

Representation of air-sea interactions on an idealised coupled atmosphere-ocean model with focus on the Western Baltic Sea

3rd Workshop on Physics Dynamics Coupling – PDC18

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- ① Coastal upwelling in Western Baltic Sea
- ② Air-sea interactions: ICON & GETM
- ③ Idealised atmosphere-ocean model
- ④ Conclusions & Outlook

Coastal upwelling – coast of Poland: May 25 – Jun 08, 2008

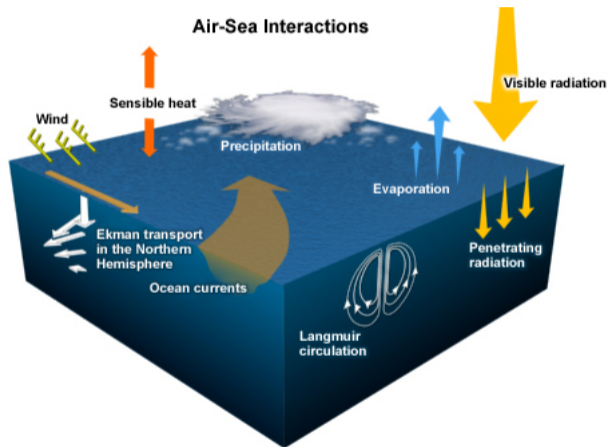


Sea surface temperature western Baltic Sea –
May/June 2008

Wind map of central Europe
at 6am UTC



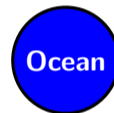
What happens at the water surface?



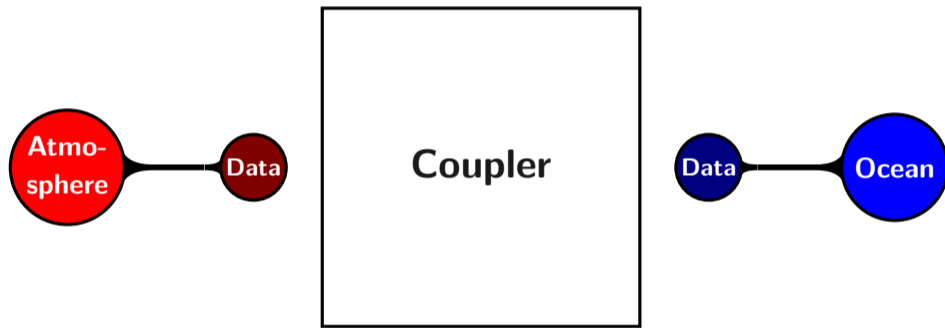
Linking of atmosphere and ocean via transfer of heat and momentum and gas exchange, i.e.

- Waves and currents in the ocean caused by wind
- Dissolution of greenhouse gases like carbon dioxide into the ocean
- Heat absorption (due to radiation) and emission by the ocean

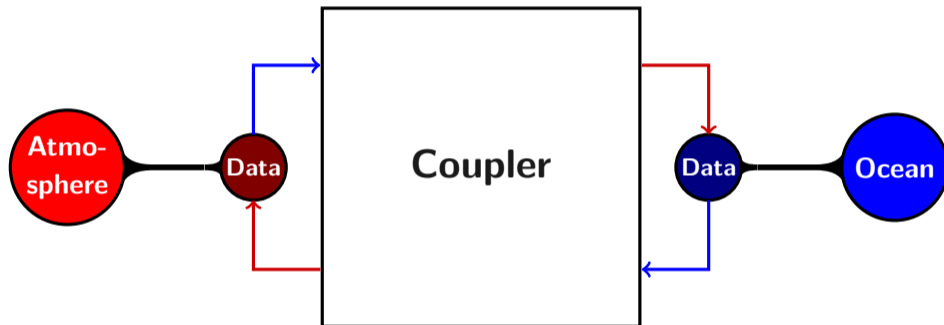
How are atmosphere and ocean models online coupled?



How are atmosphere and ocean models online coupled?

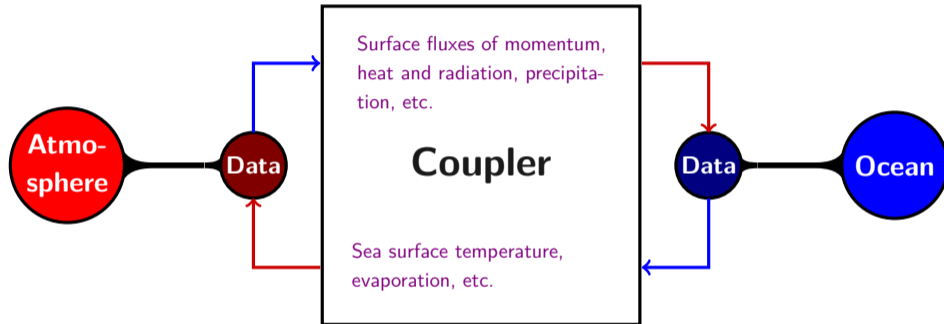


How are atmosphere and ocean models online coupled?



- Which variables will be exchanged?
- Which time intervals will be suitable for a data exchange?
- Which interpolation method will best fit for a data exchange?

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Coupling scheme for ICON and GETM

ICON

ESMF

GETM

Coupling scheme for ICON and GETM

ICON

- Local mass conservation by flux-form for continuity equation Zängl et al., 2015
- Compressible non-hydrostatic set of equations on global domains

ESMF

GETM

Coupling scheme for ICON and GETM

ICON

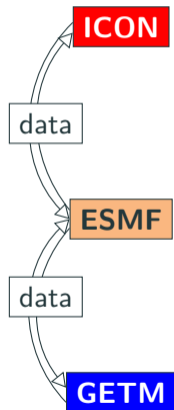
- Local mass conservation by flux-form for continuity equation Zängl et al., 2015
- Compressible non-hydrostatic set of equations on global domains

ESMF**GETM**

- Drying and flooding processes for coastal and estuarine domains
- Hydrostatic set of equations with Boussinesq approximation and eddy viscosity assumption

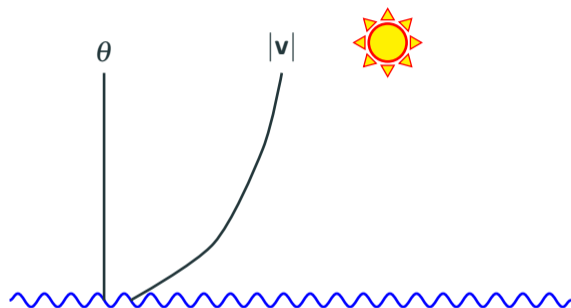
Burchard et al., 2004

Coupling scheme for ICON and GETM



- Local mass conservation by flux-form for continuity equation Zängl et al., 2015
- Compressible non-hydrostatic set of equations on global domains
- Data exchange: momentum and surface heat flux, evaporation, etc.
- Horizontal interpolation of data at air-sea interface
- Drying and flooding processes for coastal and estuarine domains
- Hydrostatic set of equations with Boussinesq approximation and eddy viscosity assumption Burchard et al., 2004

