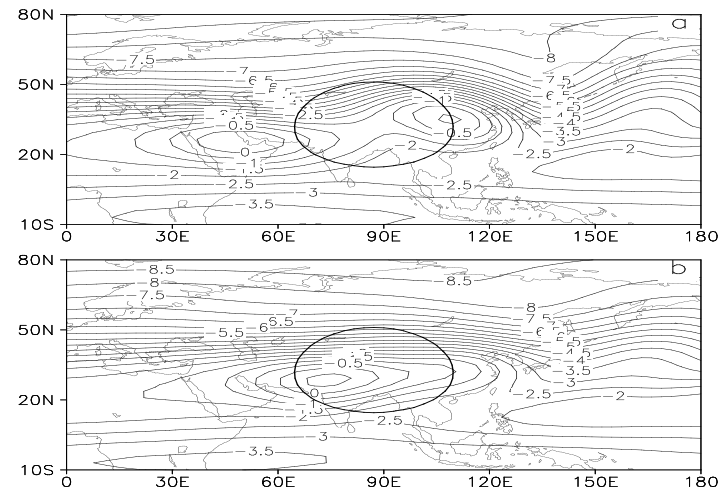


Tao and Zhu (1964)

FIG. 6. Two major patterns of the Tibetan high at (a) 200 mb; and (b) 100 mb.

Instability of low PV strip as it extends westwards from idealised Tibetan Plateau heating in a 3-D baroclinic model

Liu, Hoskins & Blackburn (2007)



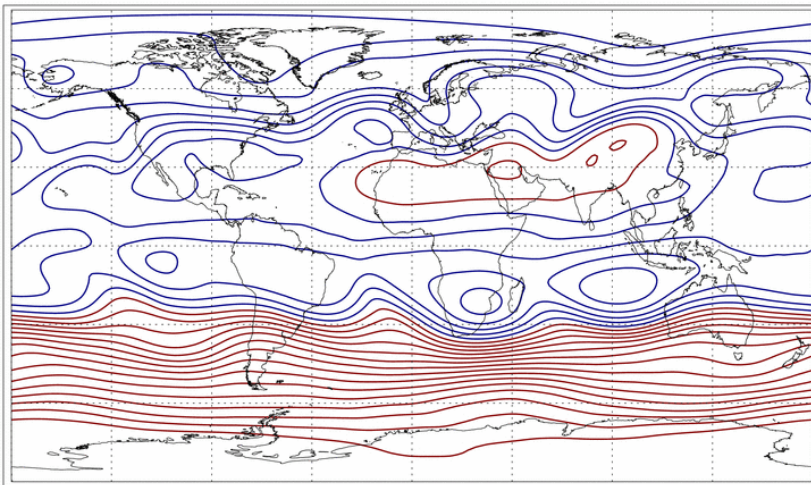
The events of July 2010

Mean fields at 250 hPa for 23-30 July

23/07/10 through
30/07/10

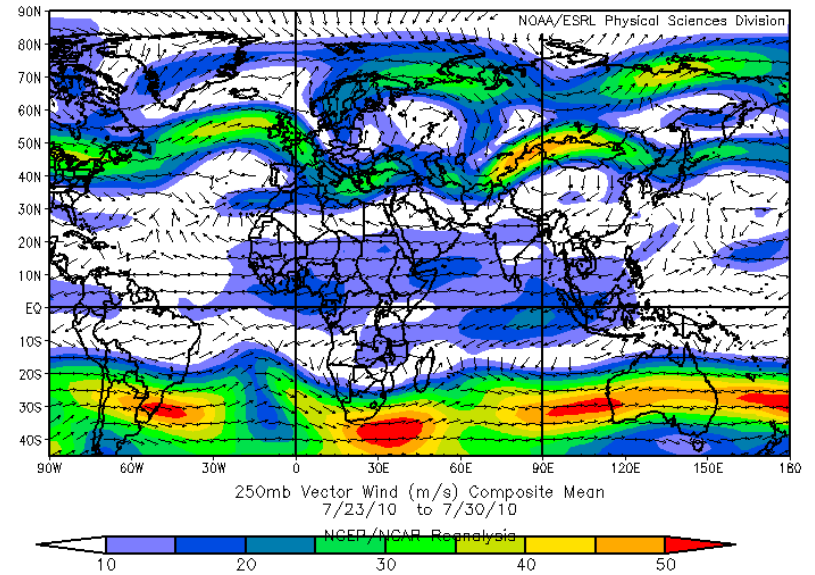
Streamfunction at 250 hPa
(contours every $10 \times 10^6 \text{ m}^2 \text{ s}^{-1}$)