Air-sea interactions: ICON & GETM – uncoupled



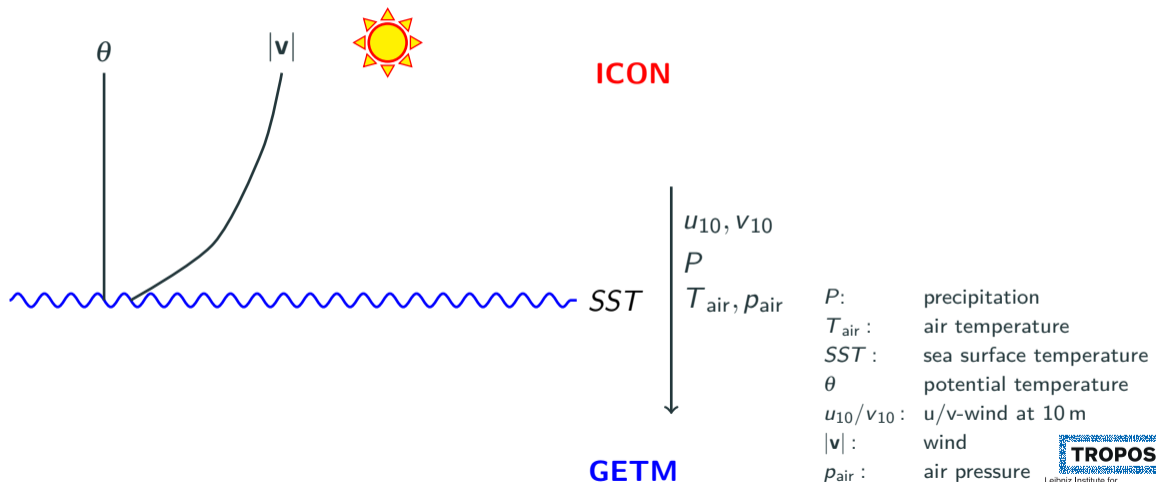
ICON

GETM

θ potential temperature

$|v|$: wind

Air-sea interactions: ICON & GETM – uncoupled



Realisation of air-sea interactions in ICON & GETM

ICON:

Momentum:

$$\tau_s^x = -\rho \cdot C_m^d \cdot |\mathbf{v}| \cdot u$$

$$\tau_s^y = -\rho \cdot C_m^d \cdot |\mathbf{v}| \cdot v$$

Heat:

$$Q = Q_s + Q_l + Q_b + Q_{SW}$$

GETM:

Momentum:

$$\tau_s^x = \rho \cdot C_m^d \cdot |\mathbf{v}| \cdot u$$

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Realisation of air-sea interactions in ICON & GETM

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Heat:

$$Q = Q_s + Q_l + Q_b + Q_{SW}$$

No mass exchange with **ocean** via precipitation and evaporation due to exact local mass conservation.

GETM:

Momentum:

$$\tau_s^x = \rho \cdot C_m^d \cdot |\mathbf{v}| \cdot u$$

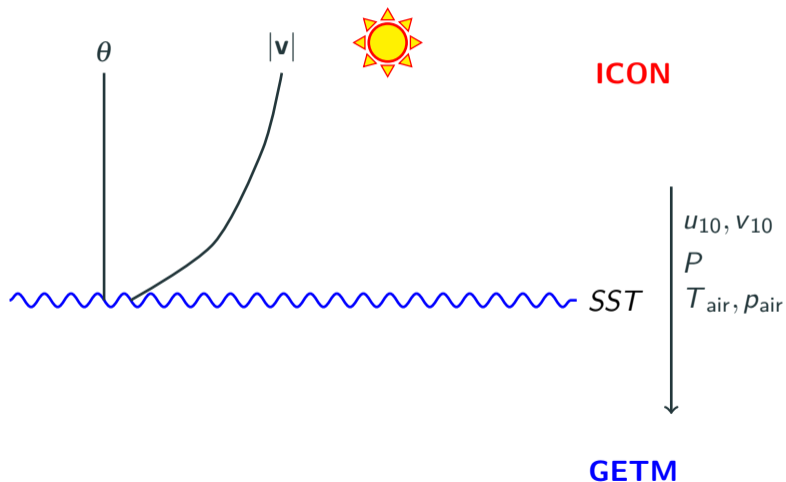
$$\tau_s^y = \rho \cdot C_m^d \cdot |\mathbf{v}| \cdot v$$

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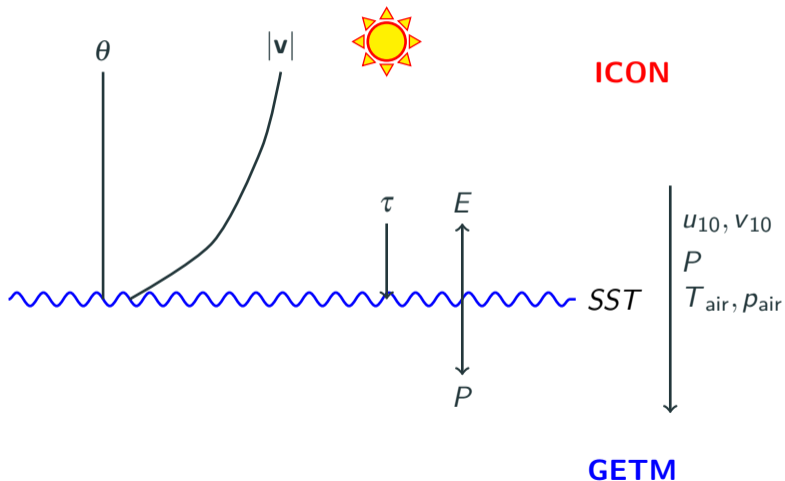
Considering of precipitation and evaporation for salinity flux.

Air-sea interactions: ICON & GETM – uncoupled



P : precipitation
 T_{air} : air temperature
 SST : sea surface temperature
 θ : potential temperature
 u_{10}/v_{10} : u/v-wind at 10 m
 $|\mathbf{v}|$: wind
 p_{air} : air pressure

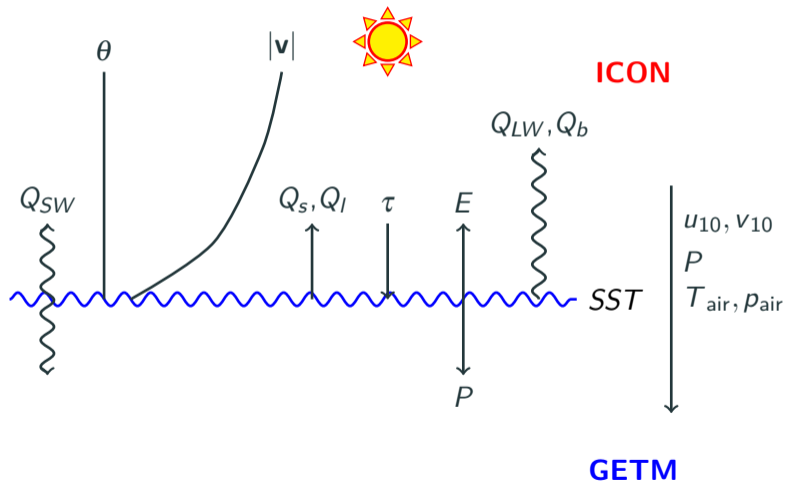
Air-sea interactions: ICON & GETM – coupled



- τ : shear stress
- E : evaporation
- P : precipitation
- T_{air} : air temperature
- SST : sea surface temperature
- θ : potential temperature
- u_{10}/v_{10} : u/v-wind at 10 m
- $|v|$: wind
- p_{air} : air pressure



Air-sea interactions: ICON & GETM – coupled



Q_s, Q_l :	sensible/latent heat flux
Q_{SW} :	solar short wave radiative flux
Q_{LW} :	terrestrial long wave radiative flux
Q_b :	long wave net radiative flux
τ :	shear stress
E :	evaporation
P :	precipitation
T_{air} :	air temperature
SST :	sea surface temperature
θ :	potential temperature
u_{10}/v_{10} :	u/v-wind at 10 m
$ v $:	wind
p_{air} :	air pressure

Idealised atmosphere-ocean model: objectives

- Development of idealised model for
 - 1D: Studying mass, momentum and energy coupling between **atmosphere** and **ocean** with a water/air column model system
 - 2D: Constructing an idealised coupled model system with straight coast and upwelling favourable winds
 - 3D: Fully coupled idealised **atmosphere-ocean** experiment (Baltic Sea)

Idealised atmosphere-ocean model: objectives

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 - 1D: Studying mass, momentum and energy coupling between **atmosphere** and **ocean** with a water/air column model system
 - 2D: Constructing an idealised coupled model system with straight coast and upwelling favourable winds
 - 3D: Fully coupled idealised **atmosphere-ocean** experiment (Baltic Sea)
- Utilising different coupling strategies
 - a) Online coupling with coupler (e.g. ESMF)
 - b) Derivation and application of numerical methods with multirate approaches for **atmosphere-ocean** models

Idealised atmosphere-ocean model: properties

- Mass and momentum conservation and energy consistency
- Unified parameterisation of air-sea interactions
- Applying parameterisation for radiative energy intake in **ocean**
- Utilising turbulence closure scheme for **atmosphere** and **ocean**
- Possible different discretisation for **atmosphere** and **ocean**, i.e. horizontal interpolation at air-sea interface as part of discretisation

Idealised atmosphere-ocean model: properties

- Mass and momentum conservation and energy consistency
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- Utilising turbulence closure scheme for atmosphere and ocean
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Idealised atmosphere-ocean model: continuity equation

- 1 **Atmosphere** components: dry air (d), water vapour (v), liquid water (l), ice (i), rain drops (r) and snow (sn)

Wacker et al., 2006, Bott, 2008

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② **Ocean** components: fresh water (f) and salinity (sa)

Burchard et al., 2004

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

d

v

l

i

r

sn



Ocean

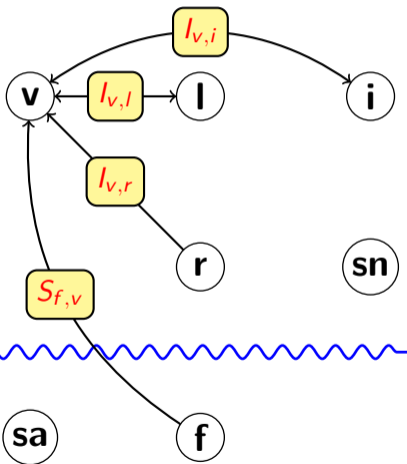
sa

f

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

(d)



Water vapour (v):

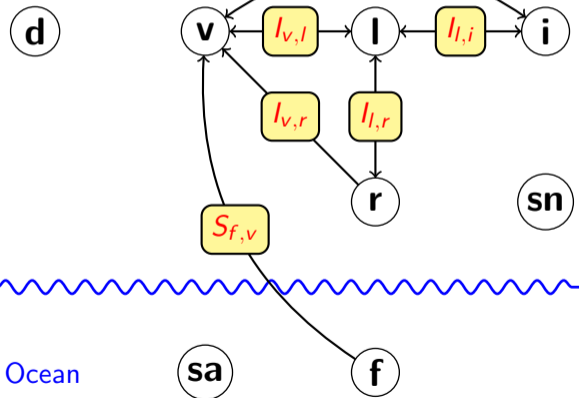
$$I_v = -I_{v,l} - I_{v,i} - I_{v,r}$$

$$S_v = S_{f,v}$$

Ocean

Idealised atmosphere-ocean model: source and sink connections

Atmosphere



Water vapour (v): $I_v = -I_{v,l} - I_{v,i} - I_{v,r}$

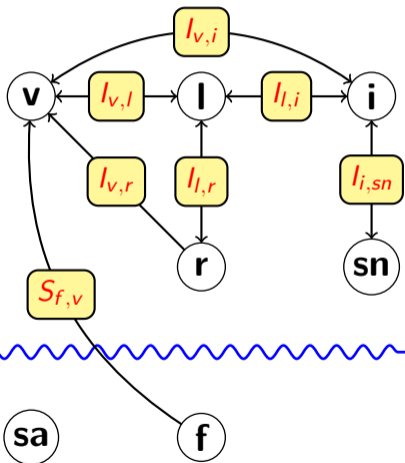
$S_v = S_{f,v}$

Liquid water (l): $I_l = I_{v,l} - I_{l,i} - I_{l,r}$

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

(d)



Water vapour (v): $I_v = -I_{v,l} - I_{v,i} - I_{v,r}$

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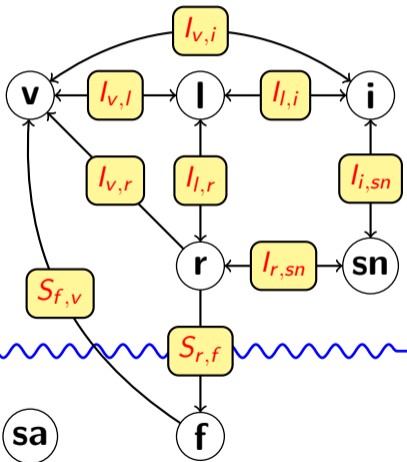
Liquid water (l): $I_l = I_{v,l} - I_{l,i} - I_{l,r}$

Ice (i): $I_i = I_{v,i} + I_{l,i} - I_{i,sn}$

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

(d)



Water vapour (v): $I_v = -I_{v,l} - I_{v,i} - I_{v,r}$

$S_v = S_{f,v}$

Liquid water (l): $I_l = I_{v,l} - I_{l,i} - I_{l,r}$

Ice (i): $I_i = I_{v,i} + I_{l,i} - I_{i,sn}$

Rain drops (r): $I_r = I_{v,r} + I_{l,r} - I_{r,sn}$

$S_r = -S_{r,f}$

Ocean

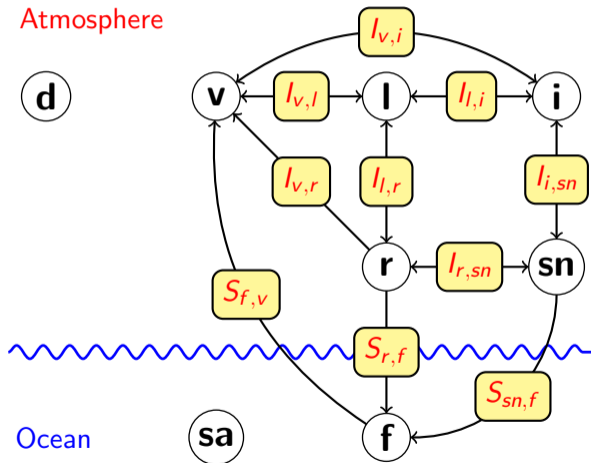
(sa)

(f)

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

(d)



Water vapour (v): $I_v = -I_{v,l} - I_{v,i} - I_{v,r}$

$$S_v = S_{f,v}$$

Liquid water (l):

$$I_l = I_{v,l} - I_{l,i} - I_{l,r}$$

Ice (i):

$$I_i = I_{v,i} + I_{l,i} - I_{i,sn}$$

Rain drops (r):

$$I_r = I_{v,r} + I_{l,r} - I_{r,sn}$$

$$S_r = -S_{r,f}$$

Snow (sn):

$$I_{sn} = I_{r,sn} + I_{i,sn}$$

$$S_{sn} = -S_{sn,f}$$

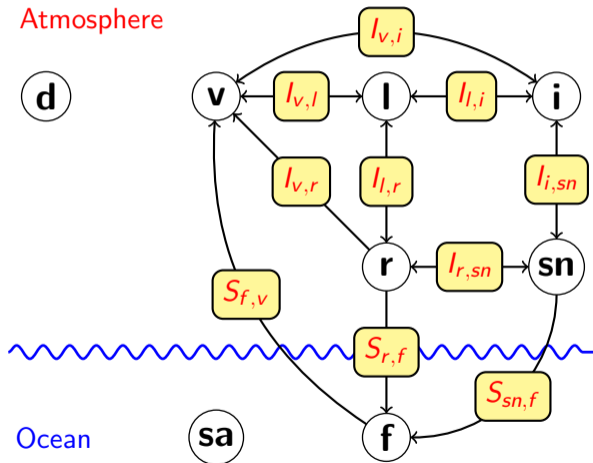
Ocean

(sa)

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

(d)



Ocean

(sa)

Water vapour (v): $I_v = -I_{v,l} - I_{v,i} - I_{v,r}$

$$S_v = S_{f,v}$$

Liquid water (l): $I_l = I_{v,l} - I_{l,i} - I_{l,r}$

Ice (i): $I_i = I_{v,i} + I_{l,i} - I_{i,sn}$

Rain drops (r): $I_r = I_{v,r} + I_{l,r} - I_{r,sn}$

$$S_r = -S_{r,f}$$

Snow (sn): $I_{sn} = I_{r,sn} + I_{i,sn}$

$$S_{sn} = -S_{sn,f}$$

Fresh water (f): $S_f = S_{r,f} + S_{sn,f} - S_{f,v}$

Idealised atmosphere-ocean model: continuity equation

- ① **Atmosphere** components: dry air (d), water vapour (v), liquid water (l), ice (i),
rain drops (r) and snow (sn) Wacker et al., 2006, Bott, 2008