Max: 157.730
Min: -84.6647



Units: $10^6 \text{ m}^2 \text{ s}^{-1}$

Stream function

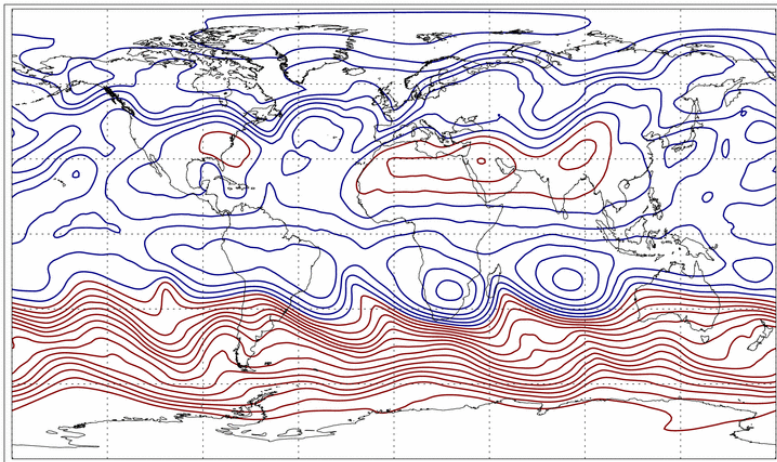


Wind

23/07/10 at 06:00 GMT
Friday

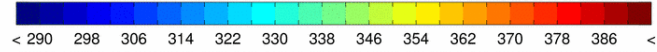
Streamfunction at 250 hPa
(contours every $10 \times 10^6 \text{ m}^2 \text{ s}^{-1}$)

Max: 166.844
Min: -83.3583



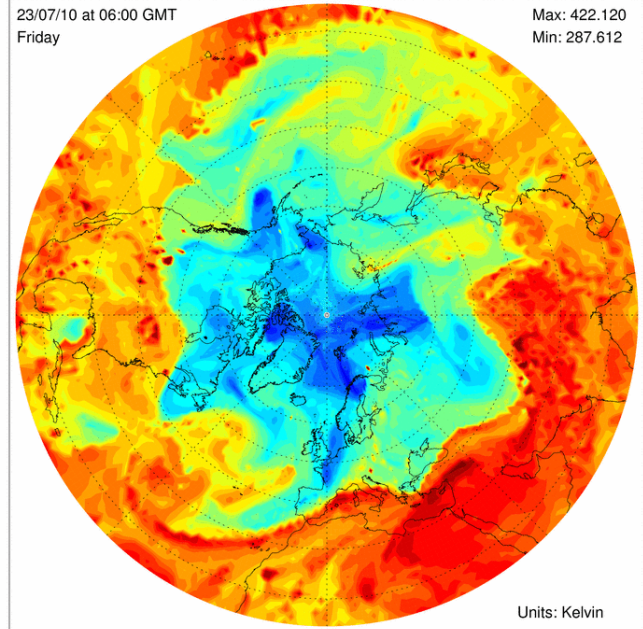
Units: $10^6 \text{ m}^2 \text{ s}^{-1}$

Theta on PV=2



23/07/10 at 06:00 GMT
Friday

Max: 422.120
Min: 287.612



Units: Kelvin

Theories of low-frequency variability over Europe

1. Synoptic eddy organisation & feedback
2. Rossby waves – horizontal propagation
 - Theory
 - MJO
 - Summer 2007
3. Teleconnection patterns
 - NAO
 - stratosphere connection
 - climate noise?
4. Blocking
 - N Hemisphere
 - European
 - Greenland
 - Summer 2003
5. Events related to the Asian Summer Monsoon
 - Summer 2002
 - bi-weekly oscillation
 - Summer 2010

The Hypothesis:

- NAO+ can be viewed as describing the basic (2-jet) state
- NAO- can be viewed as describing the state in which Rossby wave breaking is frequent (1-jet)

The West Pacific Pattern is similar but, particularly on decadal time-scales, is influenced by ENSO, PNA...

External forcings have a direct effect that makes wave-breaking more or less likely or changes its location and so produces “an NAO” response.