- ② **Ocean** components: fresh water (f) and salinity (sa) Burchard et al., 2004



Idealised atmosphere-ocean model: continuity equation

- 1 **Atmosphere** components: dry air (d), water vapour (v), liquid water (l), ice (i),
rain drops (r) and snow (sn) Wacker et al., 2006, Bott, 2008
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- 3 No internal and external source and sink terms for dry air and salinity

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Wacker et al., 2006, Bott, 2008
- 2 **Ocean** components: fresh water (f) and salinity (sa)
Burchard et al., 2004
- 3 No internal and external source and sink terms for dry air and salinity
- 4 No internal source and sink term for fresh water

Idealised atmosphere-ocean model: continuity equation

- ① **Atmosphere** components: dry air (d), water vapour (v), liquid water (l), ice (i), rain drops (r) and snow (sn)
Wacker et al., 2006, Bott, 2008
- ② **Ocean** components: fresh water (f) and salinity (sa)
Burchard et al., 2004
- ③ No internal and external source and sink terms for dry air and salinity
- ④ No internal source and sink term for fresh water

Mass conservation of **atmosphere-ocean** system:

⇒ exchange of mass at air-sea interface

⇒ **atmosphere** and **ocean**, each on its own not mass conserving

⇒ compressible and non-hydrostatic set of equation

Idealised atmosphere-ocean model: continuity equation

Atmosphere:

- Dry air (d):

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \cdot \mathbf{v}_d) = 0$$

- All other components:

$$\frac{\partial \rho_k}{\partial t} + \nabla \cdot (\rho_k \cdot \mathbf{v}_k) = I_k + S_k$$

Idealised atmosphere-ocean model: continuity equation

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- All other components:

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$$\frac{\partial \rho^A}{\partial t} + \nabla \cdot (\rho^A \cdot \mathbf{v}^A) = \sum [I_k + S_k] = S$$

Idealised atmosphere-ocean model: continuity equation

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$$\frac{\partial \rho^A}{\partial t} + \nabla \cdot (\rho^A \cdot \mathbf{v}^A) = \sum [I_k + S_k] = S$$

Ocean:

- Fresh water (f):

$$\frac{\partial \rho_f}{\partial t} + \nabla \cdot (\rho_f \cdot \mathbf{v}_f) = S_f$$

- Salinity (sa):

$$\frac{\partial \rho_{sa}}{\partial t} + \nabla \cdot (\rho_{sa} \cdot \mathbf{v}_{sa}) = 0$$

Idealised atmosphere-ocean model: continuity equation

Atmosphere:

- Dry air (d):

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \cdot \mathbf{v}_d) = 0$$

- All other components:

$$\frac{\partial \rho_k}{\partial t} + \nabla \cdot (\rho_k \cdot \mathbf{v}_k) = I_k + S_k$$

$$\frac{\partial \rho^A}{\partial t} + \nabla \cdot (\rho^A \cdot \mathbf{v}^A) = \sum [I_k + S_k] = S$$

Ocean:

- Fresh water (f):

$$\frac{\partial \rho_f}{\partial t} + \nabla \cdot (\rho_f \cdot \mathbf{v}_f) = S_f$$

- Salinity (sa):

$$\frac{\partial \rho_{sa}}{\partial t} + \nabla \cdot (\rho_{sa} \cdot \mathbf{v}_{sa}) = 0$$

$$\frac{\partial \rho^O}{\partial t} + \nabla \cdot (\rho^O \cdot \mathbf{v}^O) = S_f$$

Idealised atmosphere-ocean model: continuity equation

Atmosphere:

- Dry air (d):

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \cdot \mathbf{v}_d) = 0$$

- All other components:

$$\frac{\partial \rho_k}{\partial t} + \nabla \cdot (\rho_k \cdot \mathbf{v}_k) = I_k + S_k$$

$$\frac{\partial \rho^A}{\partial t} + \nabla \cdot (\rho^A \cdot \mathbf{v}^A) = \sum [I_k + S_k] = S$$

Ocean:

- Fresh water (f):

$$\frac{\partial \rho_f}{\partial t} + \nabla \cdot (\rho_f \cdot \mathbf{v}_f) = S_f$$

- Salinity (sa):

$$\frac{\partial \rho_{sa}}{\partial t} + \nabla \cdot (\rho_{sa} \cdot \mathbf{v}_{sa}) = 0$$

$$\frac{\partial \rho^O}{\partial t} + \nabla \cdot (\rho^O \cdot \mathbf{v}^O) = S_f$$

Mass conserving: $\frac{\partial (\rho^A + \rho^O)}{\partial t} + \nabla \cdot (\rho^A \cdot \mathbf{v}^A + \rho^O \cdot \mathbf{v}^O) = S + S_f = 0$

$$\Rightarrow S = -S_f$$

Idealised atmosphere-ocean model: further assumptions

Atmosphere:

- Treatment as ideal gas
- No pressure forces on hydrometers, i.e. only on dry air and water vapour
- Equation of state: $p = \rho^A \cdot R \cdot T = \rho^A \cdot R_d \cdot T_v$



Ocean:

- Handling of salinity as tracer
- Linearised equation of state: $\rho^O = \rho_0^O \cdot (1 + \alpha \cdot (\theta - \theta_I) + \beta \cdot (sa - sa_I))$

Idealised atmosphere-ocean model: momentum equation of atmosphere

Momentum equation (**atmosphere**):

$$\frac{\partial(\rho^A \mathbf{v}^A)}{\partial t} + \nabla \cdot (\rho^A \mathbf{v}^A \cdot \mathbf{v}^A T) = -\nabla p^A - \rho^A \cdot \nabla \phi - 2 \cdot \Omega \times \rho^A \mathbf{v}^A + \nabla \cdot \boldsymbol{\tau}^A + \mathbf{v}^A \cdot S$$

$$+ \sum [(\mathbf{v}_k - \mathbf{v}^A) \cdot (I_k + S_k)] - \sum \left[\nabla \cdot (\rho_k (\mathbf{v}_k - \mathbf{v}^A) \cdot (\mathbf{v}_k - \mathbf{v}^A) T) \right]$$

Idealised atmosphere-ocean model: momentum equation of atmosphere

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Differences to **ICON**:

Lange, 2002, Gassmann et al., 2008

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Differences to **ICON**:

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- Mass conservation: $S = 0$

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Differences to **ICON**:

Lange, 2002, Gassmann et al., 2008

- Mass conservation: $S = 0$
- a) $\sum [(\mathbf{v}_k - \mathbf{v}^A) \cdot (I_k + S_k)] = 0$ (conservation of momentum due to chemical reactions)

Idealised atmosphere-ocean model: momentum equation of atmosphere

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Differences to **ICON**:

Lange, 2002, Gassmann et al., 2008

- Mass conservation: $S = 0$
- a) $\sum [(\mathbf{v}_k - \mathbf{v}^A) \cdot (I_k + S_k)] = 0$ (conservation of momentum due to chemical reactions)
- b) $\mathbf{v}_k \approx \mathbf{v}^A \Rightarrow \sum \left[\nabla \cdot (\rho_k (\mathbf{v}_k - \mathbf{v}^A) \cdot (\mathbf{v}_k - \mathbf{v}^A) T) \right] \ll \nabla \cdot (\rho^A \mathbf{v}^A \cdot \mathbf{v}^A T) \Rightarrow$ negligible

Idealised atmosphere-ocean model: momentum equation of atmosphere

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$$\frac{\partial(\rho^A \mathbf{v}^A)}{\partial t} + \nabla \cdot (\rho^A \mathbf{v}^A \cdot \mathbf{v}^A T) = -\nabla p^A - \rho^A \cdot \nabla \phi - 2 \cdot \Omega \times \rho^A \mathbf{v}^A + \nabla \cdot \tau^A$$

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Lange, 2002, Gassmann et al., 2008

- Mass conservation: $S = 0$
- a) $\sum [(\mathbf{v}_k - \mathbf{v}^A) \cdot (I_k + S_k)] = 0$ (conservation of momentum due to chemical reactions)
- b) $\mathbf{v}_k \approx \mathbf{v}^A \Rightarrow \sum [\nabla \cdot (\rho_k (\mathbf{v}_k - \mathbf{v}^A) \cdot (\mathbf{v}_k - \mathbf{v}^A) T)] \ll \nabla \cdot (\rho^A \mathbf{v}^A \cdot \mathbf{v}^A T) \Rightarrow$ negligible

Idealised atmosphere-ocean model: momentum equation of ocean

Momentum equation ([ocean](#)):

$$\frac{\partial(\rho^O \mathbf{v}^O)}{\partial t} + \nabla \cdot (\rho^O \mathbf{v}^O \cdot \mathbf{v}^{OT}) = -\nabla p^O - \rho^O \cdot \nabla \phi - 2 \cdot \Omega \times \rho^O \mathbf{v}^O + \nabla \cdot \boldsymbol{\tau}^O + \mathbf{v}_f \cdot S_f - \sum \left[\nabla \cdot (\rho_k (\mathbf{v}_k - \mathbf{v}^O) \cdot (\mathbf{v}_k - \mathbf{v}^O)^T) \right]$$

Idealised atmosphere-ocean model: momentum equation of ocean

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Differences to [GETM](#):

Burchard et al., 2004

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Differences to [GETM](#):

Burchard et al., 2004

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Idealised atmosphere-ocean model: momentum equation of ocean

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Differences to [GETM](#):

Burchard et al., 2004

- Boussinesq approximation leads to mass conservation, i.e. $S_f = 0$
- $\mathbf{v}_k = \mathbf{v}^O \Rightarrow \sum \left[\nabla \cdot (\rho_k (\mathbf{v}_k - \mathbf{v}) \cdot (\mathbf{v}_k - \mathbf{v})^T) \right] = 0$

Idealised atmosphere-ocean model: momentum equation of ocean

Momentum equation (ocean):

$$\frac{\partial(\rho^O \mathbf{v}^O)}{\partial t} + \nabla \cdot (\rho^O \mathbf{v}^O \cdot \mathbf{v}^{OT}) = -\nabla p^O - \rho^O \cdot \nabla \phi - 2 \cdot \Omega \times \rho^O \mathbf{v}^O + \nabla \cdot \boldsymbol{\tau}^O$$

Differences to GETM:

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- Boussinesq approximation leads to mass conservation, i.e. $S_f = 0$
- $\mathbf{v}_k = \mathbf{v}^O \Rightarrow \sum \left[\nabla \cdot (\rho_k (\mathbf{v}_k - \mathbf{v}) \cdot (\mathbf{v}_k - \mathbf{v})^T) \right] = 0$

Idealised atmosphere-ocean model: energy equation of atmosphere

Energy equation (**atmosphere**):

$$\begin{aligned}
 & \sum \left[\frac{\partial(\rho_k K_k)}{\partial t} + \nabla \cdot (\rho_k K_k \cdot \mathbf{v}_k) \right] + \frac{\partial(\rho^A \phi)}{\partial t} + \nabla \cdot (\rho^A \phi \cdot \mathbf{v}^A) + \frac{\partial(\rho^A e^A)}{\partial t} + \nabla \cdot (\rho^A e^A \cdot \mathbf{v}^A) \\
 = & \sum \left[-(\mathbf{v}_k - \mathbf{v}^A) \cdot \nabla \rho_k + (\mathbf{v}_k - \mathbf{v}^A) \cdot (\nabla \cdot \boldsymbol{\tau}_k) + (K_k - K^A) \cdot (I_k + S_k) \right] \\
 & - \nabla \cdot (\rho^A \cdot \mathbf{v}^A) + \nabla \cdot (\boldsymbol{\tau}^A \cdot \mathbf{v}^A) - \nabla \cdot Q^A + (K^A + \phi + h^A) \cdot S
 \end{aligned}$$

Idealised atmosphere-ocean model: energy equation of atmosphere

Energy equation (**atmosphere**):

$$\begin{aligned} & \sum \left[\frac{\partial(\rho_k K_k)}{\partial t} + \nabla \cdot (\rho_k K_k \cdot \mathbf{v}_k) \right] + \frac{\partial(\rho^A \phi)}{\partial t} + \nabla \cdot (\rho^A \phi \cdot \mathbf{v}^A) + \frac{\partial(\rho^A e^A)}{\partial t} + \nabla \cdot (\rho^A e^A \cdot \mathbf{v}^A) \\ & = \sum \left[-(\mathbf{v}_k - \mathbf{v}^A) \cdot \nabla p_k + (\mathbf{v}_k - \mathbf{v}^A) \cdot (\nabla \cdot \boldsymbol{\tau}_k) + (K_k - K^A) \cdot (I_k + S_k) \right] \\ & \quad - \nabla \cdot (\rho^A \cdot \mathbf{v}^A) + \nabla \cdot (\boldsymbol{\tau}^A \cdot \mathbf{v}^A) - \nabla \cdot Q^A + (K^A + \phi + h^A) \cdot S \end{aligned}$$

Differences to **ICON**:

Lange, 2002, Gassmann et al., 2008

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Differences to **ICON**:

Lange, 2002, Gassmann et al., 2008

- Mass conservation: $S = 0$

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Differences to **ICON**:

Lange, 2002, Gassmann et al., 2008

- Mass conservation: $S = 0$
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Energy equation (**atmosphere**):

$$\begin{aligned} & \sum \left[\frac{\partial(\rho_k K_k)}{\partial t} + \nabla \cdot (\rho_k K_k \cdot \mathbf{v}_k) \right] + \frac{\partial(\rho^A \phi)}{\partial t} + \nabla \cdot (\rho^A \phi \cdot \mathbf{v}^A) + \frac{\partial(\rho^A e^A)}{\partial t} + \nabla \cdot (\rho^A e^A \cdot \mathbf{v}^A) \\ & = \sum \left[-(\mathbf{v}_k - \mathbf{v}^A) \cdot \nabla p_k + (\mathbf{v}_k - \mathbf{v}^A) \cdot (\nabla \cdot \boldsymbol{\tau}_k) + (K_k - K^A) \cdot (I_k + S_k) \right] \\ & \quad - \nabla \cdot (\rho^A \cdot \mathbf{v}^A) + \nabla \cdot (\boldsymbol{\tau}^A \cdot \mathbf{v}^A) - \nabla \cdot Q^A \end{aligned}$$

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- $\mathbf{v}_k \approx \mathbf{v}^A \Rightarrow K_k \approx K^A \Rightarrow$ negligible

Idealised atmosphere-ocean model: energy equation of atmosphere

Energy equation (**atmosphere**):

$$\frac{\partial (\rho^A (K^A + \phi + e^A))}{\partial t} + \nabla \cdot (\rho^A (K^A + \phi + e^A) \cdot \mathbf{v}^A) = -\nabla \cdot (p^A \cdot \mathbf{v}^A) + \nabla \cdot (\boldsymbol{\tau}^A \cdot \mathbf{v}^A) - \nabla \cdot Q^A$$

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Idealised atmosphere-ocean model: energy equation of ocean

Energy equation (ocean):

$$\begin{aligned}
 & \sum \left[\frac{\partial(\rho_k K_k)}{\partial t} + \nabla \cdot (\rho_k K_k \cdot \mathbf{v}_k) \right] + \frac{\partial(\rho^O \phi)}{\partial t} + \nabla \cdot (\rho^O \phi \cdot \mathbf{v}^O) + \frac{\partial(\rho^O e^O)}{\partial t} + \nabla \cdot (\rho^O e^O \cdot \mathbf{v}^O) \\
 = & \sum \left[-(\mathbf{v}_k - \mathbf{v}^O) \cdot \nabla p_k + (\mathbf{v}_k - \mathbf{v}^O) \cdot (\nabla \cdot \boldsymbol{\tau}_k) \right] \\
 & - \nabla \cdot (\rho^O \cdot \mathbf{v}^O) + \nabla \cdot (\boldsymbol{\tau}^O \cdot \mathbf{v}^O) - \nabla \cdot Q^O + (K_f + \phi + h^O) \cdot S_f
 \end{aligned}$$

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 & = \sum \left[-(\mathbf{v}_k - \mathbf{v}^O) \cdot \nabla p_k + (\mathbf{v}_k - \mathbf{v}^O) \cdot (\nabla \cdot \boldsymbol{\tau}_k) \right] \\
 & \quad - \nabla \cdot (\rho^O \cdot \mathbf{v}^O) + \nabla \cdot (\boldsymbol{\tau}^O \cdot \mathbf{v}^O) - \nabla \cdot Q^O + (K_f + \phi + h^O) \cdot S_f
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Differences to GETM:

Burchard et al., 2004

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Differences to GETM:

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Energy equation (ocean):

$$\begin{aligned} & \sum \left[\frac{\partial(\rho_k K^O)}{\partial t} + \nabla \cdot (\rho_k K^O \cdot \mathbf{v}_k) \right] + \frac{\partial(\rho^O \phi)}{\partial t} + \nabla \cdot (\rho^O \phi \cdot \mathbf{v}^O) + \frac{\partial(\rho^O e^O)}{\partial t} + \nabla \cdot (\rho^O e^O \cdot \mathbf{v}^O) \\ & = -\nabla \cdot (\rho^O \cdot \mathbf{v}^O) + \nabla \cdot (\boldsymbol{\tau}^O \cdot \mathbf{v}^O) - \nabla \cdot Q^O \end{aligned}$$

Differences to [GETM](#):

Burchard et al., 2004

- Boussinesq approximation leads to mass conservation, i.e. $S_f = 0$
- $\mathbf{v}_k = \mathbf{v}^O \Rightarrow K_k = K^O$

Idealised atmosphere-ocean model: energy equation of ocean

Energy equation (ocean):

$$\frac{\partial (\rho^O (K^O + \phi + e^O))}{\partial t} + \nabla \cdot (\rho^O (K^O + \phi + e^O) \cdot \mathbf{v}^O) = -\nabla \cdot (\rho^O \cdot \mathbf{v}^O) + \nabla \cdot (\boldsymbol{\tau}^O \cdot \mathbf{v}^O) - \nabla \cdot Q^O$$

Differences to GETM:

Burchard et al., 2004

- Boussinesq approximation leads to mass conservation, i.e. $S_f = 0$
- $\mathbf{v}_k = \mathbf{v}^O \Rightarrow K_k = K^O$

Idealised atmosphere-ocean model: air-sea interactions

- Mass and momentum conservation and energy consistency
- Unified parameterisation of air-sea interactions
- Applying parameterisation for radiative energy intake in ocean
- Utilising turbulence closure scheme for atmosphere and ocean
- Possible different discretisation for atmosphere and ocean, i.e. horizontal interpolation at air-sea interface as part of discretisation

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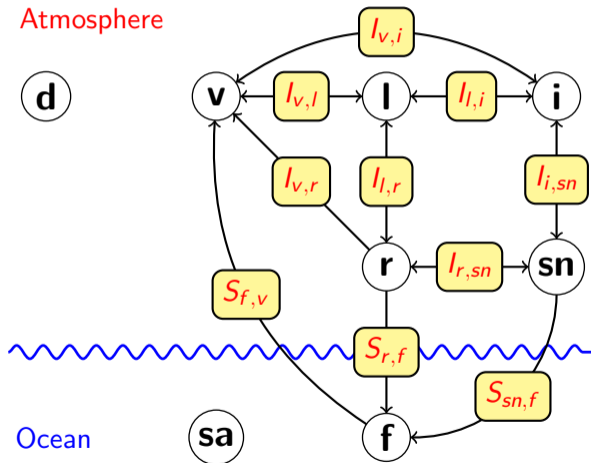
Idealised atmosphere-ocean model: air-sea interactions

- Mass exchange due to
 - a) Precipitation: $S_{r,f} + S_{sn,f}$
 - b) Evaporation: $S_{f,v}$

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

(d)



Ocean

(sa)

Water vapour (v): $I_v = -I_{v,l} - I_{v,i} - I_{v,r}$

$$S_v = S_{f,v}$$

Liquid water (l): $I_l = I_{v,l} - I_{l,i} - I_{l,r}$

Ice (i): $I_i = I_{v,i} + I_{l,i} - I_{i,sn}$

Rain drops (r): $I_r = I_{v,r} + I_{l,r} - I_{r,sn}$

$$S_r = -S_{r,f}$$

Snow (sn): $I_{sn} = I_{r,sn} + I_{i,sn}$

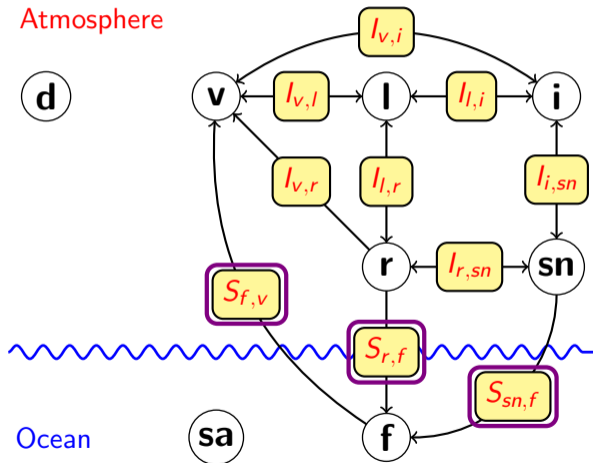
$$S_{sn} = -S_{sn,f}$$

Fresh water (f): $S_f = S_{r,f} + S_{sn,f} - S_{f,v}$

Idealised atmosphere-ocean model: source and sink connections

Atmosphere

(d)



Water vapour (v): $I_v = -I_{v,l} - I_{v,i} - I_{v,r}$

$$S_v = S_{f,v}$$

Liquid water (l):

$$I_l = I_{v,l} - I_{l,i} - I_{l,r}$$

Ice (i):

$$I_i = I_{v,i} + I_{l,i} - I_{i,sn}$$

Rain drops (r):

$$I_r = I_{v,r} + I_{l,r} - I_{r,sn}$$

$$S_r = -S_{r,f}$$

Snow (sn):

$$I_{sn} = I_{r,sn} + I_{i,sn}$$

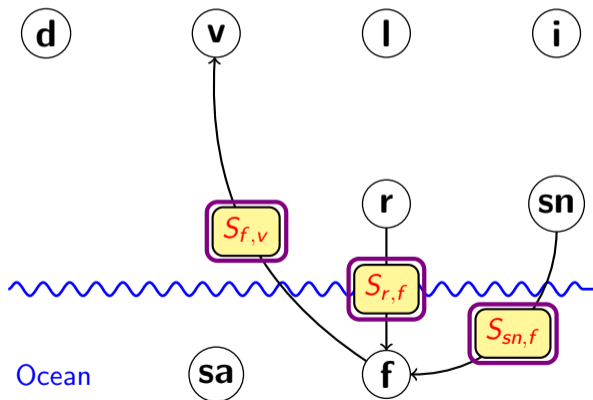
$$S_{sn} = -S_{sn,f}$$

Fresh water (f):

$$S_f = S_{r,f} + S_{sn,f} - S_{f,v}$$

Idealised atmosphere-ocean model: source and sink connections

Atmosphere



Water vapour (v):

$$S_v = S_{f,v}$$

Liquid water (l):

Ice (i):

Rain drops (r):

$$S_r = -S_{r,f}$$

Snow (sn):

$$S_{sn} = -S_{sn,f}$$

Fresh water (f):

$$S_f = S_{r,f} + S_{sn,f} - S_{f,v}$$

Idealised atmosphere-ocean model: air-sea interactions

- Mass exchange due to
 - a) Precipitation: $S_{r,f} + S_{sn,f}$
 - b) Evaporation: $S_{f,v}$

Idealised atmosphere-ocean model: air-sea interactions

- Mass exchange due to
 - a) Precipitation: $S_{r,f} + S_{sn,f}$
 - b) Evaporation: $S_{f,v}$

Note:

Mass conservation is assumed, i.e. precipitation leaves the **atmosphere** and enters the **ocean**, for evaporation vice versa.

Idealised atmosphere-ocean model: air-sea interactions

- Mass exchange due to
 - a) Precipitation: $S_{r,f} + S_{sn,f}$
 - b) Evaporation: $S_{f,v}$

Note:

Mass conservation is assumed, i.e. precipitation leaves the **atmosphere** and enters the **ocean**, for evaporation vice versa.

- Heat exchange and radiative energy intake: formulation of $\nabla \cdot Q^A$ and $\nabla \cdot Q^O$

Idealised atmosphere-ocean model: air-sea interactions

- Mass exchange due to
 - a) Precipitation: $S_{r,f} + S_{sn,f}$
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Note:

Mass conservation is assumed, i.e. precipitation leaves the **atmosphere** and enters the **ocean**, for evaporation vice versa.

- Heat exchange and radiative energy intake: formulation of $\nabla \cdot Q^A$ and $\nabla \cdot Q^O$
 - ① Treatment as external forcing of internal energy for individual **atmosphere** and **ocean** models:

$$Q^A = Q_s + Q_l + Q_b^A + Q_{LW}^A + Q_{SW}^A \quad \text{and} \quad Q^O = -Q_s - Q_l + Q_b^O + Q_{LW}^O + Q_{SW}^O$$

Idealised atmosphere-ocean model: air-sea interactions

- Mass exchange due to

a) Precipitation: $S_{r,f} + S_{sn,f}$

b) Evaporation: $S_{f,v}$

Note:

Mass conservation is assumed, i.e. precipitation leaves the **atmosphere** and enters the **ocean**, for evaporation vice versa.

- Heat exchange and radiative energy intake: formulation of $\nabla \cdot Q^A$ and $\nabla \cdot Q^O$

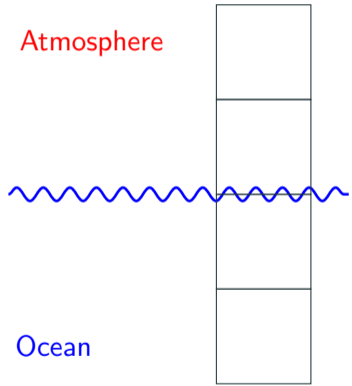
- Treatment as external forcing of internal energy for individual **atmosphere** and **ocean** models:

$$Q^A = Q_s + Q_l + Q_b^A + Q_{LW}^A + Q_{SW}^A \quad \text{and} \quad Q^O = -Q_s - Q_l + Q_b^O + Q_{LW}^O + Q_{SW}^O$$

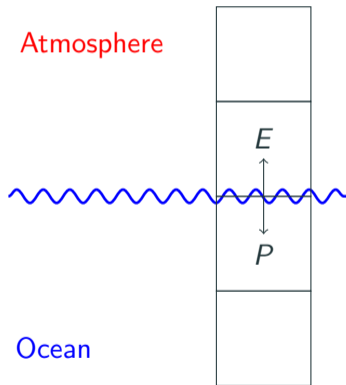
- Atmosphere-ocean** model: *radiative energy intake as external forcing of internal energy:*

$$Q = Q_b^A + Q_b^O + Q_{LW}^A + Q_{LW}^O + Q_{SW}^O + Q_{SW}^O$$

Idealised atmosphere-ocean model: vertical discretisation

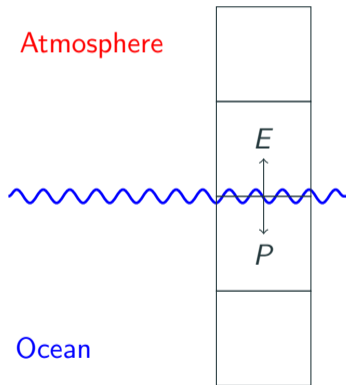


Idealised atmosphere-ocean model: vertical discretisation



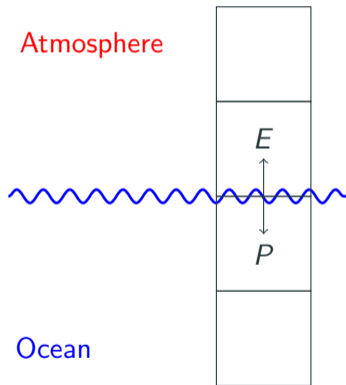
- Rise and sink of sea level with precipitation (P) and evaporation (E)

Idealised atmosphere-ocean model: vertical discretisation



- Rise and sink of sea level with precipitation (P) and evaporation (E)
- Fixed vertical layer at $z = 0$ either in **atmosphere** or **ocean**

Idealised atmosphere-ocean model: vertical discretisation



- Rise and sink of sea level with precipitation (P) and evaporation (E)
- Fixed vertical layer at $z = 0$ either in **atmosphere** or **ocean**
- Adaptive vertical discretisation necessary

Conclusions

- Coupling of **atmosphere-ocean** systems only recommended with unified parameterisation of air-sea interactions
- Mass conservation only for **atmosphere-ocean** systems and **not** for individual subsystems
- Idealised **atmosphere-ocean** model with further assumptions reformable to coupled **ICON-GETM** model
- Heat fluxes as external source for internal energy in **atmosphere** and **ocean** models, but not for whole **atmosphere-ocean** models
- Radiative energy intake always as external source for internal energy

Outlook

- Applying turbulence closure scheme for idealised model
- Formulation of heat fluxes for idealised model with use of a coupler
- Investigation of different discretisation approaches for needs of idealised model
- Validation of idealised model against benchmark tests for **atmosphere** and **ocean** parts

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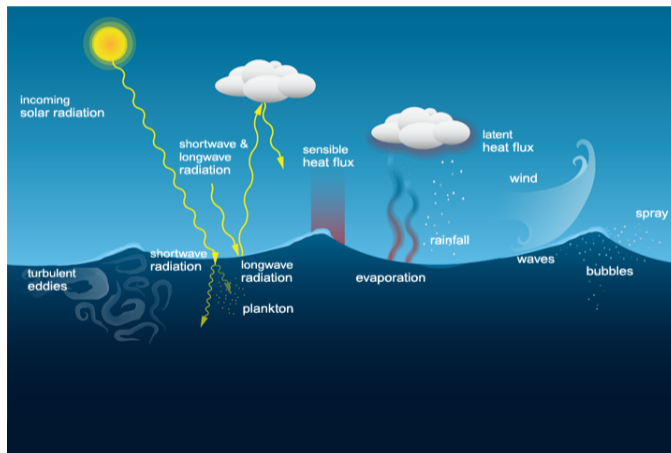
References II

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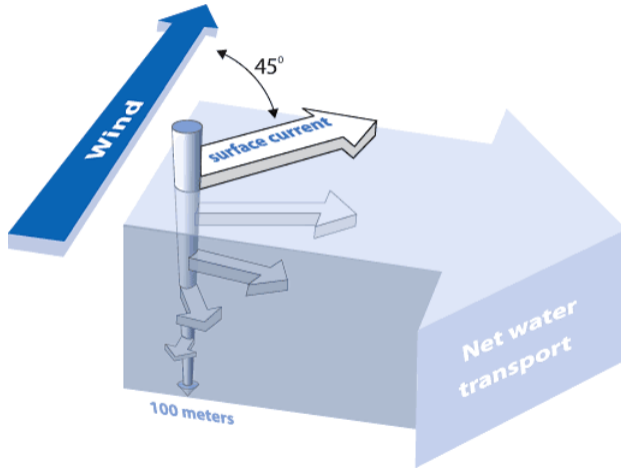
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Air-sea interactions



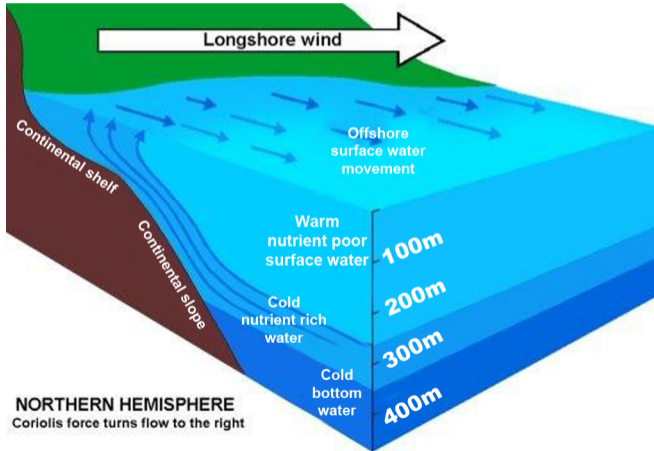
Ekman transport in water



- Rotation of 45° of surface current due to Coriolis force (Coriolis effect)
- Continuing of rotation into **ocean** till wind loses influence (Ekman spiral)
- Transporting of water in 90° angle of the wind (Ekman transport)
- Northern/southern hemisphere in right/left direction

www.oceanservice.noaa.gov (21.09.2016)

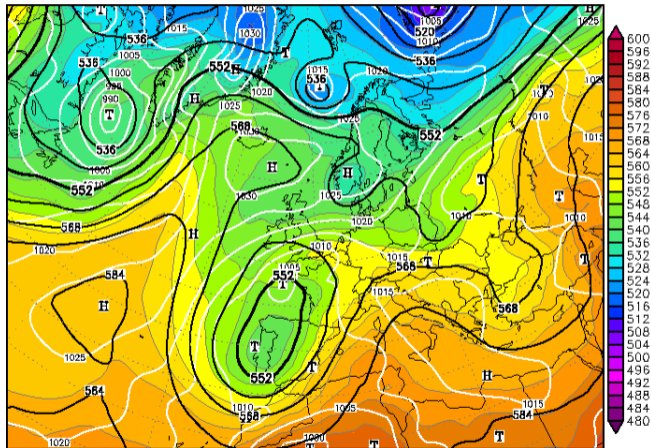
What is coastal upwelling?



- Oceanographic phenomenon
- Main drivers: wind, Coriolis effect and Ekman transport
- Brings dense, cooler and usually nutrient-rich water towards the **ocean** surface
- Higher marine productivity due to an increase in plankton
- Cooling of lower **atmosphere**

www.seos-project.eu (15.07.2016)

Coastal upwelling – coast of Poland: May 25 – Jun 08, 2008



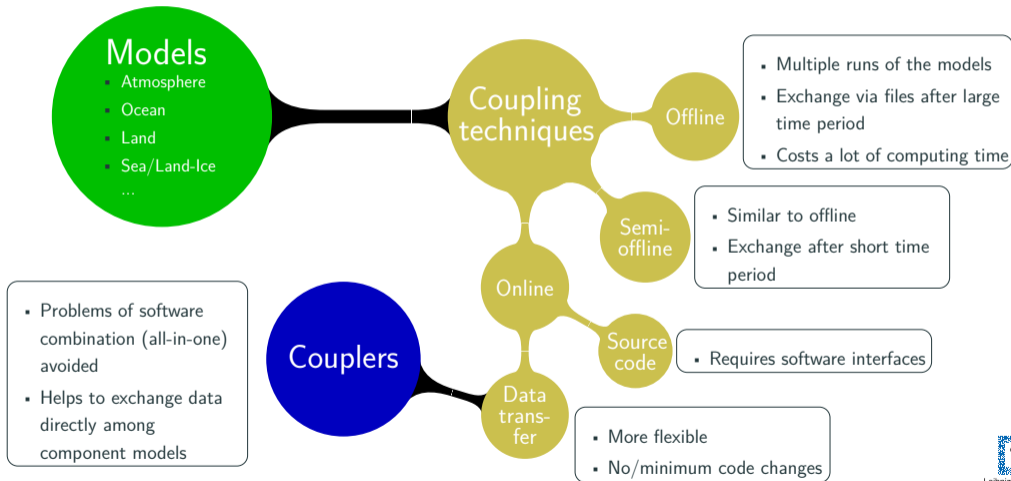
Weather map of Europe on 25th of May 2008 at 6am UTC

- Occasionally weather situation
- High pressure system over southern Scandinavia
- Wind direction mainly northeast

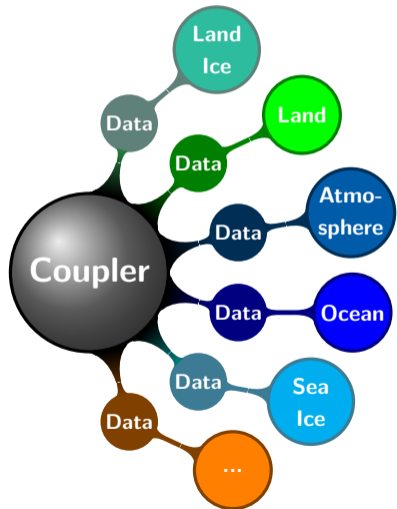
www.wetter3.de (21.09.2016)

(geopotential, relative topography and surface pressure)

Coupling techniques



Online coupling – What are the benefits of a coupler?



- Coupling of additional components to existent models, e.g. atmospheric chemistry, marine biology, carbon cycle etc.
- Developing of components independently from models
- Changing of existing code in the components minimized
- Performing of necessary interpolations
- Supporting of multiple core applications

Couplers: ESMF, MCT, OASIS, YAC

Coupled models for the Baltic Sea or coastal upwelling

Model	Atmosphere	Ocean	Reference
HIRLAM/BOBA-PROBE	HIRLAM	BOBA-PROBE	Gustafsson et al., 1998
REMO/BSMO	REMO	BSMO	Hagedorn et al., 2000
RCAO	RCA2	RCO	Döscher et al., 2002
BALTIMOS	REMO	BSIOM	Lehmann et al., 2004
COAMPS/ROMS	COAMPS	ROMS	Perlin et al., 2007
COSTRICE	COSMO-CLM	TRIMNP	Ho et al., 2012
COSMO-CLM/NEMO	COSMO-CLM	NEMO	Van Pham et al., 2014
RCA4_NEMO	RCA4	NEMO	Wang et al., 2015

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COAMPS/ROMS	COAMPS	ROMS	Perlin et al., 2007
COSTRICE	COSMO-CLM	TRIMNP	Ho et al., 2012
COSMO-CLM/NEMO	COSMO-CLM	NEMO	Van Pham et al., 2014
RCA4_NEMO	RCA4	NEMO	Wang et al., 2015

COAMPS/ROMS vs. COSMO-CLM/NEMO

	COAMPS/ROMS <small>(Perlin et al., 2007)</small>		COSMO-CLM/NEMO <small>(Van Pham et al., 2014)</small>	
	COAMPS	ROMS	COSMO-CLM	NEMO
Coupler	MCT		OASIS3	
Equation	Non-hydrostatic, compressible	Hydrostatic, free-surface	Non-hydrostatic, compressible	Hydrostatic, free-surface
Horizontal resolution	50x20 1-km by 1-km grid boxes		50 km	3 km
Vertical layers	47	40	40	56
Main achievement	Modelling of wind-driven upwelling system along the coast of Oregon		Investigation of 2 m temperature biases between observed data and (un-)coupled results	

Coupler: ESMF – Earth System Modeling Framework

- Suite of software tools for developing high-performance, multicomponent Earth science modeling applications
- Components: **atmosphere**, **ocean**, terrestrial or other physical domains and constituent processes (dynamical, chemical, biological etc.)
- Set of simple, consistent component interfaces – applicable even to couplers themselves
- Variety of data structures for transferring data between components, libraries for regridding/interpolation, time advancement and other common modeling functions

Hill et al., 2004